

ORIGINAL ARTICLE

Proximal to distal grain-size distribution of basin-floor lobes: A study from the Battfjellet Formation, Central Tertiary Basin, Svalbard

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Abstract

The grain-size distribution of sediment particles is an important aspect of the architecture of submarine fans and lobes. It governs depositional sand quality and reflects distribution of particulate organic carbon and pollutants. Documenting the grain-size distribution of these deep-marine sedimentary bodies can also offer us an insight into the flows that deposited them. Submarine lobes are commonly assumed to linearly fine from an apex, meaning there should be a proportional relationship between grain size and distance from the lobe apex. However, not much detailed quantitative work has been done to test this hypothesis. Exposure of a 5 km long dip-section of basin-floor lobes in Clinoform 12, Battfjellet Formation, Spitsbergen, enable the study of basinward grain-size evolution in lobe deposits. Furthermore, the dataset allows testing if there are any documentable grain-size differences between lobe sub-environments. For this purpose, the palaeogeography of Clinoform 12 was reconstructed and the youngest lobe, which was exposed in all collected logs, chosen to be evaluated for its grain-size trends. Photographed thin sections of 66 rock samples were analysed to obtain quantitative grain-size distributions. The results show that fining of lobe deposits occurs predominantly in the most proximal and most distal parts of the lobe, while the intermediate lobe, which is dominated by lobe off-axis deposits, is characterised by a relatively consistent grain-size range. Lobe sub-environments show statistically distinct grain-size distributions from lobe axis to lobe fringe. An explanation for these trends is the interplay of capacity and competence-driven deposition with the grain-size stratification of the flows. The outcomes of this study help to better understand the proximal to distal evolution of turbidity currents and their depositional patterns. They also provide important insights in reservoir potential of basin-floor fans at lobe scale.

KEYWORDS

flow processes, grain-size counting, palaeo-reconstruction, thin sections, turbidites

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1 | INTRODUCTION

Grain-size distribution is one of the fundamental attributes of facies properties. It reflects transport history and depositional mechanisms of sedimentary deposits. Fine-grained turbidite systems are characterised by larger depositional systems, longer run-out distances and larger flow sizes compared to coarse-grained systems (Bouma, 2000; Reading & Richards, 1994). Grain-size distribution of deep-marine bodies not only govern deposit characteristics like porosity and permeability (Porten et al., 2016), but is also thought to be associated with the distribution of organic carbon (Burdige, 2005; Hedges et al., 1997) and pollutants such as plastics (Kane & Clare, 2019; Pohl et al., 2020). However, the processes that control the grain-size distribution across submarine fans are poorly understood. The root of this poor understanding seems to stem at least in part from the paucity of documentation of quantitative grain-size distribution in natural submarine fan deposits (Bell et al., 2018; Fildani et al., 2018; Kane et al., 2017).

Recent studies (de Leeuw et al., 2018a, 2018b; Jobe et al., 2017; Paolo & Martin, 2012; Straub et al., 2012) show that grain sizes change continuously over deep-marine turbidite systems through partitioning along channel-levee complexes. This grain-size partitioning governs the range of grain sizes that is transferred to basin-floor lobe complexes (de Leeuw et al., 2018a, 2018b). In general, studies suggest that the largest amount of sand within a shelf-slope-basin floor system is located within the basin-floor deposits (Bell et al., 2018; Hubbard et al., 2014; van der Merwe et al., 2014; Prather et al., 2017), and that lobe deposits are also relatively coarse compared to the correlative channel-levee deposits (de Leeuw et al., 2018a, 2018b). The fact that basin-floor deposits are coarser grained than their coeval channel-levee deposits is explained by the stratification of flow where fining upwards within the flow hampers the loss of coarser material from the channel through overspill while increasing the loss of finer grains.

While we understand the relationship between lobes and the rest of the deep-marine system we know very little about the grain-size distribution within the lobes themselves. Sylvester and Lowe's (2004) study of vertical fining patterns within turbidites and hybrid beds was one of the first detailed studies of grain-size distribution within lobe deposits and the processes that formed them. More recently, Fildani et al. (2018) looked into detailed vertical grain-size distributions within different bed types attributed to turbidites, hybrid beds and debrites using laser particle analysis. Their study suggests that turbidite deposits in lobes contain more silty material than estimated from visual core description alone, underlining the need for integration of macroscopic facies analysis with microscopic analysis to

inform more comprehensive depositional models. Bell et al. (2018) looked into longitudinal trends of grain-size distribution within lobes from the Gerbe system (Aínsa Basin, Spain) in more detail. They observed that lobe sub-environments (lobe axis, lobe off-axis, lobe fringe and lobe distal fringe; sensu Prélat et al., 2009) show a longitudinal decrease in mean grain size from lobe axis to fringe, but could not report a specific trend. Instead, they documented that for the Gerbe system down-dip decrease in grain size between lobe axis and off-axis is minimal, whereas a more distinct decrease can be documented down-dip between lobe off-axis and fringe deposits.

The aim of this integrated outcrop and grain-size study is to assess the longitudinal trend of grain-size distribution from inferred axial to fringe lobe deposits. For this purpose, quartz-dominated sandstones preserved in the basin-floor segment of Clinotherm 12 (sensu Steel & Olsen, 2002) of the Battfjellet Formation (CTB; Central Tertiary Basin, Svalbard) at Storvola and Hyrnestabben, were used. These were previously interpreted as basin-floor fan deposits (Crabaugh & Steel, 2004; Johannessen & Steel, 2005; Steel & Olsen, 2002). The specific research objectives are: (a) detailed reconstruction of facies, architecture and hierarchy of the exhumed lobe deposits, (b) analysis of grain-size differences between lobe sub-environments, and (c) assessment of the longitudinal trend of grain-size distribution of the youngest exposed lobe. The documented spatial trends in grain size within and between the lobe sub-environments provide a basis to discuss flow processes of turbidity currents flowing over and depositing lobes. Gained insights could be used in future process modelling of lobe deposits.

2 | GEOLOGICAL SETTING

The CTB is located on south-west Spitsbergen, the largest island of the Svalbard archipelago, Arctic Norway (Figure 1A). The basin formed during the Palaeocene and early Eocene contemporaneous with the formation of the West Spitsbergen Fold and Thrust Belt (WSFTB), which is located along the western edge of Spitsbergen and closely linked to the strike-slip opening of the Norwegian-Greenland Sea and later sea floor spreading (Helland-Hansen, 1990; Helland-Hansen & Grundvåg, 2021; Kellogg, 1975; Lundin & Dore, 2002; Müller & Spielhagen, 1990; Steel et al., 1985). Dextral shearing and subsequent transpression/compression along the De Geer Shear Zone caused by the northward movement of Greenland resulted in crustal shortening and thrusting-up of the WSFTB (Leever et al., 2011; Müller & Spielhagen, 1990; Myhre et al., 1982). The initial phase possibly started in the Late Cretaceous or early

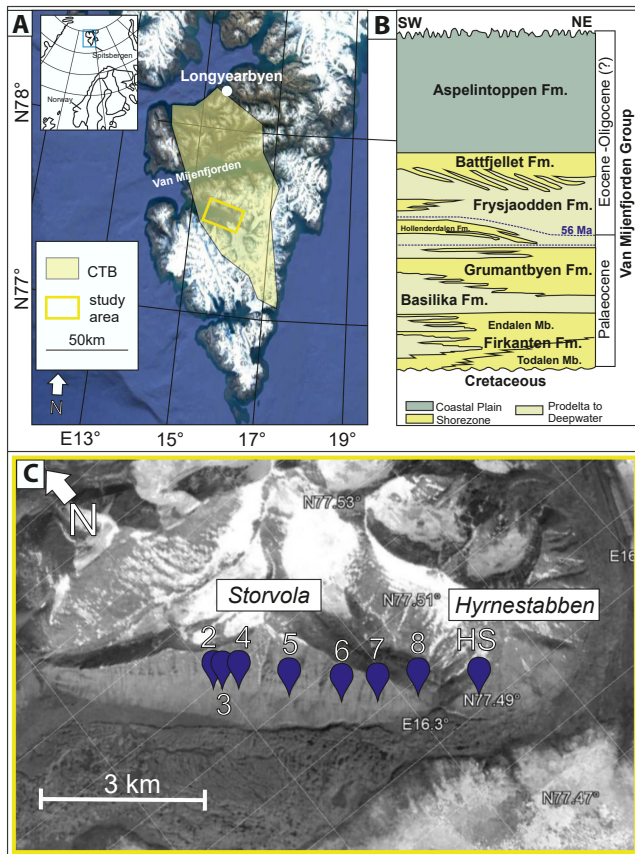


FIGURE 1 (A) Location of the Central Tertiary Basin (CTB). The square indicates the location of the study area in the Van Keulenfjorden. (B) Stratigraphy of the CTB (modified from Crabaugh & Steel, 2004). An Eocene to Oligocene age is interpreted for the Battfjellet Formation. (C) Location of outcrop sections along Storvola and Hyrnestabben used in the study. Satellite images are taken from Google Earth

Palaeocene with oblique contraction and possibly reversion of N–S normal faults in eastern Spitsbergen (Haremo et al., 1993) and was followed by the main phase of ENE-oriented contraction during latest Palaeocene–Eocene (Bergh et al., 1997; Braathen & Bergh, 1995; Bruhn & Steel, 2003; Steel et al., 1985). Some authors have linked the WSFTB to the Eurekan fold belts of North Greenland and Arctic Canada (Piepjohn et al., 2016). The NNW–SSE oriented CTB is located east of the orogenic belt (i.e. the WSFTB) and is identified as a foreland basin despite its connection to the bordering transpressive orogeny (Bruhn & Steel, 2003; Helland-Hansen, 1990; Helland-Hansen & Grundvåg, 2021).

The CTB comprises the *ca.* 2.3 km thick Van Mijenfjorden Group, which was deposited from the early Palaeocene to the Eocene and possibly early Oligocene (Figure 1B). Sediments of the Palaeocene Firkanten, Basilika and Grumantbyen formations were delivered from the east, before a regional drainage reversal occurred at the transition from the late Palaeocene to the early

Eocene (Petersen et al., 2016; Steel et al., 1981). Subsequent formations (the Hollendardalen, Frysaodden, Battfjellet and Aspelintoppen formations) were fed from the emerging orogenic wedge located to the west (Helland-Hansen, 1990; Petersen et al., 2016; Steel et al., 1981, 1985). This study will focus on the slope to basin-floor sediments of the Eocene Battfjellet Formation.

