

Target-oriented imaging using extended image volumes: a low-rank factorization approach

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ABSTRACT

Imaging in geological challenging environments has led to new developments, including the idea of generating reflection responses by means of interferometric redatuming at a given target datum in the subsurface, when the target datum lies beneath a complex overburden. One way to perform this redatuming is via conventional model-based wave-equation techniques. But those techniques can be computationally expensive for large-scale seismic problems since the number of wave-equation solves is equal to two times the number of sources involved during seismic data acquisition. Also conventional shot-profile techniques require lots of memory to save full subsurface extended image volumes. Therefore, we can only form subsurface image volumes in either horizontal or vertical directions. To exploit the information hidden in full subsurface extended image volumes, we now present a randomized singular value decomposition-based approach built upon the matrix probing scheme, which takes advantage of the algebraic structure of the extended imaging system. This low-rank representation enables us to overcome both the computational cost associated with the number of wave-equation solutions and memory usage due to explicit storage of full subsurface extended image volumes employed by conventional migration methods. Experimental results on complex geological models demonstrate the efficacy of the proposed methodology and allow practical reflection-based extended imaging for large-scale five-dimensional seismic data.

Key words: Extended-image volume, Interferometry, Low-rank factorization, Redatum, Point spread function, Randomized SVD, Target-imaging.

1 INTRODUCTION

Seismic data are the primary source of information used to image structures, evaluate rock properties and monitor their changes in the Earth's subsurface. To achieve these goals, seismic reflection imaging relies on kinematically correct background-velocity models. However, imaging from acquired data struggles to reveal the true location of the subsurface reflectors at the deeper sections, particularly in geological environments presenting complex overburden structures such

as high-velocity salt bodies or low-velocity gas chimneys. This is due to the fact that the ray-based imaging methods cannot adequately describe the wave propagation in complex geological scenarios and sophisticated wave-equation-based methods are very sensitive to the accuracy of the near-surface or overburden velocity model.

The seismic community proposed various methods to overcome these difficulties. These methods can be roughly divided into two main categories. The first method is called virtual source technology (Bakulin and Calvert 2004, 2008), where borehole seismic data are acquired with receivers

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being deployed beneath the complex part of the overburden and sources are placed at the surface. Virtual source data at the receivers locations are created using time reversal techniques applied to the surface-to-downhole data, where time-gated data are passed through the time reversal process and convolved with the surface-to-downhole data. The resulting virtual source data are then imaged with conventional methods using the velocity information below the receivers. Although this method is beneficial for removing artifacts caused by the complex overburden, it requires receivers to be placed in the borehole, which is not practical in exploration surveys.

The second method is called interferometric redatuming, which has been studied extensively in the literature (Snieder 2004; Schuster and Zhou 2006; Snieder, Sheiman and Calvert 2006; Wapenaar and Fokkema 2006; Mehta *et al.* 2007; Schuster 2009; Curtis and Halliday 2010; Brogini, Snieder and Wapenaar 2012; Wapenaar, Brogini and Snieder 2012; van der Neut and Herrmann 2013). In interferometric redatuming, seismic data are acquired by placing both sources and receivers at the surface during the acquisition. Then a virtual seismic response at the subsurface datum is created by performing a multi-dimensional deconvolution of the upgoing waves with respect to the downgoing waves, where the downgoing and upgoing wave constituents are created from the transmitted waves by inverting the model-based wavefield composition matrix (Schuster and Zhou 2006; Wapenaar, Slob and Snieder 2008). One drawback of this method is that one has to invert this composition matrix, which can be challenging because it exhibits singularities at critical angles. We can overcome this by regularizing the least-square solutions to generate the down- and upgoing wavefields as shown in van der Neut and Herrmann (2013), where the authors proposed to solve a sparsity-promotion-based least-squares formulation. Alternatively, Wapenaar and van der Neut (2010) and Vasconcelos and Rickett (2013) proposed multi-dimensional deconvolution-based imaging conditions using extended image volumes, where two-way wave equation-based extrapolation is used to generate the full non-linear reflection response at a given target-datum in the subsurface, as a function of time or frequency.

Although, multi-dimensional deconvolution based imaging conditions using extended image volumes have shown potential benefits in target-oriented imaging (van der Neut and Herrmann 2013; Vasconcelos and Rickett 2013; Ravasi *et al.* 2016), evaluating these reflection responses is prohibitively expensive in realistic three-dimensional (3D) settings. This is due to the fact that it involves solving two-way wave-equation extrapolation with a conventional shot-profile approach at

any particular datum in the subsurface. For the conventional shot-profile approach, the required number of partial differential (wave) equation's (PDE's) evaluations is equal to twice the number of sources. Since, we need to solve so many PDEs, the use of conventional two-way approaches to perform target-oriented imaging is very expensive especially for industrial scale problems in 2D and 3D.

To overcome the computational cost of building extended image volumes, analytic solutions have been proposed in ten Kroode (2012), Hou and Symes (2015), Symes (2008), Stolk, De Hoop and Symes (2009) and Stolk and De Hoop (2005) by adding horizontal subsurface offset variables. Here we propose to analyse the algebraic structure of full subsurface extended image volumes organized as a matrix. Each column of this matrix represents a source experiment that captures the reflection response in both horizontal and vertical directions since it is derived from the two-way wave-equation (Berkhout 1992; van Leeuwen, Kumar and Herrmann 2017). Following van Leeuwen, Kumar and Herrmann (2017), we build the full extended image volume by using matrix-free operations with probing vectors that cost only two PDE solves per vector. However, this still requires to compute the extended image volume for each single point source and the final resulting matrix is often too huge for storage.

By construction, we find that monochromatic full subsurface extended image volumes exhibit a low-rank structure for low-to-mid range frequencies. If we can somehow gain access to this low-rank structure, we will be able to reduce the number of required PDE solves to a number that is proportional to the rank and not the number of sources. To accomplish this, we propose a framework using randomized singular value decomposition (SVD) (Halko, Martinsson and Tropp 2011) that gives us the needed low-rank factorization of full subsurface extended image volumes. Unlike van Leeuwen, Kumar and Herrmann (2017), we now probe the extended image volume (EIV) with a small number of random vectors rather than the Dirac vectors as suggested before. This means that, instead of considering a single point source only, we shoot a random combination of all sources together. This random configuration of the sources has already been proposed for full waveform inversion in Haber, Chung and Herrmann (2012), but not for EIVs. Then, thanks to the randomized SVD algorithm (Halko, Martinsson and Tropp 2011), we are able to reveal a low-rank factorization of the EIV, consisting of two thin matrices. As a result, the computation cost is reduced and the low-rank approximation of the EIV is now storable. Moreover, from this low-rank representation of the EIV, we can extract reverse time migration (RTM) images

and/or common-image point gathers through easy algebraic operations, whereas the former probing technique (with Dirac vectors) required to recompute the EIV for each extraction. Finally, the randomized SVD algorithm has already been used for seismic data, for instance in Demanet *et al.* (2012) and Jumah and Herrmann (2014), but not yet applied to EIV.

With this technique, we only need to compute a low-rank representation of the full subsurface extended image volume, which contains all subsurface offsets for all subsurface points once. Using this low-rank representation of the EIV, which requires greatly reduced storage, we can extract information, like the RTM image from the diagonal or the common-image point gathers from columns of the full matrix, of any part of the model without recomputing it. This was not the case in earlier work by van Leeuwen, Kumar and Herrmann (2017), since the probing technique with Dirac vectors do not allow us to build and store the entire EIV for all subsurface points. For this reason, we needed to recompute the probed EIV everytime for each subsurface point in the model. As a result, the computational and storage cost of forming the EIV for all subsurface offsets and all subsurface points is now much smaller using the proposed randomized SVD formulation compared to the conventional probing.

The paper is organized as follows: we begin by introducing the multi-dimensional deconvolution-based extended imaging conditions in a two-way wave-equation setting. Next, we show a computationally affordable and memory-efficient way of computing the reflection response at a subsurface datum using probing techniques. We illustrate these features on a small section of the Marmousi model for which we can form the matrices for comparison and we show that our low-rank approach is in principle scalable with the model size. On a larger more complex model, we carry out our redatuming and show that we are able to handle the low-velocity zone in the overburden. Although, probing techniques overcome an important computational bottleneck, we only have very limited access to the full subsurface reflection response because we are restricted to the physical locations of the probing vectors. To get access to the full subsurface image volume for all subsurface points, we propose to take the randomized singular value decomposition with probing a step further by representing the image volumes themselves in a low-rank factored form. Finally, we demonstrate the computational advantages of low-rank factored form to perform the target-oriented imaging beneath the salt using a section of Sigsbee2A model (<http://www.delphi.tudelft.nl/SMAART/sigsbee2a.htm>) (The SMAART JV 2001).

