

Meetings

Terrestrial nitrogen cycling in Earth system models revisited

'Terrestrial nitrogen cycling in Earth system models' Workshop, Dartington Hall, Devon, UK, February 2016

Understanding the degree to which nitrogen (N) availability limits land carbon (C) uptake under global environmental change represents an unresolved challenge. First-generation 'C-only' vegetation models, lacking explicit representations of N cycling, projected a substantial and increasing land C sink under rising atmospheric CO₂ concentrations. This prediction was questioned for not taking into account the potentially limiting effect of N availability, which is necessary for plant growth (Hungate *et al.*, 2003). More recent global models include coupled C and N cycles in land ecosystems (C–N models) and are widely assumed to be more realistic. However, inclusion of more processes has not consistently improved their performance in capturing observed responses of the global C cycle (e.g. Wenzel *et al.*, 2014). With the advent of a new generation of global models, including coupled C, N, and phosphorus (P) cycling, model complexity is sure to increase; but model reliability may not, unless greater attention is paid to the correspondence of model process representations and empirical evidence. It was in this context that the 'Nitrogen Cycle Workshop' at Dartington Hall, Devon, UK was held on 1–5 February 2016. Organized by I. Colin Prentice and Benjamin D. Stocker (Imperial College London, UK), the workshop was funded by the European Research Council, project 'Earth system Model Bias Reduction and assessing Abrupt Climate change' (EMBRACE). We gathered empirical ecologists and ecosystem modellers to identify key uncertainties in terrestrial C–N cycling, and to discuss processes that are missing or poorly represented in current models.

'... future experiments should be specifically designed to identify constraints set by resource availability and trade-offs caused by intrinsic physiological limitations.'

Openness of the nitrogen cycle

We started by addressing how N cycling varies at large spatial scales. Nitrogen cycle openness can be quantified as the ratio between 'new

N' inputs and total internal N cycling (see Fig. 1). While consensus appears to suggest an increase in N cycle openness from boreal to tropical and in particular Savannah ecosystems (Cleveland *et al.*, 2013), models and empirical estimates diverge on the strength of this gradient (Sönke Zaehle, MPI-BGC, Germany).

These quantifications hinge partly on estimates of biological N fixation, for which a recent top-down analysis suggests a global rate of 58 Tg N yr⁻¹ (Vitousek *et al.*, 2013) – much lower than the earlier estimate by Cleveland *et al.* (1999) of 100 to 290 Tg N yr⁻¹. New measurements in tropical forests (Sullivan *et al.*, 2014) are consistent with this downward correction. However, remarkably large uncertainties remain. Sasha Reed (US Geological Survey, USA) emphasized the surprising variety of N-fixing organisms that have only recently been discovered in different habitats (Reed *et al.*, 2011) and the high degree of variability in fixation rates with stand age in tropical forests (Batterman *et al.*, 2013). Several participants pointed to remaining challenges in measuring N fixation rates in the field and, in particular, the likely underestimation of free-living N fixation in current estimates. For example Elbert *et al.* (2012) estimated N fixation in cryptogamic covers alone at 49 Tg N yr⁻¹. Moreover, Taraka Davies-Barnard (University of Exeter, UK), Andy Wiltshire (MetOffice, UK), and Karin Rebel (Utrecht University, the Netherlands) noted that from a modelling perspective, it is essential not only to quantify the magnitude of the N fixation flux, but also to understand its controls and its energetic (C) cost. Indeed, mechanisms have been identified that imply a (limited) plant control over N fixation rates (Menge & Hedin, 2009), including labile C export by plants to free-living N-fixing heterotrophs (Reed *et al.*, 2011). Finally, John Raven (Dundee University, UK) discussed key insights into the physiological processes of N fixation.

Sarah Cornell (Stockholm Resilience Centre, Sweden) pointed out that the other major component of new N inputs, atmospheric deposition, may also be underestimated. Deposition of organic N forms has generally been ignored, but can be substantial in remote areas due to their longer atmospheric lifetimes and transport ranges compared to NO_x (Neff *et al.*, 2002). In view of the increasing appreciation of organic N in plant nutrition, as noted by Torgny Näsholm (Umeå Plant Science Centre, Sweden), a re-assessment of N cycle openness in boreal systems may be required.