2.1 | Battfjellet Formation and Clinothem 12

The Eocene Battfjellet Formation comprises a series of clinothems that migrated eastwards into the foreland basin (Helland-Hansen, 1990, 1992; Helland-Hansen & Grundvåg, 2021; Steel et al., 1985). Present dip angles for the associated clinofolds are in the range of 1–4°, and have not been decreased by more than 1° by post-depositional compaction (Helland-Hansen, 1992; Plink-Björklund et al., 2001). Their amplitudes are typically reported to be in the range of 100–250 m (Helland-Hansen, 1992; Johannessen & Steel, 2005), or more rarely up to 350 m (Plink-Björklund et al., 2001).

The clinothems of the Battfjellet Formation consist of shelf to slope and basin-floor deposits (Helland-Hansen, 1990; Mellere et al., 2002; Steel & Olsen, 2002). Sand-prone sediments were trapped on the shelf-edge and upper slope, forming basinward-thinning wedges, or on the shelf as stacked shoreface to deltaic parasequences. (Grundvåg et al., 2014a; Johannessen & Steel, 2005; Plink-Björklund et al., 2001). However, some clinothems show evidence of sand transport to the basin floor (Crabaugh & Steel, 2004; Johannessen & Steel, 2005; Plink-Björklund et al., 2001; Steel & Olsen, 2002). Clinothem 12 is one such clinothem and is the focus of this study. The study area is part of the renowned Van Keulenfjorden transect on the south-western flanks of Storvola and Hyrnestabben, where Clinothem 12 is interpreted as representing slope and basin floor fan deposits (Clark & Steel, 2006; Crabaugh & Steel, 2004). The coeval shelf-delta deposits occur further to the west on the neighbouring mountain of Brogniartfjella (Johannessen & Steel, 2005; Steel & Olsen, 2002).

3 | FIELD METHODS

Eight stratigraphic logs, totalling 140 m in thickness, were collected along a 4.7 km transect of the exposed basin-floor fan of Clinothem 12 (Figure 1C). The logs record lithofacies, grain size (based on field observations), bed thickness, sedimentary structures, the presence of organic matter (i.e. plant debris), bounding surfaces and palaeocurrents from

flute marks and current ripples. The stratigraphic logs comprise the exposed vertical succession of the sand-prone basin-floor fan at the outcrop locations. The heterolithic lower part of the fan is commonly weathered and covered by scree. Logs were correlated by assigning depositional sub-environments based on their lithofacies association and bounding surfaces (Prélat et al., 2009). A grain-size sample was collected every 1–2 m vertically within the individual logs near the base of the sampled beds (at about 2–3 cm from the base) to avoid biasing towards specific bed divisions or sedimentary structures. Sampling just above the

bases of beds may under or overestimate the grain size of the entire bed as massive turbiditic sands can show inverse grading, normal grading or a lack of grading (Sylvester & Lowe, 2004). However, most often the bed bases give a good representation of the more energetic part of the flow and the consistency is needed to compare lobe sub-environments over longer distances. Within the investigated succession, one lobe with a complete proximal to distal preservation could be recognised and only those samples (66) that are located within that lobe were selected for evaluation of the longitudinal grain-size distribution (Figures 2 and 3).

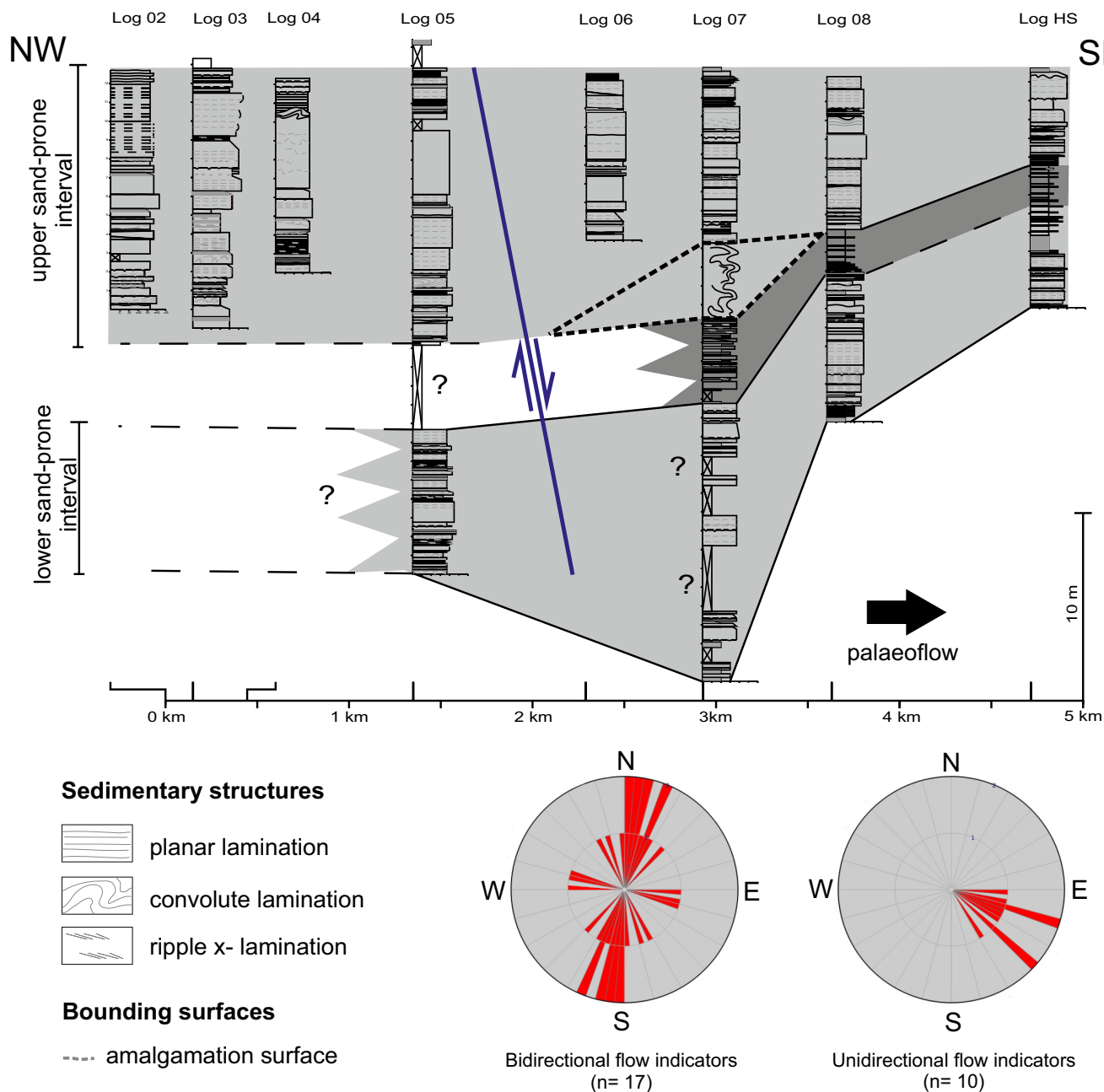


FIGURE 2 Correlation panel of Clinotherm 12. Uncertain correlations are pointed out with dashed black lines. The normal fault between log 5 and log 6 is marked by a dark blue line. The slumped unit of the upper lobe/lobe 2 is marked by a stippled black polygon. The indicated fault between log 5 and log 6 is interpreted as post-depositional and serves orientation purposes only

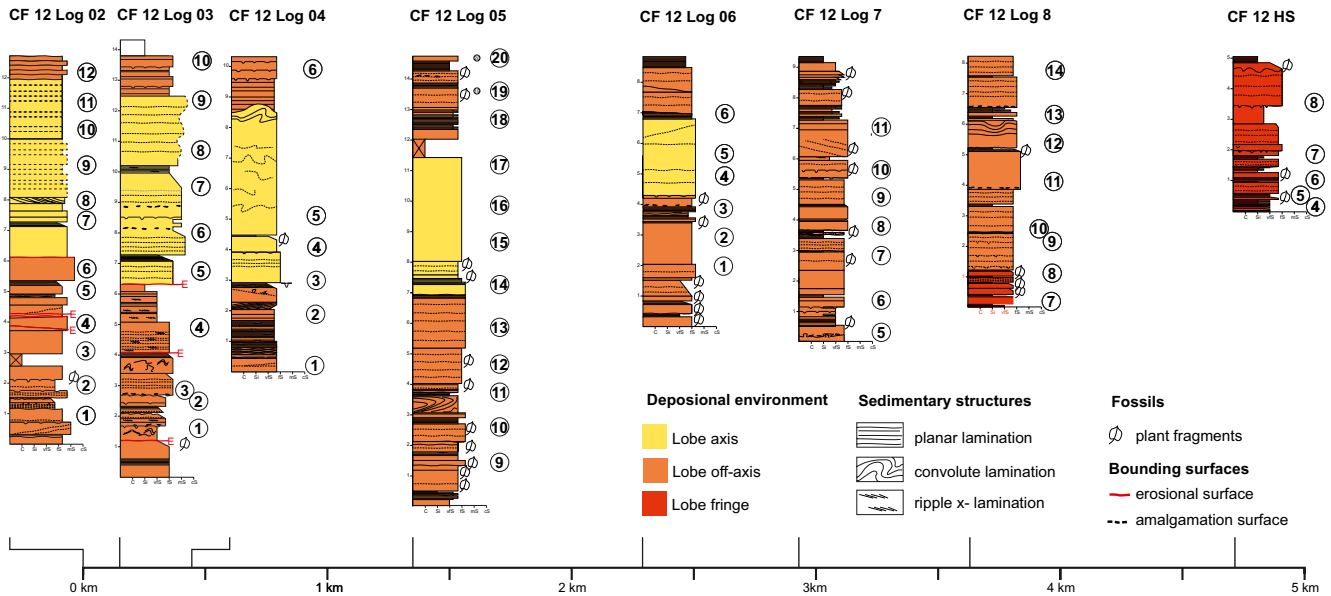


FIGURE 3 Locations of samples and sample number taken within the deposits of the youngest lobe. Colours indicate the interpreted lobe sub-environment from field data alone

3.1 | Depositional architecture of the basin floor fan

3.1.1 | Thickness, lithofacies and palaeoflow directions

Cliniothem 12 is formed by two sand-prone intervals that are separated by a thin-bedded heterolithic package and silt-prone package. Both sand-prone packages have an approximately 15 m maximum thickness. The separating heterolithic/siltstone-prone interval has a maximum thickness of 4.5 m. The maximum exposed total thickness of 33 m is measured in log 7 (Figure 2). Poor exposure up-dip of log 7 prevents assessment of the total thickness in logs 2–6. In addition, the lowermost, heterolithic part of the sections is commonly scree covered, hindering accurate thickness measurements of the distal lobe fringe facies. A thinning of the cliniothem can be observed distally of Log 7. The total thickness of the deposits thin from approximately 33 m (log 7) to 13 m at Hyrnestabben (estimated thinning rate of 11 m/km). In the intermediate part of the studied section, a 4 m thick folded heterolithic sandstone unit can be observed at the base of the upper sand package. A detachment zone to its base suggests that part of the underlying succession may have been eroded as the deposits were remobilised.