2 EXTENDED IMAGE VOLUMES VIA DECONVOLUTION

Before we describe the proposed methodology of extracting the target-oriented reflection response from full subsurface extended image volumes, we first review the governing equations. Given the time-harmonic source and receiver wavefields, denoted by $u(\mathbf{x}, \mathbf{x}_s, \omega)$ and $v(\mathbf{x}', \mathbf{x}_s, \omega)$, the reflection response r (Berkhout 1992; Vasconcelos, Sava and Douma 2010; Thomson 2012) of the medium below the datum ∂D_d satisfies the following relationship

$$v(\mathbf{x}', \mathbf{x}_s, \omega) = \int_{\mathbf{x} \in \partial D_d} r(\mathbf{x}', \mathbf{x}, \omega) u(\mathbf{x}, \mathbf{x}_s, \omega) d^2 \mathbf{x}, \quad (1)$$

where $\mathbf{x}, \mathbf{x}' \in \mathcal{D} \subset \mathbb{R}^n$ ($n = 2$ or 3) represent any subsurface position, $\omega \in \Omega \subset \mathbb{R}$ is the temporal frequency and $\mathbf{x}_s \in \mathcal{D}_s \subset \mathbb{R}^{n-1}$ the location of sources deployed at the surface. For more details on how to derive the subsurface reflection response in the acoustic case, we refer to Wapenaar and van der Neut (2010). The source and receiver wavefields are obtained by solving the following two-way wave-equations

$$\begin{aligned} H(m)u(\mathbf{x}, \mathbf{x}_s, \omega) &= q(\mathbf{x}, \mathbf{x}_s, \omega), \\ H(m)^* v(\mathbf{x}', \mathbf{x}_s, \omega) &= \int_{\mathcal{D}_r} d\mathbf{x}_r d(\mathbf{x}_r, \mathbf{x}_s, \omega) \delta(\mathbf{x}' - \mathbf{x}_r), \end{aligned} \quad (2)$$

where $H(m) = \omega^2 m(\mathbf{x})^2 + \nabla^2$ is the Helmholtz operator with Sommerfeld boundary conditions, m the squared slowness, the symbol $*$ represents the conjugate-transpose, q denotes the source function and d the reflection-data at receiver positions $\mathbf{x}_r \in \mathcal{D}_r \subset \mathbb{R}^{n-1}$.

To do numerical computations, we discretize our domain of interest \mathcal{D} with a rectangular grid with a total of N_x grid points. We will denote N_s the number of sources, N_r the number of receivers and N_f the number of frequencies. The discretized source and receiver wavefields \mathbf{U} and \mathbf{V} become three-dimensional (3D) tensors of size $N_x \times N_s \times N_f$. For the i th frequency, \mathbf{U}_i and \mathbf{V}_i represent complex valued source and receiver wavefields organized as $N_x \times N_s$ matrices, where each column of these matrices contains a monochromatic source experiment. The reflection response can also be expressed as a 3D tensor \mathbf{R} of size $N_x \times N_x \times N_f$, where a slice \mathbf{R}_i at the i th frequency is a $N_x \times N_x$ matrix, which satisfies

$$\mathbf{V}_i = \mathbf{R}_i \mathbf{U}_i. \quad (3)$$

The discrete form of equation (2) is then obtained by solving

$$\begin{aligned} \mathbf{U}_i &= \mathbf{H}_i(\mathbf{m})^{-1} \mathbf{P}_s^T \mathbf{Q}_i, \\ \mathbf{V}_i &= (\mathbf{P}_r \mathbf{H}_i(\mathbf{m})^{-1})^* \mathbf{D}_i, \end{aligned} \quad (4)$$

where \mathbf{H}_i is a discretization of the two-way Helmholtz operator ($\omega_i^2 \mathbf{m} + \nabla^2$) for temporal frequency ω_i and for a gridded squared slowness \mathbf{m} . We use the notations $*$ for the conjugate-transpose, $^{-1}$ for its inverse and T for the transpose. Each column of the $N_s \times N_s$ matrix \mathbf{Q}_i represents the discretized source function, and the $N_r \times N_s$ matrix \mathbf{D}_i contains the monochromatic observed wavefields, without free-surface multiples, after removing the direct wave and source and receiver side ghosts. Matrices \mathbf{P}_s and \mathbf{P}_r sample the wavefields at the source and receiver positions. Their transpose injects the sources and receivers wavefields into the grid at the source and receiver locations. For the sake of simplicity, we drop the frequency index i for the remainder of paper.

To perform the target-oriented imaging, we need access to the reflection response \mathbf{R} at the subsurface datum ∂D of interest for given source and receiver wavefields. Because \mathbf{U} is not a full rank (the maximum rank of \mathbf{U} is N_s), we cannot straightforwardly compute its inverse to calculate \mathbf{R} . For this reason, we compute the subsurface reflection response \mathbf{R} in the least-squares sense as follows:

$$\underset{\mathbf{R}}{\text{minimize}} \quad \frac{1}{2} \|\mathbf{V} - \mathbf{R}\mathbf{U}\|_F^2, \quad (5)$$

where $\|\mathbf{A}\|_F^2 = \sum_{i,j} |a_{i,j}|^2$ denotes the Frobenius norm squared. One way to solve equation (5) is by cross-correlating both sides of equation (3) with \mathbf{U} , which results in the following normal equation

$$\underbrace{\mathbf{V}\mathbf{U}^*}_{\mathbf{E}} = \mathbf{R} \underbrace{(\mathbf{U}\mathbf{U}^*)}_{\mathbf{\Gamma}}. \quad (6)$$

Here,

$$\mathbf{E} = \mathbf{V}\mathbf{U}^* = \mathbf{H}^{-*} \mathbf{P}_r^T \mathbf{D} \mathbf{Q}^* \mathbf{P}_s \mathbf{H}^{-*}$$

represents the monochromatic extended image for all subsurface offsets and for all subsurface points computed as the cross-correlation between the source and receiver wavefields, whereas

$$\mathbf{\Gamma} = \mathbf{U}\mathbf{U}^* = \mathbf{H}^{-1} \mathbf{P}_s^T \mathbf{Q} \mathbf{Q}^* \mathbf{P}_s \mathbf{H}^{-*}$$

corresponds to the monochromatic wavefield point-spread function (PSF), which can be seen as the radiation pattern of the virtual sources (van der Neut *et al.* 2010). Both \mathbf{E} and $\mathbf{\Gamma}$ are matrices of size $N_x \times N_x$. Given the above relationship, \mathbf{E} can be interpreted as a defocusing (lower resolution) of the reflection response \mathbf{R} by the PSF function $\mathbf{\Gamma}$ for all subsurface offsets and for all subsurface points that lie on the subsurface datum ∂D_d .

To retrieve \mathbf{R} from equation (6), we need to compute the inverse of $\mathbf{\Gamma}$ and apply it as a filter to the cross-correlated

source and receiver wavefields \mathbf{E} . This filtering process to recover the focused \mathbf{R} is known as multi-dimensional deconvolution (MDD) (Wapenaar, Slob and Snieder 2008). Opposed to cross-correlation (Wapenaar and Fokkema 2006), the retrieval of \mathbf{R} with multidimensional deconvolution (Herrmann and Wang 2008; Wapenaar and van der Neut 2010), compensates for illumination, finite aperture, and intrinsic losses, etc. In classical settings, the process to generate the reflection response at the subsurface datum ∂D_d corresponds to the following steps: (i) solve the wave-equation twice for each surface source to calculate the \mathbf{U} and \mathbf{V} matrices, respectively, (ii) perform the cross-correlations as per equation (6) to generate \mathbf{E} and $\mathbf{\Gamma}$, (iii) perform the MDD in the least-squares sense to retrieve \mathbf{R} . As a consequence, we have to solve the wave-equation for all sources and receivers within the seismic survey to compute the reflection response at a particular datum in the subsurface. This operation can easily become computationally and storage wise too expensive even for relatively small two-dimensional cases.

To overcome the computational cost of calculating a least-squares estimate for \mathbf{R} at a particular datum in the subsurface, we follow the probing technique proposed by van Leeuwen, Kumar and Herrmann (2017). This technique is designed to perform matrix-free operations to extract information from matrices that cannot be formed explicitly because they are too big and dense. In this way, we remove most of the computational and storage requirements to work with the full-subsurface extended image volumes. Following van Leeuwen, Kumar and Herrmann (2017), in order to probe we multiply equation (6) from the right by a tall $N_x \times K$ ($K \ll N_x$) probing matrix $\mathbf{W} = [\mathbf{w}_1, \dots, \mathbf{w}_K]$ yielding

$$\begin{aligned} \tilde{\mathbf{E}} &= \mathbf{E}\mathbf{W} = \mathbf{H}^{-*} \mathbf{P}_r^T \mathbf{D} \mathbf{Q}^* \mathbf{P}_s \mathbf{H}^{-*} \mathbf{W}, \\ \tilde{\mathbf{\Gamma}} &= \mathbf{\Gamma}\mathbf{W} = \mathbf{H}^{-1} \mathbf{P}_s^T \mathbf{Q} \mathbf{Q}^* \mathbf{P}_s \mathbf{H}^{-*} \mathbf{W}. \end{aligned} \quad (7)$$

In this expression, K is the number of probing vectors. We denote probed quantities by the symbol $\tilde{\cdot}$. For redatuming, we chose the columns \mathbf{W} , \mathbf{w}_j for $j = 1 \dots K$, to be given by $\mathbf{w}_j = [0, \dots, 0, 1, 0, \dots, 0]$ – that is a vector containing a single point scatterer at the j th grid location. For the sake of simplicity, we will call Dirac vectors the \mathbf{w}_j defined above. Also, if we take $K = N$ the total number of grid points in the background, probing matrix \mathbf{W} is the identity matrix and the probed EIV $\tilde{\mathbf{E}}$ is exactly equal to \mathbf{E} . The same argument holds for $\tilde{\mathbf{\Gamma}}$ and $\mathbf{\Gamma}$. Therefore, computing the EIV with the Dirac vectors describing the entire background is equivalent to perform conventional MDD. Now, to extract the EIV for a specified localized area in the background, we only consider

the corresponding Dirac vectors in the probing. This is mathematically equivalent to extract the corresponding area from the entire EIV once fully computed.