Adrien Finzi (Boston University, MA, USA) presented a global compilation of mineralization rates and compared these with the fraction of total gross primary productivity (GPP) allocated belowground. The derived pattern reveals a high C-cost of P acquisition but a low apparent-C cost of N acquisition in the tropics, consistent with 'open' N cycling. By contrast, Ivan Janssens and Sara Vicca (both University of Antwerp, Belgium) reported low biomass production efficiency (ratio of biomass production to GPP) in tropical forests, and suggested this reflects high C costs of nutrient acquisition. The spatial scale of the studies varies

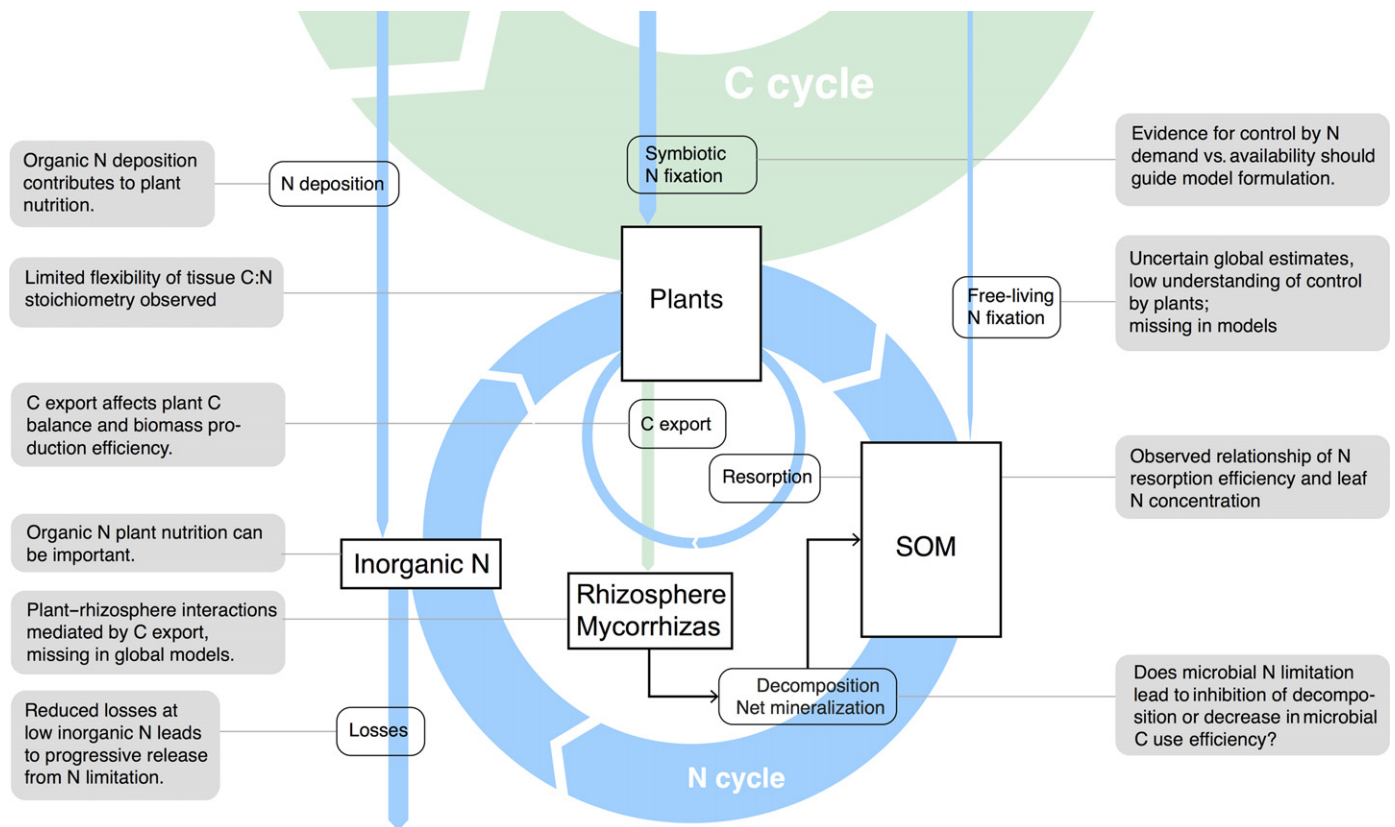


Fig. 1 Summary of findings and research priorities emerging from the workshop 'Terrestrial nitrogen cycling in Earth system models'. The thickness of arrows qualitatively illustrates the relative magnitude of fluxes. Nitrogen (N) cycle openness is illustrated by the relative flux magnitudes of losses or inputs (arrows for N deposition, symbiotic N fixation, and free-living N fixation) vs internal cycling (circles). SOM, soil organic matter.

substantially, and various possibilities exist to explain this discrepancy (higher rates of autotrophic respiration among others). Additional research is necessary to reconcile these two reports. Their studies do, however, point to large scale gradients of N cycle openness, total belowground allocation, and biomass production efficiency as prime benchmarks for a new generation of C–N models.