A limited set of palaeoflow directions ($n = 27$) measured from flute marks and current ripples of Cliniothem 12 deposits resulted in two different directions: 195 and 118°, respectively (Figure 2). Flute marks account for the southern palaeoflow indicators, while palaeoflow measurements taken from ripple laminations point to the

south-east. Crabaugh and Steel (2004) reported east to south-east directed flute casts for Clinioform 12.

A bimodal flow direction in confined basins is probably caused by deflection and reflection of flow from side slopes and topographic highs (Haughton, 1994; Kneller et al., 1991; Kneller & McCaffrey, 1999; Pohl & McCann, 2014; Spotts, 1964). Crabaugh and Steel (2004) suggested flow diversion within the basin-floor fan of Clinioform 14 which changed from basinward to slope parallel at the most distal part and proposed that tectonically induced topography influenced fan deposition. They propose that the basin-floor fan of Cliniothem 12 is not affected by such tectonic activity during sedimentation. However, influence of confinement on flow directions cannot be entirely discarded.

The main palaeoflow direction assumed from the measurements in this study is to the south-east and conforms to basinward growth of Cliniothem 12 towards the south-east reported by Crabaugh and Steel (2004). The positioning of individual lobes forming the basin-floor fan is probably more variable.

3.1.2 | Lithofacies

Lithofacies encountered in the outcrop (Figure 4) are summarised in Table 1. Four depositional sub-environments were interpreted based on characteristic lithofacies associations: lobe axis, lobe off-axis, lobe fringe and lobe distal fringe (sensu Etienne et al., 2012; Grundvåg et al., 2014b; Prélat & Hodgson, 2013; Marchand et al., 2015; Masalimova et al., 2016; Prélat et al., 2009; Sychala et al., 2015, 2017a) (Figure 5).

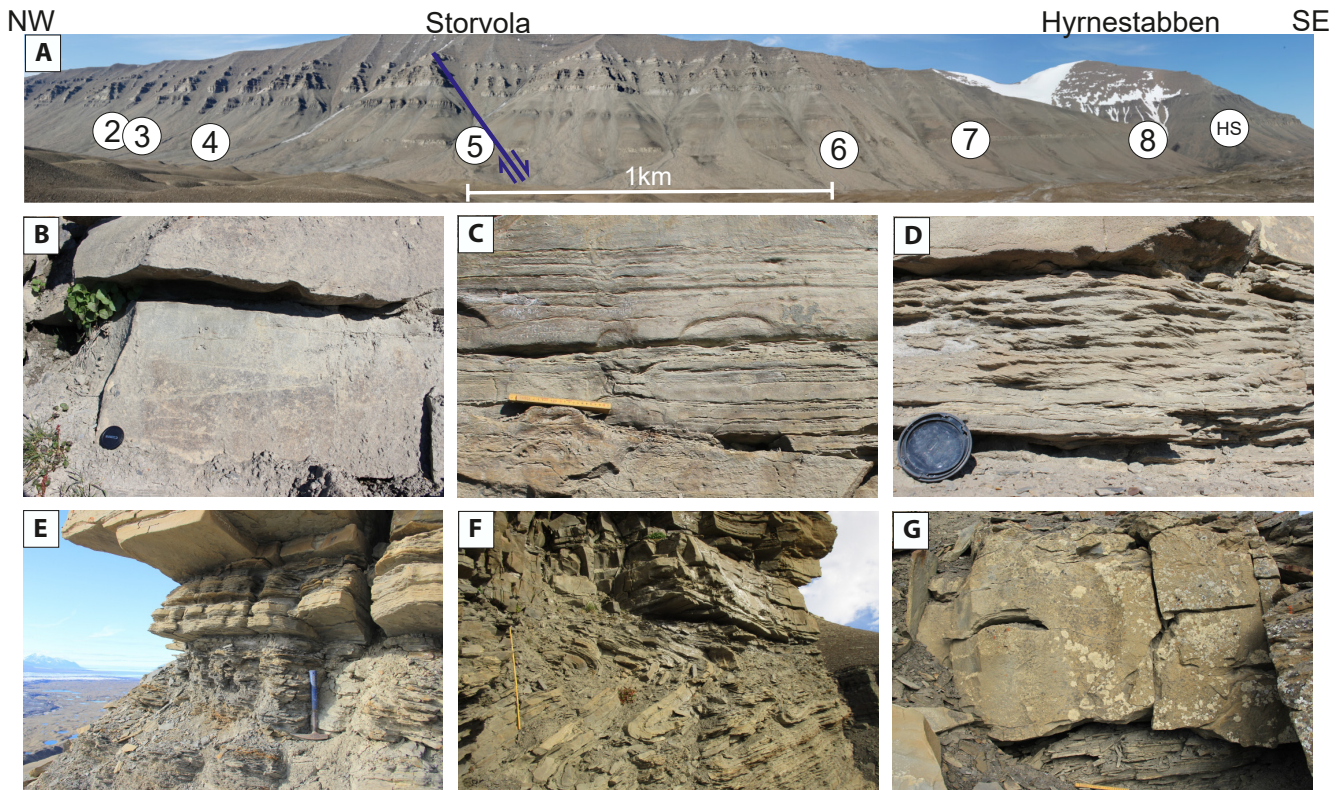


FIGURE 4 (A) Photopanel of the study area with Storvola in the foreground and Hyrnestabben in the background to the right. Log locations are indicated with their numbers. Indicated normal fault is post-depositional and for orientation purposes only. (B) Example of F1: structureless sandstone. Lens cover as scale (*ca* 7 cm diameter). (C) Example of F2: planar laminated sandstone. Folded rule (20 cm) as scale. (D) Example of F3: Structured sandstone. Here the deposit shows ripple laminations. Lens cover as scale (*ca* 7 cm diameter). (E) The log section at Hyrnestabben (log HS) shows F6 (siltstones) at its base which are overlain by F4 (heterolithic interval). Geological hammer (30 cm) as scale. (F) Folded sandstones at the base of the youngest lobe in the section of log 7. Rule (100 cm) as scale. (G) Hybrid bed (F5) at log HS. Note the irregular transition from clean sandstone to the debritic division which is weathered back into the slope. Folded rule (20 cm) as scale

Lobe axis

Lobe axis deposits are comprised of thick-bedded (0.75–1.2 m) structureless sandstone (F1, [Figure 4B](#)), thin-bedded to medium bedded (0.1–0.4 m) highly amalgamated sandstone (F1), planar laminated and stepped laminated sandstone (F2, [Figure 4C](#)). The lobe axis association is characterised by >95% sandstone. Bed bases locally cut down up to 15 cm into underlying beds. Siltstones (F6, [Figure 4E](#)) are rarely observed. Amalgamation is very common and results in sandstone packages to up to 3.4 m thick. Deposits of the lobe axis are mostly tabular, but can also show low-relief pinching and swelling due to minor compensational stacking of beds.

Lobe off-axis

Lobe off-axis deposits are dominated by planar laminated sandstone (F2, [Figure 4C](#)), structured sandstone (F3, [Figure 4D](#)) and subordinated structureless sandstone (F1, [Figure 4B](#)). Siltstone deposits (F6, [Figure 4E](#)) separating individual sand beds are more common than in

the lobe axis. This facies association is characterised by 82%–100% sandstone. Deposits are mostly thin-bedded to medium bedded (0.1–0.5 m), but occasionally thicker beds are present. Beds are tabular and can be traced for several tens of metres across the extent of the studied outcrop segments.

Lobe fringe

Lobe fringe deposits comprise planar laminated sandstone (F2, [Figure 4C](#)), structured sandstone (F3, [Figure 4D](#)), heterolithic packages (F4, [Figure 4E](#)), hybrid beds (F5, [Figure 4G](#)) and siltstone (F6, [Figure 4E](#)). From all the facies associations, the lobe fringe facies association is the most variable in facies and bed thicknesses, which range from thin-bedded to thick-bedded (0.05–1.1 m). Lobe fringe deposits are characterised by <70% sandstone. The transition of lobe fringe to distal lobe fringe is marked by the change from variable thickness packages to thin-bedded intervals with very low sand percentages (<20%).

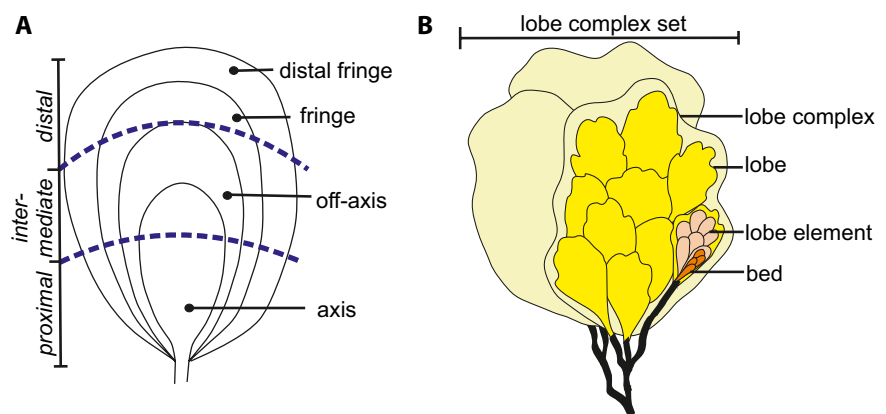
TABLE 1 Observed lithofacies in Clinotherm 12, their processes and depositional environment