By virtue of the choice of Dirac probing vectors, the \mathbf{w}_j plays the role of a point source located at the datum of interest ∂D_d . As a result of this type of probing, each column of $\tilde{\mathbf{E}}$ represents a common-image point (CIP) gather and each column of $\tilde{\mathbf{\Gamma}}$ represents the source-side blurring kernel at the subsurface locations corresponding to the \mathbf{w}_j 's. In Algorithm 1, we describe how to compute $\tilde{\mathbf{E}}$ and $\tilde{\mathbf{\Gamma}}$ in a computationally efficient manner. According to van Leeuwen, Kumar and Herrmann (2017), Algorithm 1 proposes a way to compute the image volume through matrix–vectors multiplications, which reduces the cross-correlation of the wavefields \mathbf{U} and \mathbf{V} to the cross-correlation of the data matrices $\tilde{\mathbf{Q}} = \mathbf{Q}^* \tilde{\mathbf{D}}$, see step 2. With this technique, we avoid the loop over all shots that are necessary to build wavefields \mathbf{U} and \mathbf{V} in the conventional computation. The remaining cost lies in steps 1 and 5, where we need to solve two PDEs (one per step) for each probing vector, whereas one has to solve $2N_s$ PDEs to form both wavefields \mathbf{U} and \mathbf{V} . As a result, Algorithm 1 proves to be faster and requires less storage capacity as long as the number of probing vectors K is much smaller than the number of sources N_s . This may often be the case when we consider target-oriented imaging for a small area of the background.

Algorithm 1. Computation of $\tilde{\mathbf{E}}$ and $\tilde{\mathbf{\Gamma}}$ using matrix-free probing techniques (van Leeuwen, Kumar and Herrmann 2017) with given sampling vectors $\mathbf{W} = [\mathbf{w}_1, \dots, \mathbf{w}_K]$. The computational cost is $3N_x$ wave-equation solves to which the cost of correlating the wavefields should be added

Inputs: source matrix \mathbf{Q} , data matrix \mathbf{D} and probing matrix \mathbf{W}

Outputs: probed extended image volume $\tilde{\mathbf{E}}$, and probed point-spread function $\tilde{\mathbf{\Gamma}}$

1. compute $\tilde{\mathbf{U}} = \mathbf{H}^{-*} \mathbf{W}$ and sample this wavefield at the source locations $\tilde{\mathbf{D}} = \mathbf{P}_s \tilde{\mathbf{U}}$,
 2. correlate the results with source function $\tilde{\mathbf{Q}} = \mathbf{Q}^* \tilde{\mathbf{D}}$,
 3. use the result as data weights $\tilde{\mathbf{L}}_1 = \mathbf{D} \tilde{\mathbf{Q}}$ and $\tilde{\mathbf{L}}_2 = \mathbf{Q}^* \tilde{\mathbf{Q}}$,
 4. inject the wavefield $\tilde{\mathbf{L}}_1$ at the receiver locations and compute the extended image as $\tilde{\mathbf{E}} = \mathbf{H}^{-*} \mathbf{P}_r^T \tilde{\mathbf{L}}_1$,
 5. inject the wavefield $\tilde{\mathbf{L}}_2$ at the source locations and compute the PSF as $\tilde{\mathbf{\Gamma}} = \mathbf{H}^{-1} \mathbf{P}_s^T \tilde{\mathbf{L}}_2$.
-

Now that we have access to probed versions of the relevant quantities ($\tilde{\mathbf{E}}$ and $\tilde{\mathbf{\Gamma}}$), we are now in a position to determine the reflection response from the relation: $\tilde{\mathbf{E}} = \tilde{\mathbf{R}} \tilde{\mathbf{\Gamma}}$. Note that compared to the earlier expressions, the matrix $\tilde{\mathbf{R}}$ only

contains the reflection response restricted to the datum of interest ∂D_d probed by \mathbf{W} . As before, we need to compute the inverse of $\tilde{\mathbf{\Gamma}}$ to compute $\tilde{\mathbf{R}}$. Since $\tilde{\mathbf{\Gamma}}$ is rank deficient with a maximum rank of K , we compute the reflection response via

$$\tilde{\mathbf{R}} = \tilde{\mathbf{E}} \tilde{\mathbf{\Gamma}}^\ddagger; \quad (8)$$

where \ddagger represents the damped Moore–Penrose pseudo-inverse. After applying this pseudo-inverse, each column of $\tilde{\mathbf{R}}$ now contains a CIP gather after source-side deblurring at the subsurface locations represented by the \mathbf{w}_j 's located on the datum of interest ∂D_d . To summarize, target-oriented imaging involves the following steps: (i) compute $\tilde{\mathbf{R}}$ at a given datum ∂D_d using equation (8) for each frequency; (ii) take the inverse Fourier-transform along the time-axis to get the time-domain reflection response; and (iii) perform time-domain reverse time-migration to generate the image in the area of interest below the datum ∂D_d . Before we demonstrate the validity of this approach, let us first make a detailed comparison of the computational costs of our approach versus those that use more traditional techniques.

3 COMPUTATIONAL ASPECTS

The key contribution of this work lies in illustrating the advantages probing techniques have compared to the cost involved on methods that only rely on multi-dimensional deconvolutions (MDDs). For this purpose, we compare the computational cost and memory usage of calculating the reflection response at a single scattering point at the datum ∂D_d somewhere in the subsurface. As we can see from Algorithm 1, solving the wave-equation constitutes the main computational bottleneck in retrieving the reflection response because we have to solve as in any shot-profile scheme, $2N_s$ partial differential equation (PDE). Conversely, with the probing techniques of equation (7), we only need to compute three PDEs per subsurface point irrespective of the number of sources N_s in the acquisition survey. As a result, probing techniques can significantly reduce the computational cost as long as the number of probing vectors K is small compared to the number of sources – that is $K \ll N_s$. Note that for now K represents the number of points in the domain where we want to compute the common-image point (CIP), therefore we collect the corresponding K Dirac probing vectors to retrieve that information for the chosen points, without building the entire EIV.

To provide further evidence of the computational and memory savings due to probing, we retrieve the matrix $\tilde{\mathbf{R}}$ from a central part of the Marmousi model (Versteeg 1994) (Fig. 1). We choose a small section so we can form $\tilde{\mathbf{R}}$ explicitly

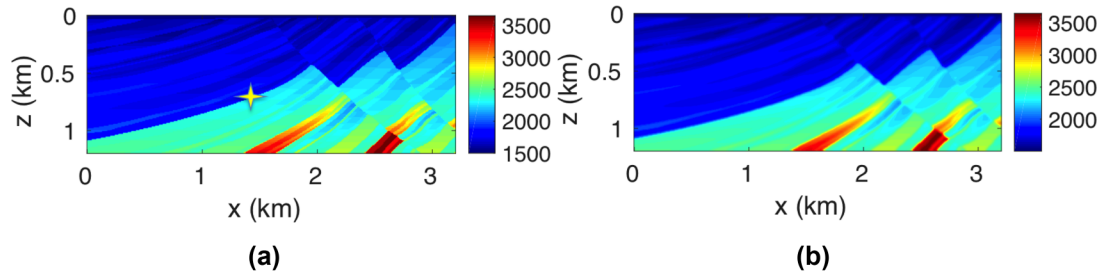


Figure 1 A section of the Marmousi model. (a) true and (b) smooth velocity model (m/s). The yellow star represents the position of a subsurface datum point, where we compute the reflection matrix \mathbf{R} .

for comparison. Still this toy section does contain strong lateral variations and dipping reflectors. We use a fixed-spread acquisition configuration with a model grid spacing of 5 m, 401 co-located sources and receivers with 10 m spacing. We only perform comparison for a single frequency at 10 Hz. The source wavelet is a Ricker wavelet with a peak frequency of 30 Hz. The results are displayed in Table 1 where we report the computational time (in seconds) and memory (in MB) required to compute $\tilde{\mathbf{R}}$ at a single point in the subsurface. Although we get the same CIP gather using the classical approach versus the proposed approach, probing reduces the computational time by a factor of 10 and memory requirement by a factor of 20. This is a first illustration of the computational and memory benefits of probing techniques over classical dense matrix techniques.

To understand how our approach scales, we also plot a comparison between computational times of probing and the conventional approach as we increase the size of the Marmousi model (Fig. 2). This example confirms that our method indeed scales better. We can see that the difference in computational time is substantial with respect to the model size even for this small two-dimensional (2D) example. We expect the benefit of our probing technique to be even more pronounced when we switch from 2D to 3D seismic data acquisition where both the size of the model and the number of sources grow drastically.

Table 1 Computational complexity of the probing technique versus classical technique for computing the reflection response in terms of the number of sources N_s and sample points N_x

	# of PDE Solves	Time(s)	Memory(MB)
Conventional	$2N_s$	190	710
This paper	$3N_x$	15	35

Notes: Comparison of the computational time (in seconds) and memory (in megabytes) while computing \mathbf{R} at a datum point on a central part of the Marmousi model. For this example, $N_s = 401$ and $N_x = 1$.