Nitrogen constraints on plant CO₂ responses

The positive effect of elevated CO₂ on leaf-level photosynthesis implies a shift in the balance of supply and demand for C vs N. We went on to discuss respective effects on plant–soil interactions. Sönke Zaehle reported that current C–N models, in contrast to measurements in Free Air CO₂ Enrichment (FACE) experiments, achieve a positive response in net primary productivity (NPP) primarily by increasing C : N at the tissue level (Zaehle *et al.*, 2014); however, the models fail to reproduce the observed shift towards belowground C allocation and increased N uptake. Ivan Janssens reported that leaf N concentrations across European forests is co-determined by phylogeny and, while spatial differences in stoichiometry within plant species can be large, temporal responses to environmental change appear small (Sardans *et al.*, 2015). It became clear that stoichiometric flexibility in different plant tissues and leaf-level photosynthetic capacity – as well as their response to changes in atmospheric CO₂ and soil nutrient availability – warrant special attention from empiricists, as well as in the formulation of models.

While overestimating the role of stoichiometric flexibility, current C–N models allow limited scope for shifts in C allocation when N availability is low. Recent research has highlighted the importance of mycorrhizas (Phillips *et al.*, 2013) and plant-controlled rhizosphere priming effects (Cheng *et al.*, 2014; Finzi *et al.*, 2015). Oskar Franklin (IIASA, Laxenburg, Austria) pointed out that mycorrhizal associations may aggravate N limitation under low N availability due to their N immobilization, thereby reducing a positive CO₂ fertilization effect (Franklin *et al.*, 2014). But Sara Vicca and César Terrer (Imperial College London, UK) showed contrasting new results from a meta-analysis of experimental data. Under low N, a positive biomass production response under elevated CO₂ does appear in plants associated with ectomycorrhizal fungi, but not in plants associated with arbuscular mycorrhizas. A similar mechanism may be at work for N-fixing species. Tuula Larmola (Natural Resources Institute, Helsinki, Finland) reported that in an environment with very low N deposition, *Sphagnum*-associated N fixation may far exceed atmospheric deposition and thus enable sustained C accumulation in peatlands (Larmola *et al.*, 2014). Sönke Zaehle showed results of a global modelling study indicating that assumptions regarding the controls on N fixation have major implications for its response to rising CO₂. He posed the question 'How do we put the brakes on N fixation?'. But interestingly, the implications of resolving N fixation controls over the global land C balance are limited (Meyerholt *et al.*, 2016). This may be linked to the fact that C–N models that resolve the

dynamics of a soil inorganic N pool, and simulate N loss as a function of the pool's size, suggest that even in the absence of a strong amplification of N fixation there is a progressive release from N limitation on a decadal timescale, due to increased ecosystem N retention (Walker *et al.*, 2015).

In view of the C costs for N fixation and subsidising rhizosphere activity and mycorrhizas, Adrien Finzi asked the question: 'Is N limitation just C limitation in disguise?'. Successfully modelling plant C allocation is key to predicting effects of environmental change. Simple models based on optimality principles have been proposed (Franklin *et al.*, 2012) and may guide the development of a next generation of C–N models. Along the way, it will be important to make good use of data from manipulation experiments to test models at the level of individual processes. Conversely, future experiments should be specifically designed to identify constraints set by resource availability and trade-offs caused by intrinsic physiological limitations.

The representation of soil nitrogen cycling in models

The CENTURY paradigm for modelling transformations of C and N in soils received a mixed response from participants. It was agreed that this modelling approach has stood the test of time, above all as a means for modelling the effects of land-use changes on the dynamics of soil C. However, Ivan Janssens noted that the observed response to N deposition is a reduction of heterotrophic respiration relative to NPP (Janssens *et al.*, 2010); opposite to what CENTURY-type models predict. Consistent with reduced respiration with increasing N availability, Stefano Manzoni (Stockholm University, Sweden) drew attention to evidence that the C-use efficiency of microbes may decrease with increasing organic matter C : N (Manzoni *et al.*, 2012). This supports an increasingly widespread view that more explicitly representing microbial function in soil models will be required to explain these observations and adequately account for plant–soil interactions.

In this respect, and in general, appropriate compromises have to be found in the dilemma between model realism and tractability. This should be guided by key observational phenomena required to be captured by models, data availability for model evaluation, and the principle that everything should be made as simple as possible, but not simpler.

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