Lithofacies	Grain size	Thickness range	Description	Process interpretation	Depositional environment
Structureless sandstone (F1)	fs to ms	0.25–1 m	Ungraded to graded; top can show (a) a silty cap, and/or (b) coal fragments (occurrence increases from CF12-3 to CF12-7); bases are sharp or loaded; often groove and flute casts on base; outsized clasts of up to 3 mm; dewatering structures; mudclasts (1–2 cm) at discrete horizons; up to 3.4 m thick in highly amalgamated intervals; 10–15 cm of erosion can occur at the base	Deposited by high-density turbidity currents (Kneller & Branney, 1995) with high aggradation rates (Arnott & Hand, 1989; Leclair & Arnott, 2005; Talling et al., 2012)	Lobe axis
Planar laminated sandstone (F2)	fs to ms	0.25–0.5 m	Planar laminations are either flat or slightly inclined; Planar laminations can be at the top of F1 beds or throughout the bed; Stepped planar laminations are observed from the base of beds	Deposited by high-density turbidity currents. Planar laminations are produced by reworking of the bed (Allen, 1982; Best & Bridge, 1992; Southard, 1991), Stepped planar laminations are produced by rapid sediment fallout (Hiscott, 1994; Sumner et al., 2008)	Lobe axis or lobe off-axis
Structured sandstone (F3)	vfs to fs	0.05–0.25 m	Wavy to sinusoidal lamination; ripple cross-lamination; isolated ripples; climbing-ripple lamination; stratified laminations; flame structures; horizontal bioturbation; bed bases and tops are commonly sharp and flat	Deposited by low-density turbidity currents. Wavy or sinusoidal lamination indicates deposition from waning currents with very high rates of suspension fallout (Allen, 1973; Jopling & Walker, 1968; Hunter, 1977). Current-ripple lamination produced by reworking through dilute flows along the bed (Allen, 1982; Best & Bridge 1992; Southard, 1991). Climbing ripple lamination forms under bedload transport associated with high aggradation rates (Allen, 1973; Hunter, 1977; Jobe et al., 2012)	Lobe off-axis
Heterolithic package (F4)	Silt and vfs	0.2–0.5 m intervals	(a) Lenticular sandstone bodies in siltstone, (b) intercalated tabular vfs and siltstone, or (c) dominantly siltstone (70–90%) with minor sandy beds (2–5 cm thick); sandstone beds can show planar, ripple or wavy lamination	Deposits of distal, sluggish, lowvolume flows (cf. Jobe et al., 2012). Ripple lamination forms beneath dilute turbulent flows via reworking of the bed under moderate aggradation rates, whereas climbing-ripple lamination forms under high aggradation rates (Allen, 1971, 1982; Southard, 1991)	Lobe fringe
Bipartite beds (F5)	csilt to fs	0.15–1.2 m	Lower sandstone division, usually structureless, rip-up clasts (up to 30 cm) on discrete horizons; upper muddier matrix with abundance of mudchips and coal fragments; poorly sorted; can be slightly graded at their tops	Co-genetic turbidity currents (lower division) and cohesive debris flows (upper divisions)—hybrid beds <i>sensu</i> Haughton et al. (2003), Talling et al. (2004), Haughton et al. (2009) and Hodgson (2009)	Lobe fringe

TABLE 1 (Continued)

Lithofacies	Grain size	Thickness range	Description	Process interpretation	Depositional environment
Siltstone (F6)	fsilt to csilt	0.01–0.02 m	Planar bedded siltstone; mostly structureless; some beds show faint lamination	Deposited by dilute turbidity currents. Planar lamination is a product of traction (Mutti, 1992; Stow & Piper, 1984; Talling et al. 2012). Structureless beds are formed by direct suspension fallout (Bouma, 1962)	Distal lobe fringe to background sedimentation
Folded sandstone (F7)	vfs to fs	0.55–4 m intervals	Folded and overturned beds of 5–15 cm thickness; some ripple laminated beds, but mostly structureless and normal graded; Bases of chaotic and folded units are sharp to erosive, while bed tops are sharp	Syn-depositional deformation due to remobilisation of local thin-bedded stratigraphy. The low amount of disaggregation supports an interpretation of slump deposits	Not indicative of a specific environment

FIGURE 5 (A) Simplified model indicating the various sub-environments in a lobe (redrawn from Prélat et al., 2009). Stippled blue lines indicate proximal, intermediate and distal lobe areas. (B) Plan-form view of fivefold lobe hierarchy: bed to bed set, lobe element, lobe, lobe complex and lobe complex set (modified from Prélat et al., 2009, redrawn from Sychala et al., 2017a)



Distal lobe fringe

Distal lobe fringe deposits are dominated by structureless and planar laminated siltstones (F6, Figure 4E). Thin-bedded very fine sandstone beds can be observed to be intercalated within the silt-prone intervals. Distal lobe fringe deposit can form intervals up to 4.5 m thick.

Geographically the intermediate lobe is neither close to the feeder channel nor to the lobe pinch-out. Finally, the distal lobe is characterised by deposits from the lobe fringe and lobe distal fringe only. It is located down-dip and marked by its association with the down-dip pinch-out of the sandy lobe deposits.

3.1.3 | Proximal and distal lobes

The terms 'proximal', 'intermediate' and 'distal' lobe are descriptive and reference spatial areas of the deposits rather than indicating a facies-based sub-environment. These terms are not officially defined and are therefore not used consistently by different authors. Figure 5A shows the definition of these terms for this study. The proximal lobe is located up-dip, close to the feeder channel and is dominated by lobe axis deposits, but can also encompass lobe off-axis and lobe fringe deposits laterally. The intermediate lobe comprises dominantly lobe off-axis deposits, but may also comprise deposits from all other lobe sub-environments.

3.1.4 | Hierarchy of Clinothem 12

Lobe deposits commonly exhibit a fivefold hierarchy (Deptuck, 2008; Grundvåg et al., 2014b; Prélat et al., 2009; Sychala et al., 2017a): (a) a 'bed' represents a single depositional event; (b) one or more beds stack to form a 'lobe element'; (c) several lobe elements that are divided by thin siltstone intervals form a 'lobe'; (d) one or more genetically related lobes stack to form a 'lobe complex'; and (e) one or more lobe complexes within the same lowstand systems track stack to form lobe complex sets (Figure 5B). The studied section of Clinothem 12 conforms to the hierarchical level of lobe complex. The lower and upper

sandstone-prone intervals conform to two lobes which are separated by lobe fringe deposits (thin-bedded heterolithic package), suggesting that there were genetically related sandstone-prone lobe deposits lateral to the exposed study area. Occurrence of hybrid beds in the most distal logged section at Hyrnestabben (log HS; Figure 1C) points towards the youngest lobe being exposed in a near dip-section as frontal lobes are more likely to exhibit hybrid bed deposits (Spychala et al., 2017a, 2017b). However, it cannot be excluded that the exposure may be slightly oblique to the main palaeocurrent direction.

3.2 | Grain-size evaluation

3.2.1 | Procedure

Thin sections with a thickness of 30 μm were produced from 66 samples (one per sample) and these thin sections were photographed with a high-resolution camera under a Leica microscope. Samples were counted by grid counting manually outlined grains using image processing software (ImageJ) which automatically records the long and short-axis and orientation of the long-axis based on the best fitting ellipse of the grains. The grid size was kept at 0.03 mm^2 . Per sample 300 grains (19,800 grains in total) were measured to provide sufficient accuracy even for very poorly sorted sandstones (Johnson, 1994). To correct for underestimation of the spherical-grain mean values due to cutting effects, the recorded long and short-axis are multiplied by 1.1318 (Johnson, 1994). The true nominal grain diameter (the diameter of a sphere with the same volume as the considered particle) was calculated following Johnson (1994):

$$D' = d' + 0.4(a' - d')^2 \quad (1)$$

where D' is the true nominal diameter (mm), and $d' = (a' b')^{1/2}$, and a' and b' are the corrected long and short-axes (mm), respectively. True nominal grain diameters were converted to ϕ -scale ($\phi = -\log_2(D/\text{mm})$; Krumbein, 1934) to give the most satisfactory representation of the distribution (Johnson, 1994). Determining grain size from photographed thin sections causes a lower precision for grains below the thin section thickness (de Leeuw, 2017; Johnson, 1994; Sylvester & Lowe, 2004) as they might overlap each other. A lower limit for grain-size measurement was placed at 30 μm in this study. This limit excludes fine-silt from grain-size measurements. The obtained grain-size distributions in ϕ are approximately normally distributed and suggest a log-normal distribution on the linear scale. Descriptive statistical measures (mean [M], standard deviation [σ], skewness [S], and kurtosis [K]) are calculated using MATLAB.

To test the hypothesis that grain-size distributions of lobe sub-environments are statistically different an analysis of variance (ANOVA) was applied to the mean grain size of the distributions. ANOVA compares the 'variance between/among' grouped data with the 'variance within' a group. For this study the grouping will be based on lobe sub-environments: lobe axis, lobe off-axis and lobe fringe. Although grain size is a factor that plays into the assessment of lobe sub-environments, evaluation of the grain-size distribution between these 'groups' can help us determine if their transitional character is dominant or if they actually form individual clusters.

3.2.2 | Measurement errors

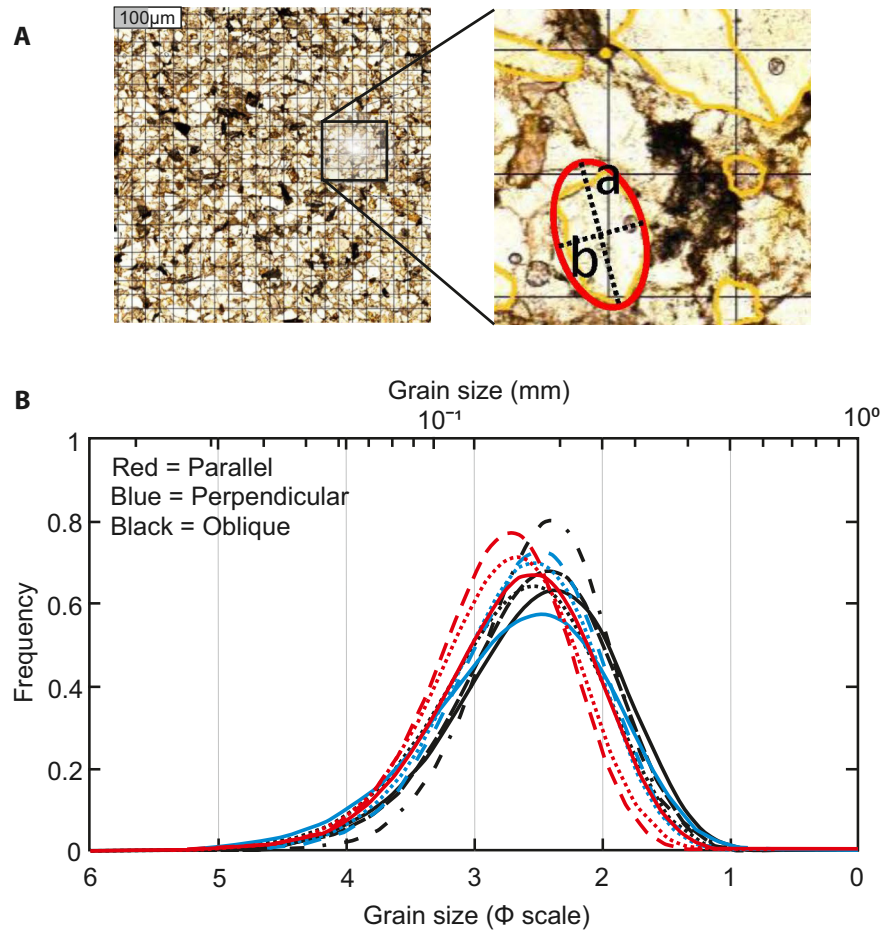
Two types of measurement error were investigated by dissecting a single sample in multiple orientations (orientation error) and counting grain sizes by multiple operators (operator error). The observed differences between groups of samples need to clearly exceed these measurement errors for the results to indicate significant differences.