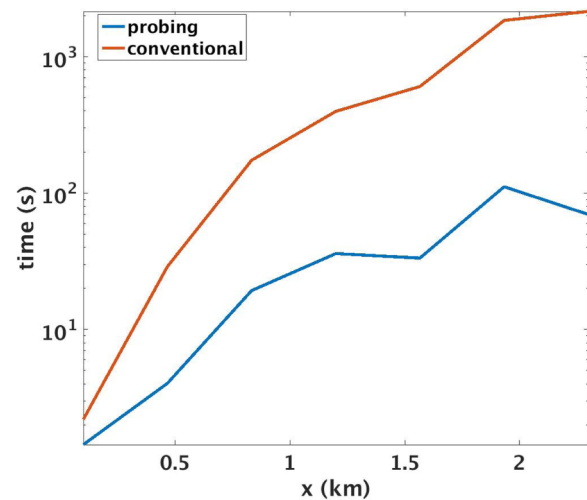


Figure 2 Comparison of computational time as the size of the Marmousi model increases: clearly probing techniques are a computationally feasible approach to compute the reflection matrix \mathbf{R} whereas the classical approach proves to be quickly expensive for the large size of the model.

3.1 Target-oriented imaging with probing

We further demonstrate the advantages of probing for redatuming and target-oriented imaging. To do so, we first consider a simple velocity model consisting of horizontal layers with low-velocity zones and a complex basement structure with a fault (see Fig. 3). The aim here is to remove the effect of a low-velocity zone to better image the basement of interest. For this example, the velocity model is 4.6 km deep and 7.5 km wide, sampled at 5 m. The synthetic data are computed from 376 sources and receivers sampled at 20 m, respectively. We used the range of frequencies from 5 to 45 Hz sampled at 0.25 Hz, where the source-signature is a Ricker wavelet with a central frequency of 20 Hz. The subsurface datum ∂D_d is represented by the dashed white line in Fig. 3(a). To capture the redatuming data created by the complex basement

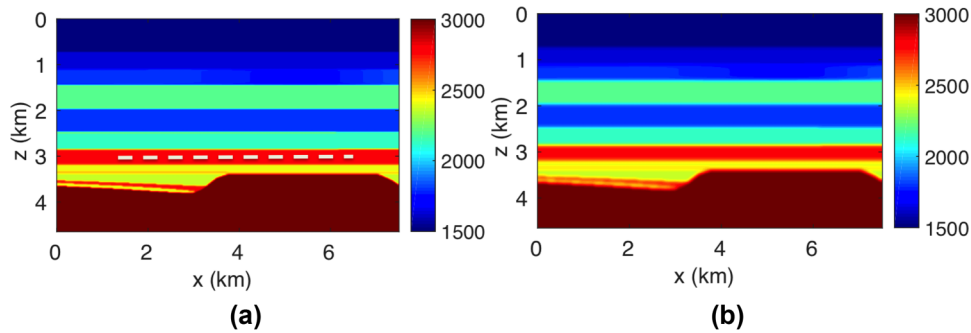


Figure 3 Velocity model (m/s) with fault: (a) true and (b) smooth velocity model with the position of subsurface datum (dashed white line).

with fault, we take $K = 50$ virtual subsurface locations with a sampling interval of 20 m, which covers the area of interest, and compute the reflection response according to the method outlined above.

Figure 4(a,c, d,f) show the frequency-space domain and time-space domain reflection responses at datum point $x = 3$ and $z = 2.5$ km before and after source-side blurring correction. To calculate the frequency-space domain reflection response, we solve equation (8) for all the frequencies, followed by restricting it to the subsurface datum ∂D_d , which will result in a reflection matrix \mathbf{R} where rows correspond to frequencies and columns correspond to horizontal offset δx . For K probing vectors, \mathbf{R} will be a three-dimensional (3D) tensor of size $N_{\delta x} \times K \times N_{\omega}$. We perform the inverse Fourier-transform along the time-axis to go from frequency-space domain to time-space domain. In Fig. 4(c), we observe that the energy across the frequency spectrum of \mathbf{R} is well distributed compared to the spectrum of \mathbf{E} . This results in a better resolution of reflection events as observed in time-space domain sections in Fig. 4(f). In this relatively small numerical example, the computation of reflection response by probing is almost three times faster than its shot-profile counterpart.

The time-domain reverse-time migration images before and after blurring correction are shown in Fig. 5(a,b). We can see that the blurring correction makes the reflectors sharper but creates ringing effects, which is a side effect of multi-dimensional deconvolution (MDD). To remove some of these artifacts, Herrmann and Wang (2008) and van der Neut and Herrmann (2013) proposed a curvelet-based sparsity promotion approach to perform the MDD, which is more stable with respect to missing data, finite aperture and bandwidth limitation. We will address this issue in a future paper and concentrate on the computational aspects.

Although probing techniques are an elegant approach to retrieve the reflection response in a computationally feasible manner, this approach only gives us limited access to

the full subsurface reflection matrix. Consequently, each new target of interest requires new computations of $\tilde{\mathbf{E}}$ and $\tilde{\mathbf{\Gamma}}$. This can become an expensive procedure when we are interested in multiple targets especially when we are dealing with large-scale 3D problems. To address this issue, we propose to exploit the low-rank structure of full subsurface extended image volumes themselves. In this approach, the idea is to factorize the full subsurface image volumes, that is \mathbf{E} and $\mathbf{\Gamma}$. Given this factorized form, we can again use probing with \mathbf{W} to extract the columns of interest from \mathbf{E} and $\mathbf{\Gamma}$. Of course, this approach is only viable if we manage to keep the number of probing vectors required by the low-rank factorization process small compared to the number of sources.

4 LOW-RANK FACTORIZATION OF IMAGE VOLUMES – A RANDOMIZED SINGULAR VALUE DECOMPOSITION APPROACH

As we mentioned before, a monochromatic full subsurface extended image volume is a matrix of size $N_x \times N_x$, where N_x is the total number of subsurface grid points. Clearly, this is a very large object, especially in three dimension, suggesting a need to exploit possible redundancies living in these volumes. Mathematically, matrices with redundancy of information exhibits a low-rank structure. Following the gist of our probing approach, we analyse the decay of the singular values of these volumes for a realistic complex geological model to see whether these extended image volumes can indeed be factored into a low-rank form. In particular, we are interested in whether these singular values decay fast so that these volumes can be approximated accurately by a rank r matrix, where $r \ll N_x$. According to equation (6), the rank of the monochromatic full subsurface extended image volume is linked to the rank of the data matrix \mathbf{D} or the source matrix \mathbf{Q} and is equal to N_s where $N_s \ll N_x$. While \mathbf{Q} is full rank