3.2.3 | Orientation error

Non-spherical grains are frequently arranged in specific patterns with respect to the palaeoflow direction (Arnott & Hand, 1989; Baas et al., 2007; Hiscott & Middleton, 1980). In turbidites, a flow-aligned fabric (with the long axis aligned with the palaeoflow direction) is favoured during relatively rapid sedimentation from suspension, whereas flow-transverse fabric is the result of grain transportation along the bed by traction during relatively low sedimentation from suspension. Oblique fabric appears to be restricted to massive sandstone turbidites but the mechanism causing this fabric is still unclear (Baas et al., 2007).

For consistency, all selected samples (66) were cut perpendicular to the bedding with most samples (51) cut parallel to the assumed main palaeoflow direction of 118°. For some samples the orientation of the samples could not be reconstructed after returning from the field, and for other samples, the limited number of palaeoflow measurements taken from few other beds, sometimes in another stratigraphic section, could be a poor reflection of the true transport direction of that sample. These uncertainties necessitate a careful assessment of the measurement errors related to the orientation of the thin sections relative to the palaeoflow orientation. For this purpose, one sample was selected and cut parallel, perpendicular and oblique (45°) to the assumed main palaeoflow direction of 118°. The data show that the measured grain size indeed varies depending on the orientation of the thin

FIGURE 6 (A) Example of outlined grains processed with ImageJ. First a scale is set to create the accurate number of grid points. Second, grains are marked manually. ImageJ measures the major (a_{mm}) and minor (b_{mm}) axis. (B) Master sample grain-size evaluation: Grain-size distributions of oriented samples with respect to the palaeoflow analysed by three different operators. Each operator counted at least 300 grains. The line styles represent the operators: solid line = Sara Wroblewska; loosely stippled line = Yvonne Sychala; closely stippled line = Thymen Ramaaker



section with respect to 118° , which illustrates the error of cutting the sample in different orientations (Figure 6). The pooled standard deviation ('weighted average standard deviation') for different orientations of the sample is 0.06ϕ .

3.2.4 | Operator error

The three thin sections of the sample were analysed by three different operators to test the effect of operator-choices in outlining individual grains. The pooled standard deviation ('weighted average standard deviation') for the error due to individual operator choices is 0.11ϕ . Based on this analysis, the decision was made to evaluate all samples by a single operator (T.A.B. Ramaaker) to ensure consistent application of grain-delineations and avoid inter-operator errors.

3.3 | Grain-size distribution

Figure 7 illustrates the grain-size distributions of 19 lobe axis samples, 40 lobe off-axis samples, and seven lobe fringe samples. With the exception of a few grains,

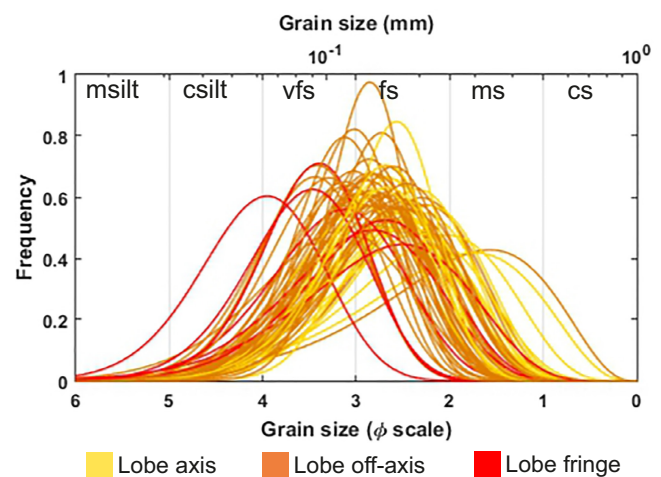


FIGURE 7 Probability density functions of the grain-size distributions of basin-floor fan deposits of Clinotherm 12 showing all 66 grain-size distributions of lobe axis (yellow), off-axis (orange) and fringe (red) samples

measured individual grain sizes range from 0 to 6ϕ (1,000 to $30 \mu\text{m}$; coarse sand to medium silt). The grain-size distributions show mean grain-size values ranging from 2.06 to 4.03ϕ (239 to $61 \mu\text{m}$) with an average mean of 2.89ϕ

(134 μm) (standard deviation of 0.33 ϕ). Fifty-six of the 66 samples have a mean grain size between 2.5 and 3.5 ϕ (177 to 88 μm) with only one sample being finer than 3.5 ϕ (4.03 ϕ). Modal grain size ranges from 1.57 to 3.95 ϕ (336 to 64 μm) and indicates an overall slightly skewed distribution towards finer material (excess of finer material).

Most samples are moderately well sorted (standard deviation of 0.5–0.7 ϕ) or moderately sorted (σ : 0.7–1.0 ϕ). The lobe off-axis samples especially are predominantly moderately well sorted. Grain-size curves of the measured samples are mainly symmetrical (skewness of -0.1 to 0.1) but can be fine skewed (0.1–0.3) indicating a tail of finer grains.

3.4 | Grain-size distribution within sub-environments

The lobe sub-environments have similar distribution curves within each group, but show significant differences between groups. Lobe axis (Figure 8A), lobe off-axis

(Figure 7B) and lobe fringe (Figure 8C,D) have mean grain-size values of 2.68 ϕ (156 μm ; standard deviation [σ (ϕ)] of 0.20), 2.91 ϕ (133 μm ; σ [ϕ]: 0.29), and 3.30 ϕ (101 μm ; σ [ϕ]: 0.43), respectively. The analysis of variance (ANOVA) of the mean grain sizes indicates that all three sub-environments are statistically different at a confidence level of 99% (for detailed information see Table 2). Although the mean grain-size values are distinctly different, there is still overlap between the grain-size distribution curves which means no sub-environment has a discrete grain-size range that separates them from the other sub-environments.

Lobe axis, lobe off-axis as well as lobe fringe have samples that are fine skewed (Figure 9). No relationship between mean grain size and bed thickness (of sampled beds), standard deviation, skewness or kurtosis could be observed to be characteristic for any sub-environment.

Lobe sub-environments were defined based on lithofacies implying a possible difference in grain orientation between lithofacies associations which could affect the

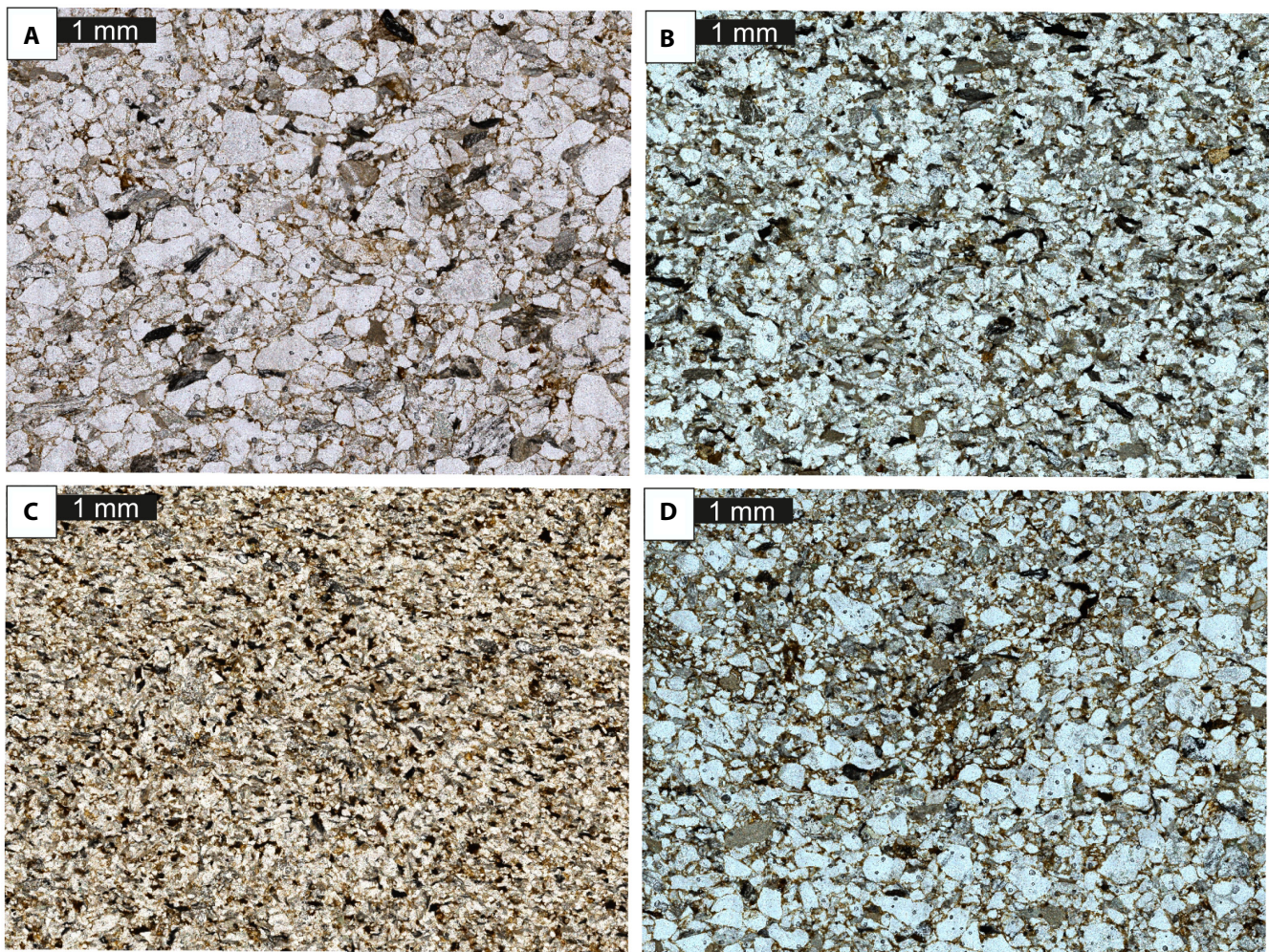


FIGURE 8 Examples of lobe sub-environment lithology and grain size in thin sections. (A) Lobe axis. Sample 6 of log 3. (B) Lobe off-axis. Sample 12 of log 5. (C) Lobe fringe. Sample 4 of log Hs. (D) Lobe fringe hybrid bed. Sample 8 of log HS

TABLE 2 One-way ANOVA table

Source	Sum-of squares (SS)	Degrees of freedom	Mean squares	F-Distribution	Prob>F
Between	2.06	2	1.03	12.69	0.00002
Within	5.12	63	0.08		
Total	7.18	65			

grain-size distribution. However, the observed difference in grain-size distribution between the different cutting orientations (0.06ϕ) is significantly smaller than the difference between lithofacies associations. Furthermore, the range within the sub-environments is significantly larger than the deviation caused by cutting orientation, which indicates that true variability is larger than measurement error both within sub-environments and between sub-environments.