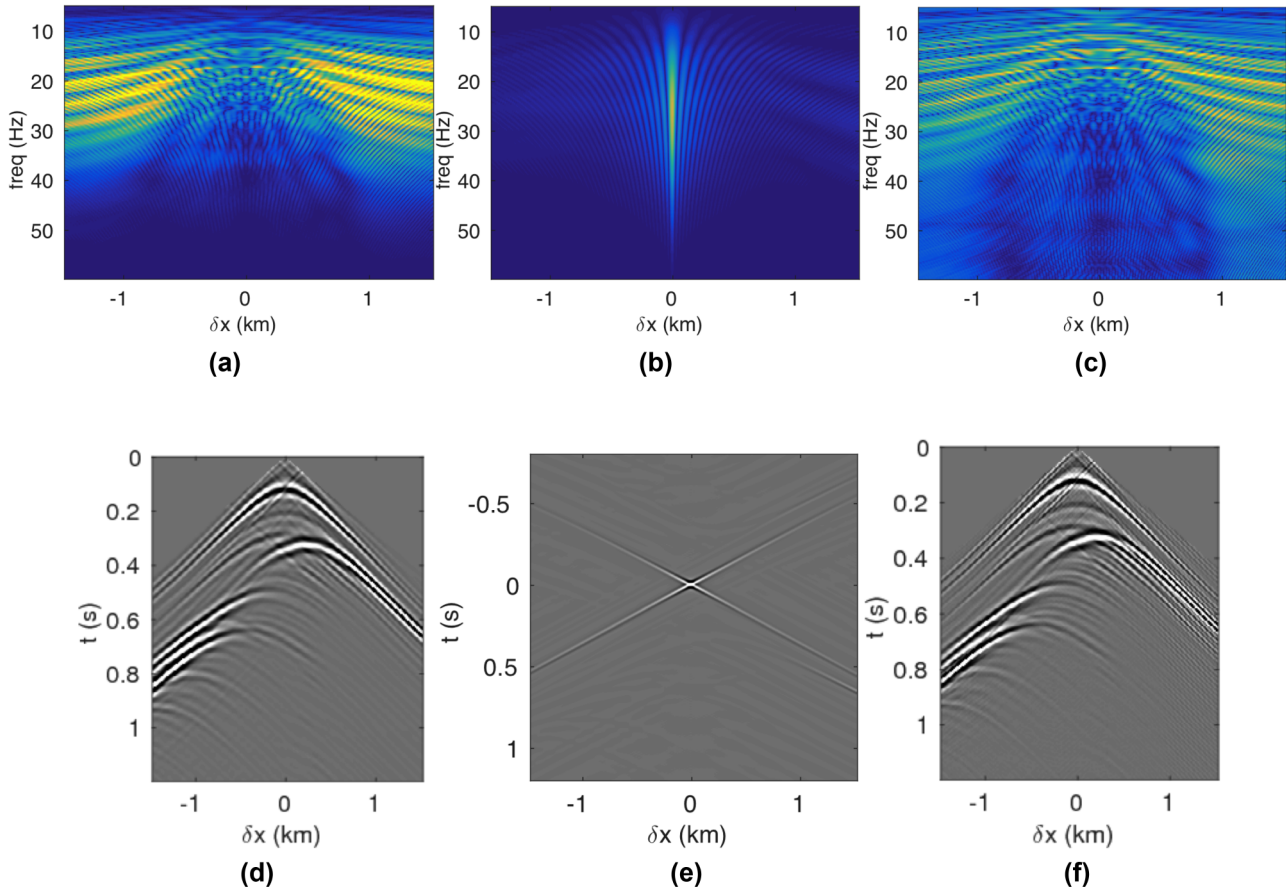


Figure 4 Redatumed data for the velocity model with a fault in the frequency-space domain (top row) and time-space domain (bottom row), with respect to the horizontal lag δx at the datum point $x = 3$ km and $z = 2.5$ km (on the dashed white line in Fig. 3 a): Subsection from $\tilde{\mathbf{E}}$ in (a,d), from $\tilde{\mathbf{I}}$ in (b,e) and \mathbf{R} matrices displayed in (c,f). We can clearly see that the reflection events are more focused in \mathbf{R} compared to $\tilde{\mathbf{E}}$.

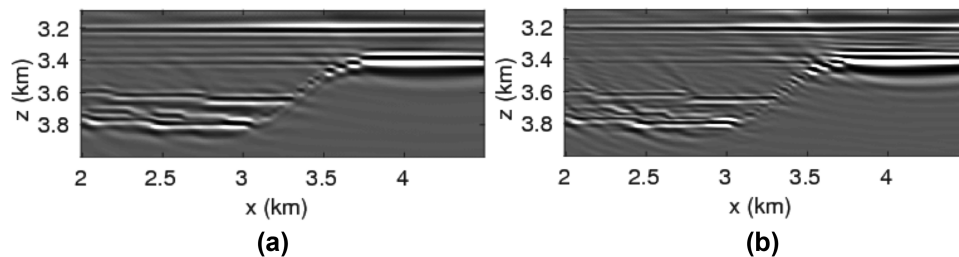


Figure 5 Reverse-time migration at a given datum in case of velocity model with fault close to the zone of interest: (a) without and (b) with blurring correction. We can see that the reflectors after blurring correction are sharper. However, MDD creates ringing artifacts that can be reduced by using regularization while solving the least-squares problem for the reflection response.

when using spatially impulsive monopole source, in which case it corresponds to the identity matrix, the singular values of \mathbf{D} typically decay quickly especially at the low to mid frequencies. This property allows us to factorize the extended image volume into a r -rank matrix with $r \ll N_s$ for the lower frequencies. For higher frequencies, the singular values of \mathbf{D}

decay less rapidly, which eventually leads to an upper bound smaller or equal to N_s for the image volumes.

To illustrate that image volumes indeed permit low-rank factorizations at lower frequencies, we again consider a small but geologically complex subset of the Marmousi model, which consists of strongly dipping reflectors with strong

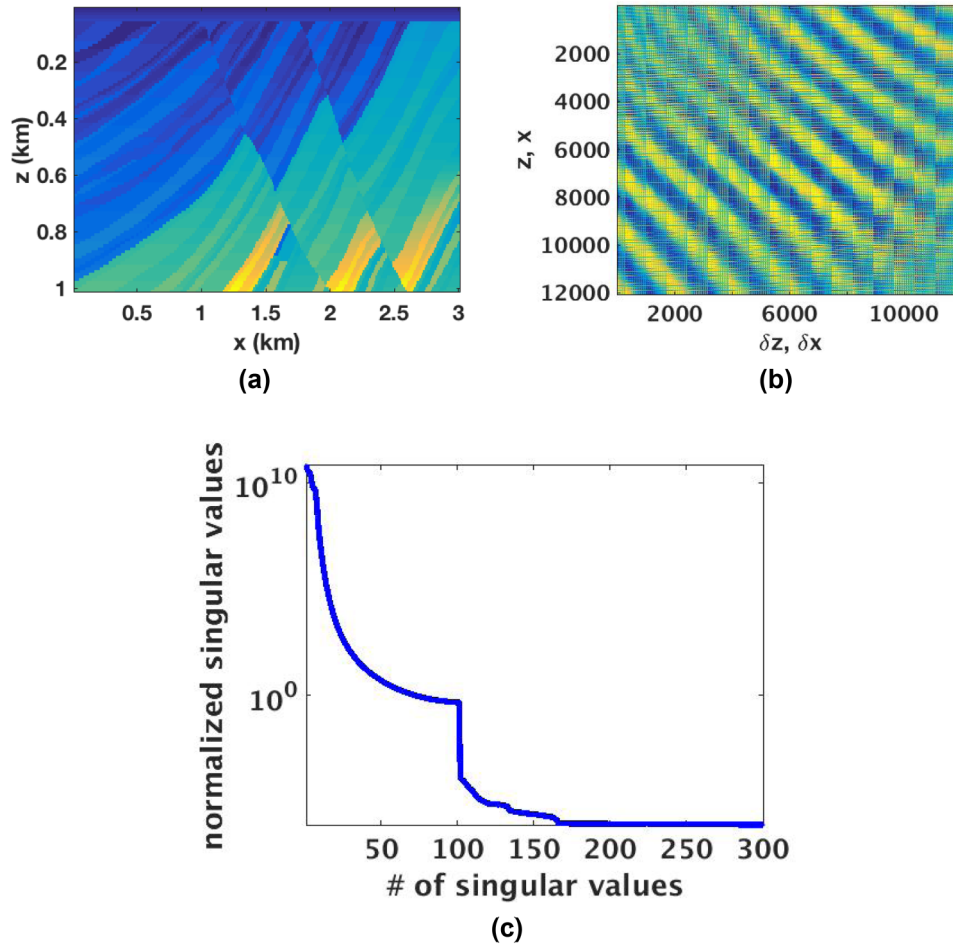


Figure 6 Low-rank approximation of full subsurface extended image volume: (a) a subsection of the Marmousi model, which is highly complex in nature with steeply dipping events, (b) corresponding full-subsurface extended image volume E at 5 Hz, (c) decay of the singular values that illustrates the low-rank nature of E . We can clearly see the fast decay of singular values and the signal-to-noise ratio of the approximated image volume is about 50 dB when we only select top 10 singular vectors.

lateral velocity variations (Fig. 6 a). We work with this small subset so that we can compare the time and storage requirements to an explicitly formed full subsurface extended image volume. Even for these complex velocity models, we observe in Fig. 6(c) a fast decay for the extended image at 5 Hz depicted in Fig. 6(b). This observation is crucial for our approach because it demonstrates that even for complex velocity models full subsurface extended image volumes exhibit a low-rank structure, which allows us to approximate the original image volume at 50 dB by only using the top 10 singular values.

From the example of Fig. 6, we can clearly see that even for complex geological structures, the full subsurface extended image volume exhibits a low-rank behaviour, a property we can exploit to compress the full subsurface extended image volume in its factorized form. We do this by only computing

the top r singular vectors of E . While such an approach has obvious benefits some important limitations of the proposed approach remain, namely (i) we need to solve $2N_x$ partial differential equation (PDE) to form and store full subsurface image volume, which is computationally prohibitively expensive, and (ii) the cost of singular value decomposition (SVD) of a dense explicit full subsurface image volume is of the order of $\mathcal{O}(N_x^3)$ using naive SVD algorithms (Holmes, Gray and Isbell 2007).

To circumvent these prohibitive computational costs, we propose to use randomized SVDs (Halko, Martinsson and Tropp 2011) to compute our low-rank approximation of full subsurface extended image volumes. Randomized SVDs (see Algorithm 2) derive from probedings and have a computational time of $\mathcal{O}(N_x^2 \times r)$, which can lead to significant gains if r is

small. As we can see from Algorithm 2, the computational costs of randomized SVDs are dominated by the costs of evaluating PDEs in lines 1 and 3. Aside from these expensive PDE solves, Algorithm 2 also involves carrying out a QR factorization and an SVD. However, the costs associated with these operations are small because they involved operations on probed quantities – that is $\mathbf{Y} \in \mathbb{C}^{N_x \times r}$ and $\mathbf{Z} \in \mathbb{C}^{r \times N_x}$ are both tall and thin matrices that require a total of only $4r$ (cf. $2 \times N_x$ PDE solves for conventional methods that loop over shots) PDE solves to evaluate. In addition to obvious computational gains, our SVD-based matrix factorization also allows us to store and form (through actions) extended image volumes while retaining most of the energy in \mathbf{E} .