3.4.1 | Longitudinal trends

A regression analyses of the mean grain size of all samples shows a moderate negative correlation ($R = -0.51$) with distance over 5 km (Figure 9). However, samples with mean grain sizes coarser than 2.5ϕ ($177 \mu\text{m}$) are only observed in the most proximal part of the basin-floor fan. Samples with mean grain size between 2.5 and 3.5ϕ (177 to $88 \mu\text{m}$) can be observed throughout the dip-section without a clear decrease in value. The finest mean grain sizes are recorded in the distal fringes of the Hyrnestabben section. Looking at longitudinal grain-size distributions of the interpreted lobe sub-environments separately a more detailed pattern can be seen. Lobe axis settings reach half-way along the exposed lobe section and show a relatively stronger correlation with distance ($R = -0.56$; Figure 10). Lobe off-axis deposits occur along the entire studied transect (more than 3.5 km) with the exception of the most distal log at Hyrnestabben (Figure 1C). Samples from these off-axis deposits only show a weak decrease in mean grain size ($R = -0.23$). For these samples no correlation between distance and any other parameter is found. Lobe fringe deposits collected from log 8 and HS (Figure 2) indicate a decrease in grain size ($R = -0.42$; Figure 10) over a distance of roughly 1 km.

3.4.2 | Vertical trends

Vertical trends through the individual logs are illustrated in Figure 11. No absolute datum is present throughout the section to retrieve absolute vertical positions within the lobes, and each log has a different thickness. Therefore, sampled beds are imaged using equal vertical spacing between beds.

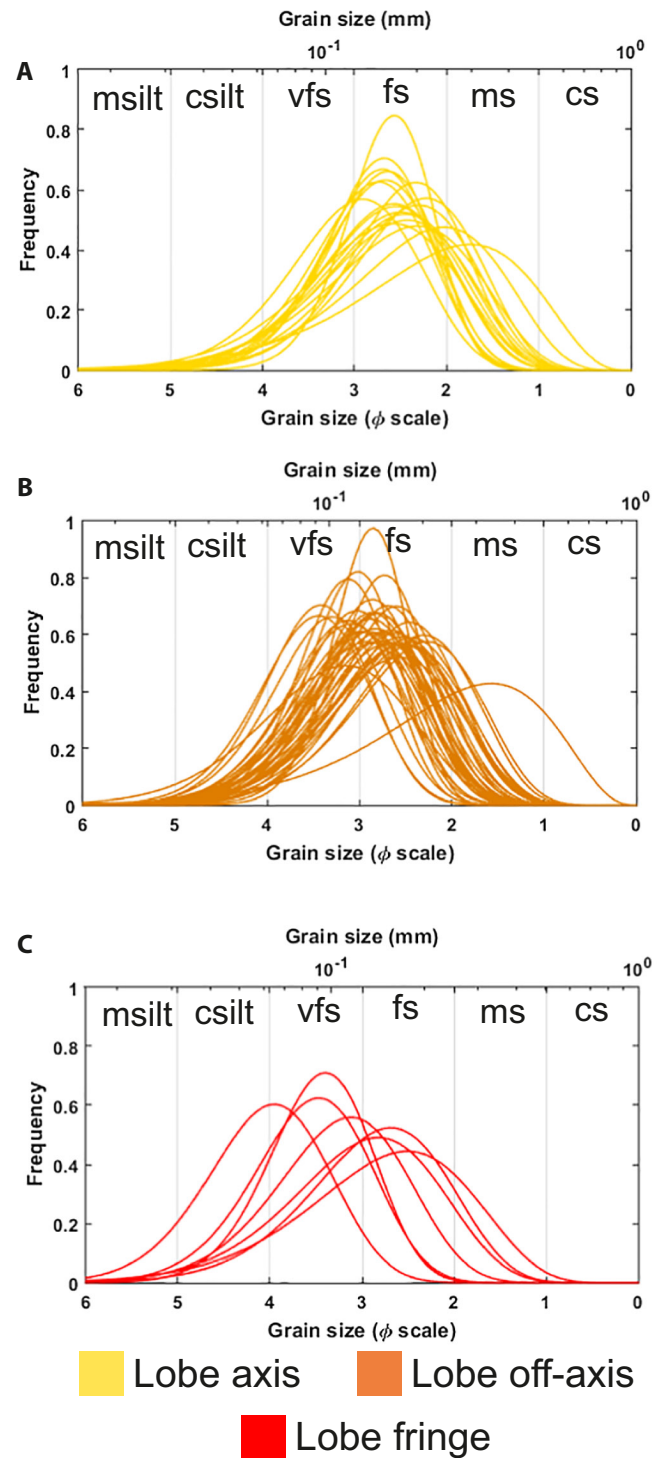


FIGURE 9 Probability density functions of the grain-size distributions of lobe sub-environment deposits of Clinotherm 12. (A) Lobe axis. (B) Lobe off-axis. (C) Lobe fringe

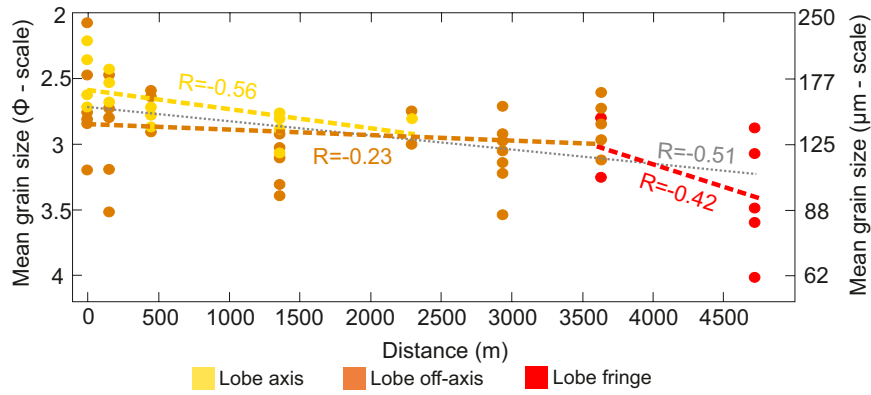


FIGURE 10 Longitudinal development of mean grain size. Black dotted line indicates the linear regression of the mean grain size of all samples with longitudinal distance ($R = -0.51$). Lobe axis, off-axis and fringe samples and their linear regression with distance ($R = -0.56$, $R = -0.23$, and $R = -0.42$, respectively) are indicated by the yellow, orange and red points and dotted lines, respectively. Note the relatively stable range in mean grain size of lobe off-axis samples

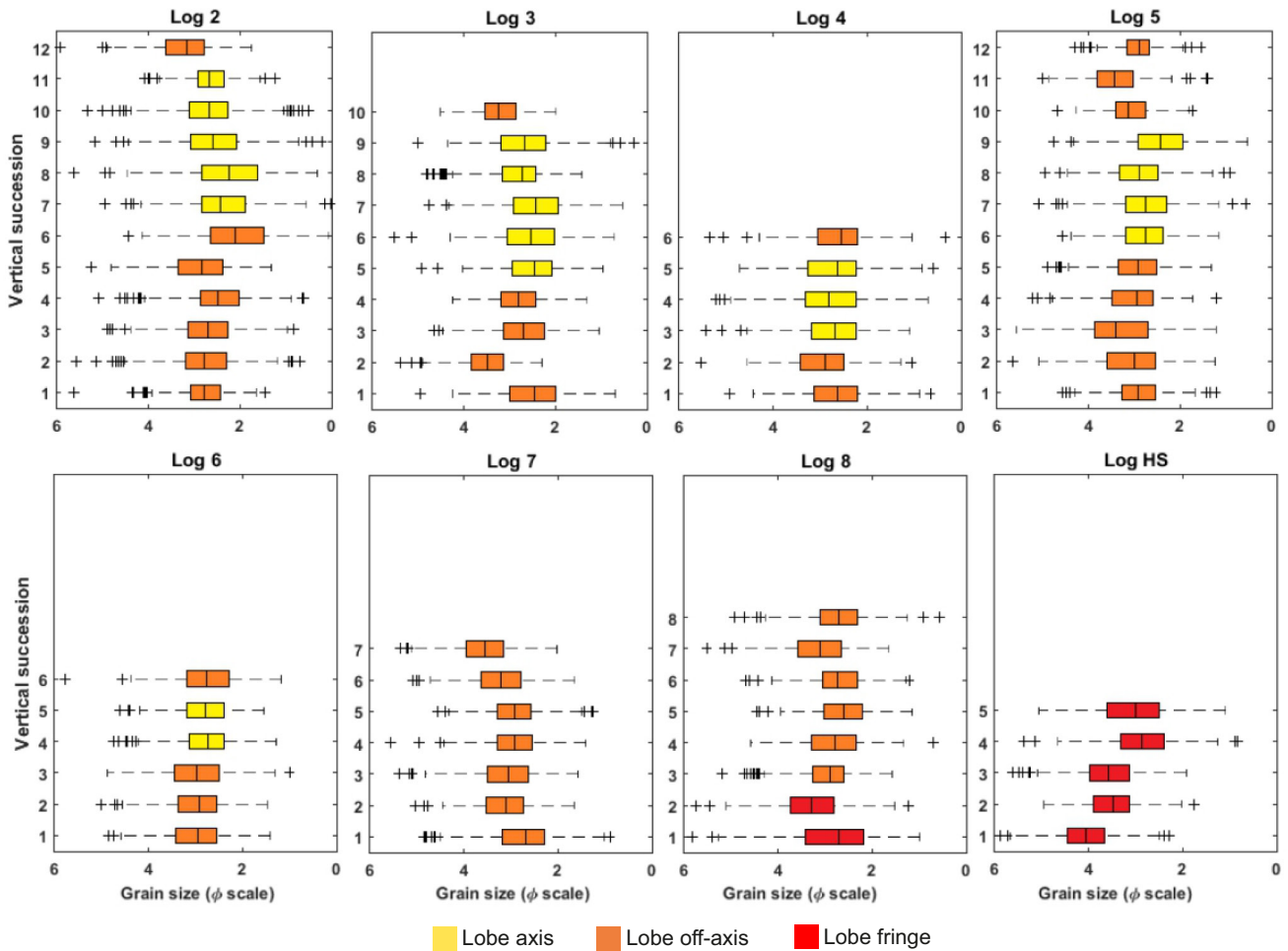


FIGURE 11 Boxplots of grain-size distributions in vertical succession (with numbers on the y-axis referring to the order of the sampled beds from base to top) of individual logs. Colours indicate the sub-environments lobe axis (yellow), off-axis (orange) and fringe (red). The vertical line within the box indicates the median and the boundaries of the box indicate the 25th and 75th percentile. The whiskers, marked as $- -$, extend 1.5 times the interquartile distance (distance between indicate the 25th and 75th percentile). The plotted whiskers extend up to adjacent value that is the most extreme data point within the calculated whisker. The outlier, marked as $+$, are values above or below the whiskers

No obvious overall vertical coarsening or fining trend is observed (Figure 11). Only the most distal location (log HS) that consists solely of lobe fringe deposits has a vertical coarsening trend with mean grain size shifting from 4.03 to 3.06 ϕ (30–119 μm). Other logs appear to show a subtle coarsening to fining trend with the coarsest samples at the middle to upper part of the log. This is clearly observed in log 2 where the lower six samples represent lobe off-axis and show a coarsening trend followed by five lobe axis samples that show an overall fining upwards. The lowermost samples of the logs are commonly relatively coarse. But these subtleties could be coincidence.

4 | DISCUSSION

Clark and Steel (2006) showed that basin-floor fans of the Battfjellet Formation consist mainly (75%) of fine sand (125–250 μm ; 3–2 ϕ). Overall, this agrees with the grain-size distributions reported here, especially when considering Clark and Steel (2006) used the coarsest part of the beds to determine the grain size of the deposits. However, visual observations made in the field can neither record the range of grain sizes of individual deposits, which generates a sampling bias towards coarser grain sizes overall, nor the subtler differences in grain-size distribution between lobe sub-environments. The fact that almost all samples are slightly skewed towards finer material conforms to the observations Fildani et al. (2018) have reported from the lobes of the Mississippi Fan, that visual descriptions alone tend to overestimate the sand fraction of deposits. These considerations should motivate the research community to increase the integration of microscopic sample analyses into outcrop studies.

4.1 | Grain-size distribution and sub-environments

The measured grain-size distributions of the lobe axis, lobe off-axis and lobe fringe are significantly different suggesting that these sub-environments have distinctly different grain-size distributions (Figure 12). Despite the fact that classifications of these sub-environments, based on macroscopic facies characteristics, are transitional, assigning lobe sub-environments to lithofacies associations proved to be a rather powerful method to differentiate the grain-size distributions within the scale of a lobe. Only one sample from the lobe off-axis group strongly differs from the other off-axis samples and is also the coarsest of all the samples. This sample, sample 6 of log 2, was not classified during the macroscopic facies analysis as lobe axis, like the overlying deposits, because an erosional surface above the

bed was chosen as lower boundary of the lobe axis at log 2. Based on the grain-size distribution this sample could also be classified as part of the overlying lobe axis. Re-assigning individual samples to other sub-environments based on their microscopic grain-size distribution could be justified in practice, but would here generate the appearance of circularity: grain size correlates with facies association, outlying grain-size distributions are re-assigned to another facies association, and therefore also correlate to facies association. Therefore, no such re-interpretations based on the grain size results were made here to illustrate the power of the a priori distinction between sub-environments based on macroscopic facies characteristics.

As discussed in the section about grain-size evaluation methods, the difference in grain size is not related to difference in grain fabric since the deviation due to orientation is much smaller than the deviation between and within the sub-environments. It is assumed that the grain fabric maturity of all samples is similar to the tested sample, implying that the obtained magnitude of difference in grain size with orientation is similar in all samples. This assumption might be invalid as the tested sample is retrieved from a massive turbiditic sandstone. This type of deposit features a broad range of long-axis orientations and possibly results in less variation with orientation compared to structured beds with more mature fabric (Baas et al., 2007). This possible larger difference in orientation is largely eliminated by sampling at the base of the beds, which mostly consist of massive sandstones.

4.2 | Proximal to distal trend

The results that are presented in this study strengthen the idea that lobes do not just simply fine from an apex, but show a more complicated non-linear fining trend (Figure 12) and conform to results described from the

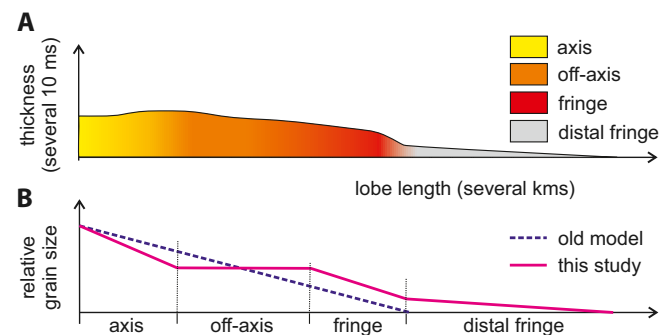


FIGURE 12 (A) Geometry and general distribution of lobe sub-environments. (B) Longitudinal grain size evolution of lobe deposits. The old linear model is indicated in a blue dashed line, whereas the non-linear trend suggested by this study is indicated by a pink solid line

Niger Delta X fan (Jobe et al., 2017), lobes from the Gerbe system (Aínsa Basin, Spain) (Bell et al., 2018) and experimental studies of Fernandez et al. (2014). In this study, lobe axis and fringe settings show a (weak) basinward fining trend, whereas lobe off-axis settings show no longitudinal decrease in grain size (Figure 12). For the Gerbe system down-dip decrease in grain size between lobe axis and off-axis was reported as minimal, whereas a more distinct decrease was documented between lobe off-axis and fringe deposits (Bell et al., 2018). The following section will discuss possible explanations for the observed fining trends.

Generally, there are two fundamental processes that govern deposition from a flow: competence and capacity. The competence of turbidity currents decreases from proximal to distal settings due to the loss in shear velocity, which is directly related to a loss of velocity (Kane et al., 2017; Kneller & McCaffrey, 2003). This decrease in flow velocity from the lobe apex is caused by the lateral spreading of flow, reduced sediment content due to sedimentation, and/or decreasing slope (Kneller, 1995), and is influenced by the topography of the lobe (Cantelli et al., 2011; Fernandez et al., 2014; Groenenberg et al., 2010; Spychala et al., 2017a). Sediment capacity describes the total amount of sediment a flow is able to transport. If sediment concentration surpasses a flow's capacity deposition is triggered by settling of the excess suspended sediment from the flow (Hiscott, 1994). It is argued that flows

that exit the channel mouth are at or near their capacity limit when lobe deposition starts.

One explanation for the decrease in the grain-size profile in the lobe axis combines grain-size stratification of the flow with the transition into lower momentum flows that is associated with the channel-lobe transition zone (Baas et al., 2004; Fildani et al., 2018). Pohl et al. (2019) show that turbidity currents tend to erode upon losing their confinement when passing from a confined channel environment into the unconfined lobe environment. As the flow incorporates more sediment through erosion, capacity-driven sedimentation is invoked (*sensu* Hiscott, 1994; Kneller & McCaffrey, 2003; Kubo, 2004). The coarse sediment fraction tends to be concentrated near the base of turbidity currents and is thus preferably deposited, whereas the medium and finer sediment fraction are more homogeneously distributed over the height of the current (Figure 13; de Leeuw et al., 2018a; Eggenhuisen et al., 2020; Garcia, 1994; Hansen et al., 2015; Kneller & Buckee, 2000; Tilston et al., 2015). Once the base (enriched in coarser grains) of the concentration profile has settled out, the off-axis sediment is deposited from a flow with a relatively homogeneous vertical grain size (Figure 13). Eventually, the flow thins and decreases in density due to continued loss of sediment resulting in the onset of competence-driven deposition in the lobe fringes (Baas et al., 2004; Kneller & McCaffrey, 2003) creating a second

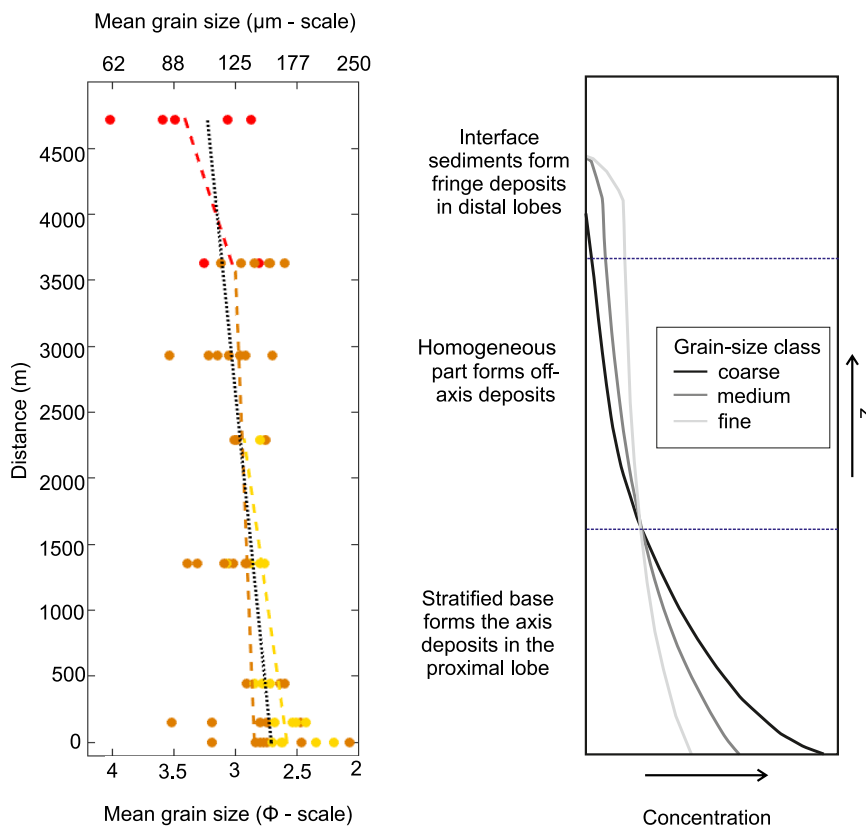


FIGURE 13 Observed longitudinal evolution of grain-size distribution from the youngest lobe of Clinotherm 12 compared with grain-size concentration profiles for different grain-size classes (after Eggenhuisen et al., 2020)

fining trend. This probably goes hand in hand with the deposition of most of the remaining sand grains, creating the sand pinch-out described from field observations of the lobe fringe (Prélat et al., 2009; Spychala et al., 2017a) long before the silt pinch-out in the distal fringe (Spychala et al., 2020). This evolution of the flow as it deposits sediment on its journey across a lobe would overall result in fining from the lobe axis to the intermediate and distal areas of the lobe off-axis and lobe fringe, which conforms to the trend seen in the deposits of Clinothem 12.

Grain-size stratification has previously been widely used to explain grain size trends in submarine channel and canyon systems (de Leeuw et al., 2018a, 2018b; Hansen et al., 2015; Hiscott et al., 1997; Hubbard et al., 2014; Jobe et al., 2017; Pirmez & Imran, 2003), but was so far not taken into consideration as an explanation of fining trends in lobe deposits. This study suggests that proximal to distal fining of lobe deposits is essentially the result of grain-size stratification in the flow exiting the channel, and the advection length of the different grain sizes (Boulesteix et al., 2020; Ganti et al., 2014; Spychala et al., 2020). The results of this study can be used to help the distribution of grain sizes in already existing simplified modelling approaches (Burgess et al., 2019; Pyrcz et al., 2005). The Clinothem 12 lobe deposits documented in this study may serve as a testing ground for such models (please refer to grain-size files in Material S1).