Mainly, Algorithm 2 consists in two stages. The first stage includes steps 1–3 and uses probing to build an orthonormalized basis of the range of the extended image volume (EIV). Typically, we start with r Gaussian random vectors \mathbf{X} and get r basis vectors $\tilde{\mathbf{B}}$ to describe most of the energy of the EIV, that is associated to the most significant singular values of \mathbf{E} . Then, in the second stage, from step 4 to step 6, we compute the singular value decomposition of a much smaller matrix $\tilde{\mathbf{Z}} = \tilde{\mathbf{B}}^* \mathbf{E}$ that inherently contains most of the energy of the EIV. Finally, we get a low-rank approximation of the EIV, consisting of two $N \times r$ matrices for the singular vectors and a real vector with r entries for the singular values. This low-rank representation is now storable, in comparison with the $N \times N$ EIV, and faster to compute than the conventional techniques.

Algorithm 2. Low-rank SVD algorithm (Halko, Martinsson and Tropp 2011) applied for image-volumes

Inputs: source matrix \mathbf{Q} , data matrix \mathbf{D}

Outputs: singular value decomposition of extended image volumes \mathbf{G} , \mathbf{S} , and \mathbf{F}

1. Generate a $(N_x \times r)$ Gaussian random probing matrix \mathbf{X}
 2. Form $\tilde{\mathbf{Y}} = \mathbf{E}\mathbf{X}$, where \mathbf{E} is a function of \mathbf{Q} and \mathbf{D} (matrix-free operation like in Algorithm 1)
 3. Construct an orthonormal basis $\tilde{\mathbf{B}}$ of range of \mathbf{E} from the QR-decomposition on $\tilde{\mathbf{Y}}$
 4. Form $\tilde{\mathbf{Z}} = \tilde{\mathbf{B}}^* \mathbf{E}$ (Transpose of matrix-free operation like Algorithm 1)
 5. Compute the singular value decomposition of the $r \times N_x$ matrix $\tilde{\mathbf{Z}}$: $[\mathbf{G}, \mathbf{S}, \mathbf{F}] = \text{svd}(\tilde{\mathbf{Z}})$
 6. Update $\mathbf{G} \leftarrow \tilde{\mathbf{B}}\mathbf{G}$
-

Note that, we observed the same low-rank behaviour for the matrix Γ , hence, we also exploit its low-rank struc-

ture in the multi-domain experimental section using the randomized SVD algorithm, where we replaced \mathbf{E} with Γ in Algorithm 2. Given the low-rank matrix factorization form for \mathbf{E} and Γ , that is $\mathbf{E} \approx \mathbf{G}_1 \mathbf{S}_1 \mathbf{F}_1^*$, and $\Gamma \approx \mathbf{G}_2 \mathbf{S}_2 \mathbf{F}_2^*$, we can extract any column from \mathbf{E} and Γ via probing with basis vectors $\mathbf{w}_j = [0, \dots, 0, 1, 0, \dots, 0]$, that is $\tilde{\mathbf{E}} \approx \mathbf{G}_1 \mathbf{S}_1 \mathbf{F}_1^* \mathbf{w}_j$ and $\tilde{\Gamma} \approx \mathbf{G}_2 \mathbf{S}_2 \mathbf{F}_2^* \mathbf{w}_j$. As mentioned before, here, \mathbf{w}_j corresponds to the grid location of the point scatterer at datum ∂D_d . Finally, our randomized SVD-based multi-domain target-oriented imaging consists of the following main steps: (i) compute the low-rank factorization form of \mathbf{E} and Γ using Algorithm 2; (ii) extract columns from \mathbf{E} and Γ corresponding to the grid location of the point scatterer at datum ∂D_d ; (iii) perform multidimensional deconvolution to get reflection response matrix \mathbf{R} .

Remark. Randomized SVDs, derived from actions of \mathbf{E} and \mathbf{E}^* are not the only option to factorize \mathbf{E} using the top r singular vectors. Krylov methods (Larsen 1998; Liu, Wen and Zhang 2013) can accomplish the same decomposition. However, randomized SVDs-based methods have at least two main advantages over Krylov-based methods. First, the matrix-vector multiplication to form $\tilde{\mathbf{Y}}$ and $\tilde{\mathbf{Z}}$ can be performed in parallel. This observation allows us to take full advantage of parallel and distributed machines, especially in cloud-computing environment. Second, randomized SVDs have enabled the development of highly efficient single-pass matrix factorization algorithms (Martinsson 2019), where the communication cost is low and we never have to store the matrices explicitly. For detailed comparison of Krylov methods and randomized SVDs, we refer the interested readers to Halko, Martinsson and Tropp (2011) and Martinsson (2019). Because of the above-mentioned reasons, we opt for randomized SVDs in this work.

4.1 Target-oriented imaging using randomized SVD

To demonstrate the computational and memory benefits of our approach with randomized singular value decomposition (SVD) over the classical approach to approximate full subsurface image volume, we compare computation times and memory use for a subset of the Marmousi model, sampled at 10 m. The simulated data consist of 301 coincident sources and receivers both sampled at 20 m. As before, we use a Ricker wavelet with a central frequency of 30 Hz. To show the benefits over the full bandwidth of the underlying signal, we compute monochromatic full subsurface image volumes for eight frequencies ranging from 5 to 50 Hz.

Table 2 Signal-to-noise ratio between the full extended image volume and the probed extended image volume using randomized SVD, for eight different frequencies

Frequency(Hz)	Real Rank	Full Time(s)	Estimated Rank	Reduced Time(s)	SNR(dB)	Memory Ratio
5	301	490	7	11	29	43
10	301	490	8	11	24	38
15	301	490	10	12	21	30
20	301	490	11	12	20	28
25	301	490	13	13	22	24
30	301	490	16	15	22	19
40	301	490	20	17	22	15
50	301	490	25	21	20	12

Notes: Comparison of the computational times and memory storage ratio in both cases. Clearly, for this small 2D example extracted from the Marmousi velocity model, the saving in computational time and memory is highly significant: with a factor of more than 25 in time and more than 15 in memory storage.

As we mentioned before, the rank of the extended image volume is indeed intrinsically linked to the data matrix from Formula (7), hence, we compute r by looking at the singular values of data. Although, we get an approximation of rank r using data matrix, this is not the optimal way of choosing the rank since the singular values decay of the data matrix is slower than the corresponding decay of the image volume. By keeping only the number of singular values higher than 10% of the largest one, we obtain the approximated rank values for each frequency. We then use this estimated rank value for each frequency and solve Algorithm 2 to compute the low-rank factorization form of the full subsurface extended image volumes. For comparison, we also form the full subsurface extended image volume using the classical approach. As summarize in Table 2, randomized SVD-based approach reduces the computational time of forming the full subsurface image volume by a factor of approximately 25. This major improvement in computational time is due to a reduction in the number of partial differential equation solve ($2r \ll 2N_x$). Moreover, we only need to store \mathbf{G} , \mathbf{S} and \mathbf{F} matrices, thus, the storage of the extended image volumes using randomized SVD is $r \times (2N_x + 1)$ (Algorithm 2). Even though the estimated rank increases with the frequency in Table 2, it is still much smaller than the number of shots needed to form the full extended image volume. Finally, we compare the low-rank approximation of the extended image volume with the full extended image volume by computing the signal-to-noise ratio that is above 20 dB for all eight frequencies. Thus, the approximation of full subsurface extended image volume using randomized SVD is satisfactory.

To illustrate the figures in Table 2, we propose to compare the resulting reverse time migration (RTM) images computing either with conventional multi-dimensional

deconvolution (MDD) or with the randomized SVD algorithm. We obtain the RTM image by extracting the main diagonal of the extended image volume (EIV). In the case of conventional MDD, we actually use the Dirac probing counterpart, which allows us to build the full EIV. As the chosen velocity model is small, we are still able to compute and store the full EIV. We display the RTM images for frequency 10 Hz in Fig. 7(a) and for 40 Hz in Fig. 7(c). For the comparison, we also extract the RTM image from the low-rank approximation of the EIV computed with Algorithm 2, at 10 Hz in Fig. 7(b) and 40 Hz in Fig. 7(d) Clearly, the low-rank approximation of the EIV gives satisfactory reconstitution of the RTM image. Even though the values on the colour scale are not relevant, we would like to stress that they are the identical at fixed frequency for a better visual comparison.

Finally, we demonstrate our approach in a more realistic setting, namely a subsection of the Sigsbee2A model (The SMAART JV 2001), which contains complex sedimentary basins interspersed with a high-contrast and high-velocity salt body. The velocity model depicted in Fig. 8 is 6 km deep and 7.5 km wide and sampled at 5 m. In total, we acquire 1502 coincident sources and receivers sampled at 5 m. As before, we use a Ricker wavelet as a source function with a central frequency of 30 Hz. To compute the reflection response matrix \mathbf{R} , we use frequencies from 5 to 55 Hz with a sampling interval of 0.125 Hz. We perform the target-oriented imaging at the subsurface datums ∂D_d represented by the black line in Fig. 8(a).

Figure 9(a) shows the estimated rank values as a function of the frequencies, which we compute by keeping the number of singular values of data higher than 10% of the largest one. We use these rank values in our randomized SVD Algorithm 2 to approximately factorize \mathbf{E} and $\mathbf{\Gamma}$. Figure 9(b) displays the

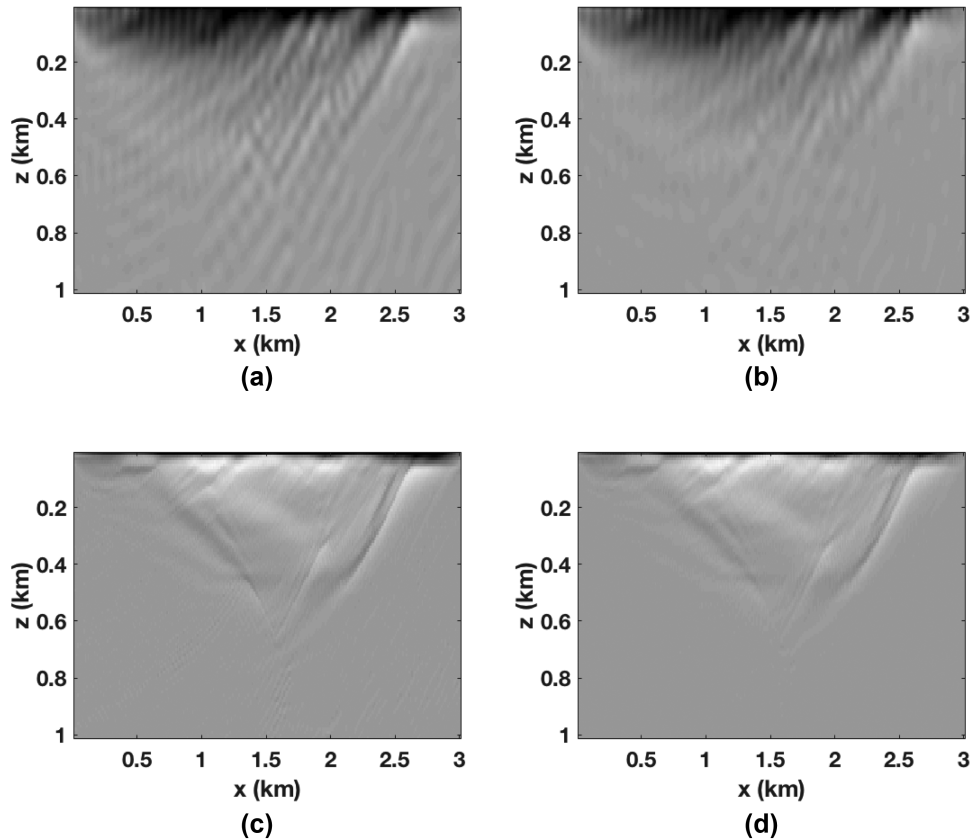


Figure 7 Extraction of the RTM image from the diagonal of the EIV for the subsection of the Marmousi velocity model 6a, at 10Hz (top row) and at 40Hz (bottom row). We compare the resulting RTM image obtained by conventional MDD or Dirac probing (on the left) with the RTM image extracted of the low-rank representation computed with the randomized SVD algorithm (on the right). Clearly, the low-rank approximation of the EIV gives satisfactory reconstitution of the RTM image. Even though the values on the colour scale are not relevant, we would like to stress that they are the identical at fixed frequency for a better visual comparison.

total number of PDE's solves needed to compute this approximation as a function of increasing frequency and compares it to the conventional method, which requires N_s PDE solves on top of computing cross-correlations between the forward and adjoint wavefield prior to summing over time. From this approach, it is evident that our randomized SVD method is computationally efficient both in its required number of PDE solves (r versus N_s) and cross-correlation costs (cross correlations in data versus model space). However, it is not unexpected that the numerical efficiency decreases as we move to higher frequencies. Monochromatic seismic data exhibit low-rank structure only at the low and intermediate frequencies, but less so at the high-frequencies, that is above 60 Hz. This behaviour can be explained because wavefields become increasingly oscillatory at high frequencies. In addition to these computational gains, we save a factor of approximately 15 in memory usage, while forming the full subsurface extended

image volumes using the proposed factorization-based approach. Apart from memory savings, Figure 10 shows the gain in the computational time to form full subsurface extended image volume using the proposed randomized SVD approach compared to the shot-profile-based extended imaging approach. The memory storage and the computational efficiency make the randomized SVD-based approach a true enabler for performing the target-oriented imaging for large-scale 3D acquisition.

Figure 11(a–c, d–f) displays the frequency-space domain and time-space domain data before and after source-side blurring correction, for the subsurface datums ∂D_d represented by the black line. We again perform the inverse Fourier transform along the frequency axis to go from frequency-space domain to time-space domain. In Fig. 11(c), we can observe that energy is distributed evenly along lower and higher end of the frequency spectrum of \mathbf{R} compared to the spectrum of

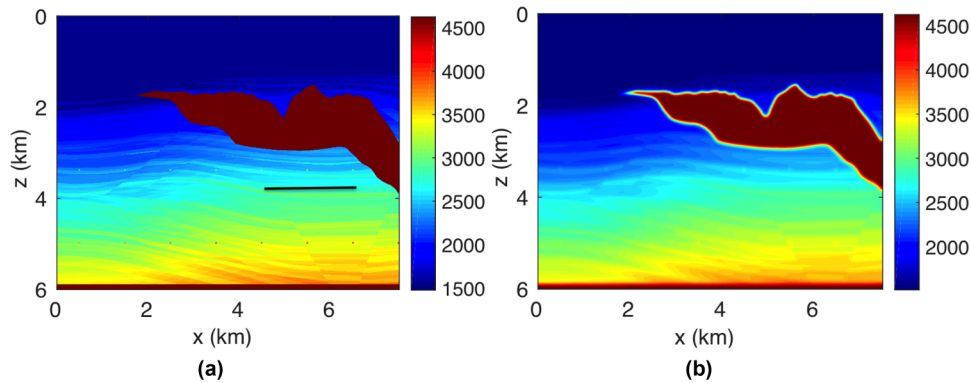


Figure 8 Sigsbee2A velocity model (The SMAART JV 2001): (a) true and (b) smooth velocity model (m/s). Position of the subsurface datum (black line) is overlay on velocity model (a).

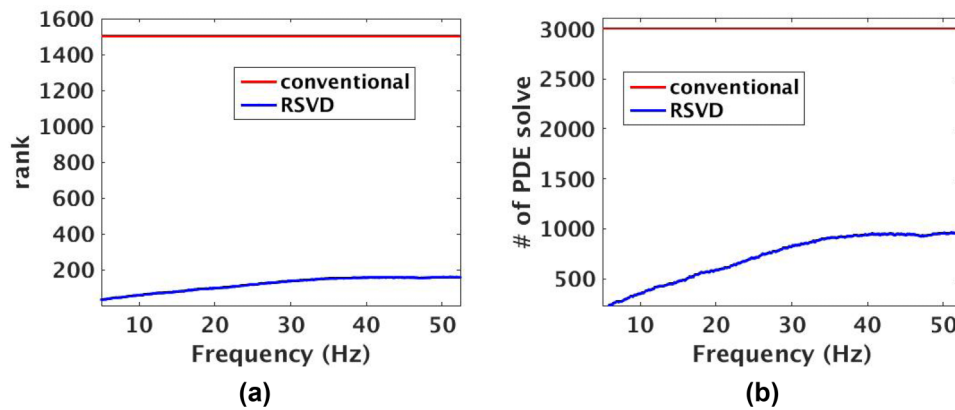


Figure 9 (a) Estimated rank values as a function of frequency for the Sigsbee2A velocity model. (b) Number of PDEs solve as a function of frequency to form the full subsurface reflection response matrix \mathbf{R} using randomized SVD approach and comparison to the classical approach.

E in Fig. 11(a). The well balanced energy distribution along the spectrum results in more sharper (high-resolution) reflection events as evident in time-domain sections (cf. Fig. 11 d,f). Figure 12(a,b) shows the reverse-time migration images before and after our blurring correction. From these Figures, we see that the increase in the energy distribution across the frequencies after the blurring correction improves the resolution of reflectors.

5 DISCUSSION

The main objective of target-oriented wave equation-based imaging is to reduce demand for computational and storage resources to create images in localized areas. While many of the proposed approaches to redatuming make claims to this effect, the majority of the approaches still require expensive loops over shots that involve simulations of forward and adjoint wavefields over the full domain. In most cases, reported gains can be attributed to the use of target-based imaging

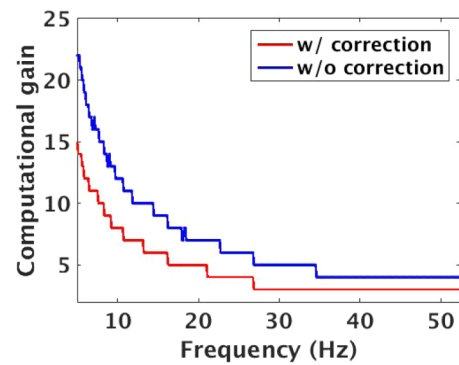


Figure 10 Computational gain in forming full reflectivity matrix $\tilde{\mathbf{R}}$ with (blue line) and without (red line) blurring corrections using the randomized SVD framework. Note that, to compute $\tilde{\mathbf{R}}$ using randomized SVD, we need to solve $4r$ PDEs without blurring correction and $6r$ PDEs with blurring correction.

conditions over the targeted domain of interest, which leads to savings in compute and storage of the wavefields.