4.3 | Explaining the spread in lobe fringe samples

Lobe fringe samples collected from Log HS (Hyrnestabben, Figures 1C and 2) show a wide spread in mean grain sizes ranging from silt to fine sand (4.03 ϕ to 2.75 ϕ). It is noteworthy that the samples at this location were collected from lithofacies that are interpreted as being deposited by two different types of sediment gravity flows: turbidity currents and transitional flows. Samples from beds that are taken from turbidites are finer and better sorted than those taken from hybrid beds (Figure 8C,D) and their deposition is controlled by different flow processes. These differences in flow processes likely also cause a positive skew of grain-size distributions within turbidites, whereas hybrid beds are more likely to show a negative skew, especially in their lower division (Kane et al., 2017) enhancing the difference in grain size further. In contrast, Fildani et al. (2018) propose that hybrid beds from the Mississippi fan have a significant larger portion of silt and finer grains than the observed turbidites or can even be observed to be bimodal. Hybrid beds have been described to be deposited by several flow processes: Longitudinal evolution from a turbidity current (sensu Haughton et al., 2003) and vertical

segregation in supersaturated flows (sensu Baas et al., 2009; Kane et al., 2017). Whether differences in grain-size distribution can reflect the exact process is speculative at this point. However, neither of these characteristics are seen in the limited dataset of samples collected from hybrid beds. Deposits from a purely turbiditic succession would show a stronger fining trend, which is now dampened by the data from the hybrid bed deposits. It is clear that more work has to be conducted to study this longitudinal grain-size differentiation between the two flow types.

4.4 | Expression of stacking patterns in grain-size data

Lobe stacking patterns like progradation and retrogradation may skew the longitudinal grain-size distribution trend seen. While progradation would probably push the extent of coarse grain sizes further into the basin, retrogradation would cause a grain-size trend to finer material. In this case, there is no major influence of progradational or retrogradational stacking patterns in the evaluated grain-size distributions for the youngest lobe of Clinothem 12. Neither vertical grain-size distribution (Figure 11) nor bed thickness plots (Figure 14) for the individual logs show a clear coarsening and thickening upward trend, nor the fining and thinning upward trend which would be characteristic of progradation and retrogradation, respectively. This conforms with the observations made from by Prélat and Hodgson (2013) who suggest that lobes can show a full range of bed thickness patterns. The data given here support that this is true for different log sections of the same lobe deposit and indicates that the stacking patterns of the studied lobe are primarily governed by compensational stacking.

However, it is likely that the influence of stacking patterns would become more apparent if grain-size distributions for lobe complexes that comprise several lobes were analysed. Further work in a well reconstructed lobe complex example is necessary to assess this question.

4.5 | Implications for hydrocarbon reservoir quality

This study aims to understand how the spatial evolution of turbidity currents is recorded in lobe deposits on a scale of approximately 5 km. The results have implications for kilometre-scale distribution of hydrocarbon reservoir parameters, as grain size is a main factor control on reservoir quality (Bell et al., 2018; Ghaffar et al., 2017; Porten et al., 2016). The amalgamated structureless sandstones that characterise lobe-axis settings have been interpreted as high-density turbidity current deposits (Bell et al., 2018;

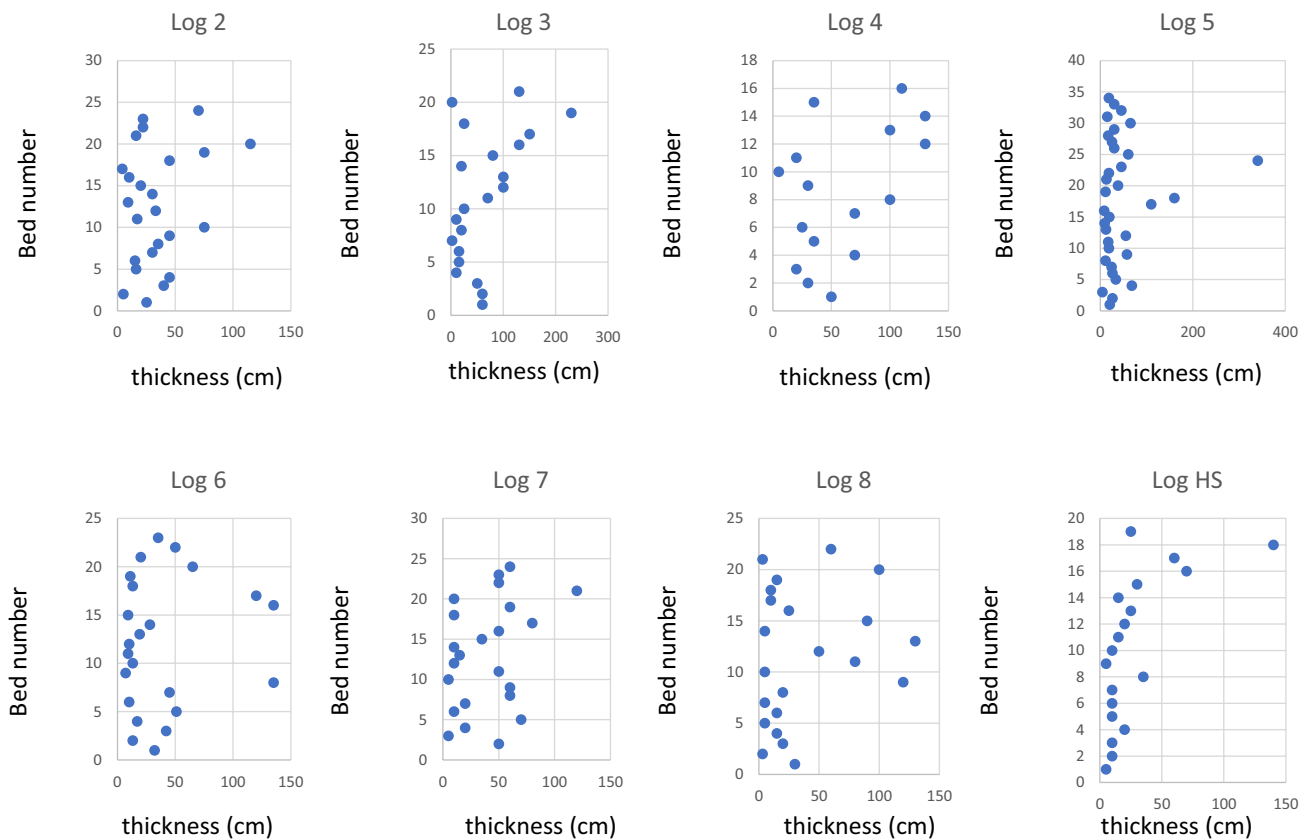


FIGURE 14 Vertical bed thickness distribution for the individual logs of Cliniothem 12. Note: Horizontal axis for log 3, log 5 and log HS deviate from the other logs

Grundvåg et al., 2014b; Prélat et al., 2009; Sychala et al., 2017a). Generally, these deposits are associated with the best reservoir characteristics as they are (a) commonly low on clay content which is connected to the highest permeabilities, and (b) have been suggested to have higher permeability because of their coarse grain size (Ghaffar et al., 2017; Porten et al., 2016). Ghaffar et al. (2017) report that permeability can decrease by at least 50% when grain size transitions from medium to fine sand. Lobe off-axes are characterised by sediments that are deposited from high-density turbidity currents as well as from low-density turbidity currents. Bell et al. (2018) did not observe a significant change in sorting between lobe axis and off-axis, which conforms to observations from this study. However, besides finer grain size, the lobe off-axis is associated with an increase in matrix content which decreases permeability (Bell et al., 2018). Lobe fringe deposits comprise fine-grained deposits from low-density turbidity currents and deposits that were deposited by flows that show turbulent as well as laminar behaviour (hybrid beds). Low-density turbidites are associated with the finest grain sizes, highest matrix content and lowest net: gross. The results of this study suggest that reservoir quality changes with distance most strongly within lobe axis and lobe fringe, albeit in

the case of lobe axis deposits the change is from very high quality to high quality. The reservoir quality of lobe off-axis deposits is demonstrated to be homogeneous without proximal to distal trends. This outcome has two implications: (a) It is justified to model lobes as high-quality reservoir elements right until their transition to lobe fringes; and (b) Reservoir quality can be uncoupled from reservoir thickness as the intermediate lobe deposits show thinning without a change in depositional sand quality. These outcomes can augment stochastic lobe modelling approaches (Pyrzcz et al., 2005) by adding the additional factor of grain-size trends to the model.

5 | CONCLUSIONS

The basin-floor fans of the Battfjellet Formation in Svalbard that outcrop in a proximal to distal orientation create the rare opportunity to analyse the longitudinal evolution of these distal deposits from lobe axis to fringe over a distance of 5 km. Within the investigated basin floor segment of Cliniothem 12, lobe sub-environments have been distinguished based on their characteristic lithofacies associations and bounding surfaces, which

reflect depositional flow properties. A transition from predominantly high-density turbidity current deposits to low-density turbidity current deposits and hybrid beds is observed. A total of 66 photographed thin sections were manually counted to determine the grain-size distribution of lobe axis, lobe off-axis and lobe-fringe deposits. These sub-environments have statistically distinct and variable grain-size distributions superimposed on an overall non-linear fining trend from lobe axis to lobe fringe. Only a single sample yielded grain-size results that could justify re-assignment of the macroscopic facies association that was ascribed a priori the thin section analysis. This success rate reinforces the confidence in established facies classification schemes for lobe deposits. Furthermore, the sub-environments show different longitudinal grain-size trends. Lobe axis and fringe settings show (weak) basinward fining trends whereas lobe off-axis settings show no variation in grain-size distribution with distance. An explanation for these trends is the interplay of capacity-driven deposition from the decelerating flow, the grain-size stratification of the flows, and advection of the flow over the lobe. Overall, the coarser grain-size fraction is enriched at the base of the flow and more prone to be deposited when capacity-driven deposition is evoked in the lobe axis. The lobe off-axis is suggested to be deposited from the more homogeneously mixed part of the flow, whereas the lobe fringe is dominated by deposition of the upper parts of the flow, resulting in the sand pinch-out in the lobe fringe.

The implications of this study are multifaceted: (a) Grain-size stratification is an important factor in the evolution and deposition of turbidity currents over lobe deposits and may be incorporated into modelling approaches of lobe deposits; (b) it is justified to model lobes as high-quality bodies until their transition to lobe fringes; and (c) reservoir quality can be uncoupled from reservoir thickness as the intermediate lobe deposits show thinning but no change in depositional sand quality.

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DATA AVAILABILITY STATEMENT

All raw grain size data are in the supporting materials. For any additional information feel free to contact the corresponding author who is happy to share these with you.

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