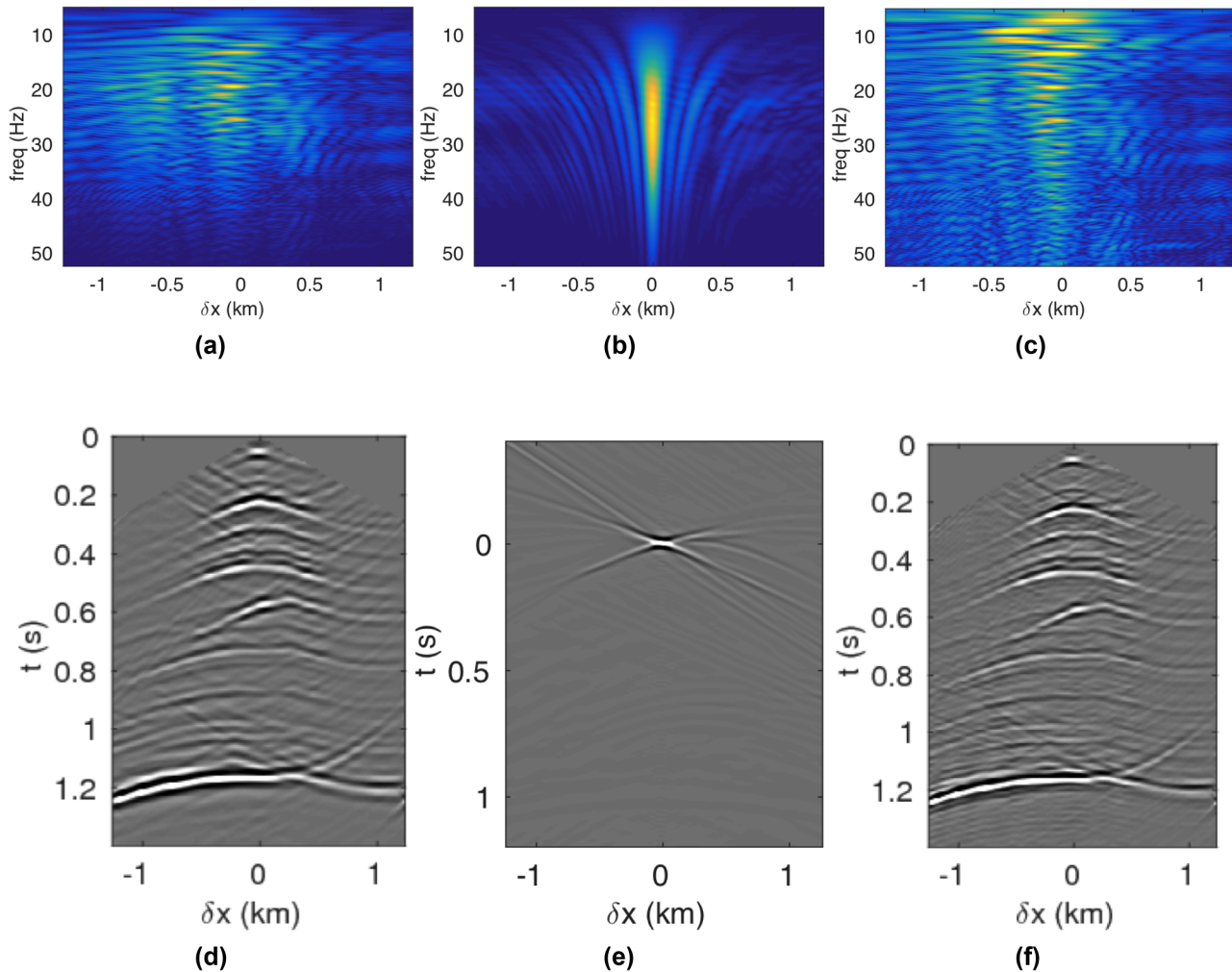


Figure 11 Redatum data for the Sigsbee2A model at the subsurface datum point $x = 2$ and $z = 0.6$ km, in frequency-space domain (top row) and in time-space domain (bottom row), with respect to the horizontal lag δx at the datum represented by the black line in Fig. 8(a): subsection from $\tilde{\mathbf{E}}$ in (a,d), from $\tilde{\mathbf{F}}$ in (b,e) and \mathbf{R} matrices displayed in (c,f). Here again, reflection events are more focused after the blurring correction.

By using techniques from randomized linear algebra, we avoid the expensive loops over shots, explicit storage of the forward and adjoint wavefields, and computations of cross-correlations that involve these wavefields. Instead, we work with low-rank factorizations of full-subsurface image volumes. We reap information needed to form these factorizations via actions of extended image volumes on random probing vectors that act as subsurface simultaneous sources. By using the wave equations themselves to carry out the heavy lifting, for example we avoid computing the cross correlations, we achieve significant computational and storage gains as long as we can approximate these image volumes by a low-rank matrix. In that case, the cost of our redatuming scheme is dominated by the rank of the factorization instead of by

the main loop over shots while storage of and operations on full wavefields are avoided. As a result, we end up with a computationally feasible and scalable scheme that by virtue of the low-rank singular-value decomposition also allows us to cheaply apply the deconvolution imaging condition.

We verified the viability of our approach in two-dimensional (2D) for monochromatic subsurface-offset volumes, which as all extended full-subsurface image volumes, are quadratic in the image size making it impossible to form them explicitly except for the smaller 2D toy problems. Experimentally, we found that an order of magnitude reduction in memory usage and computation time is achievable for the low to mid frequencies where image volumes can well be approximated by a low-rank matrix. In that regime, our method scales

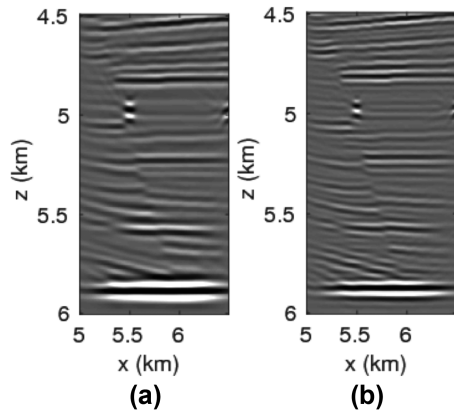


Figure 12 Reverse-time migration in case of the Sigsbee2A model at the subsurface datum represented by the black line in Fig. 8: (a) without and (b) with blurring correction. We can see that the reflectors after blurring correction are sharper.

to three-dimensional where even bigger computational gains are to be expected as we reported in van Leeuwen, Kumar and Herrmann (2017). In that case, we could never hope to form the image volumes explicitly so we have to completely rely on our approach based on randomized probing where the cost per probing vector corresponds to only two wave-equation solves. As a result, we gain significantly if the extended image volumes can well be approximated in low-rank form since the costs will scale with this rank.

While we argue that our approach beats conventional approaches, as long as the rank is smaller than the number of shots, we make this claim for the low-to-mid range frequencies where extended image volumes are indeed low rank. For the higher frequencies, we may have to exploit other approaches that include exploiting the low-rank off-diagonal structure of image volumes by hierarchical semi-separable representation (Chandrasekaran *et al.* 2006; Lin, Lu and Ying 2011; Jumah and Herrmann 2014). Another option would be to use power iterations in the randomized singular value decomposition as suggested in Halko, Martinsson and Tropp (2011) and Martinsson (2019) or in the recent work of Musco and Musco (2015) on block Krylov methods. In either case, the presented approach has the advantage that the probing can be done in parallel and involves significantly fewer resources because we only need to loop over the rank instead of over shots.

6 CONCLUSIONS

The main premise of wave-equation-based redatuming, also known as target-oriented imaging, is that it is computationally

efficient. Even though many elegant solutions to target-oriented imaging exist, their computational costs often scale linearly with the number of shots and exponentially with the size of the full model. By using probing techniques from randomized linear algebra, we mitigate these costs by representing image volumes in low-rank factored form. By construction, image volumes can be considered as redatumed data after restriction to a particular datum. The low-rank factorization gives us access to this restriction without having to form, work with or loop over all shots.


Our approach is based on a ‘double two-way wave equation’, which is an object we can probe at the costs of two wave-equation solves for each probing vector. These probeds give us access via the randomized singular-value factorization to redatumed data without the need to loop over shots, which can lead to major computational gains if the rank of the factorization is smaller than the number of shots. Since monochromatic full subsurface extended image volumes indeed exhibit a low-rank structure for low-to-mid range frequencies this leads to a viable approach that leads to significant reductions in computational costs. Rather than being dominated by the number of shots in the survey, our costs are proportional to the rank of the factorization, which is small as long as we do not work at too high frequencies. Experimental results for these low-to-mid range frequencies on synthetic two-dimensional (2D) models demonstrate that the proposed low-rank factorization-based approach is approximately 25 times faster in time and more than 15 times cheaper in memory storage compared to the conventional shot-profile approach. We further illustrate experimentally using 2D Sigsbee model that the target-oriented images using low-rank factorization-based reflection response can be computed with little to no compromise in accuracy compared to the standard shot-profile approaches. However, we could observe that the blurring correction, even though making the reflectors sharper, creates ringing effects. We believe that we could consider a curvelet-based sparsity promotion approach to perform the multi-dimensional deconvolution (Herrmann and Wang 2008; van der Neut and Herrmann 2013) in order to remove some of these artifacts.

Finally, our experimental results on complex geological models demonstrate the efficiency of the proposed methodology even for large 2D velocity models. We believe that our methodology should allow us practical reflection-based extended imaging in term of computational time and storage capacity for large-scale five-dimensional seismic data from three-dimensional velocity models.

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