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AMOC Stabilization Under the Interaction With Tipping Polar Ice Sheets

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Key Points:

- A conceptual model of interacting Atlantic Meridional Overturning Circulation (AMOC), Greenland Ice Sheet and West Antarctica Ice Sheet (WAIS) is presented
- Interactions between these tipping elements strongly modifies the stability of the whole system
- A collapse of the WAIS can prevent tipping of the AMOC

Supporting Information:

Supporting Information may be found in the online version of this article.

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Abstract Several large-scale components of the climate system may undergo a rapid transition as critical conditions are exceeded. These tipping elements are also dynamically coupled, allowing for a global domino effect under global warming. Here we focus on such cascading events involving the Greenland Ice Sheet (GIS), the West Antarctica Ice Sheet (WAIS) and the Atlantic Meridional Overturning Circulation (AMOC). Using a conceptual model, we study the combined tipping behavior due to three dominant feedbacks: the marine ice sheet instability for the WAIS, the height-surface mass balance feedback for the GIS and the salt-advection feedback for the AMOC. We show that, in a realistic parameter range of the model, a tipping of the WAIS can inhibit cascading events by preserving the AMOC stability.

Plain Language Summary In the climate system, the interaction of specific components known as tipping elements are thought to be able to induce a global domino effect, or cascading tipping. In this study, we present a conceptual model containing the most strongly interacting components, namely the Atlantic Meridional Overturning Circulation (AMOC), the Greenland Ice Sheet and the West Antarctica Ice Sheet. We find that the stability of this system as a whole is strongly modified when interactions are included. Especially, while a Greenland Ice Sheet collapse destabilizes the AMOC, the model shows that a collapse of the West Antarctica Ice Sheet might prevent a global cascading event by stabilizing the AMOC.

1. Introduction

Global warming is one of the main threats to the stability of the present-day climate system. Under this warming, specific climate system components might change abruptly when certain critical thresholds are exceeded. Examples of such tipping elements (Lenton et al., 2008) are the Greenland Ice Sheet (GIS), the Atlantic Meridional Overturning Circulation and the West Antarctic Ice Sheet (WAIS). A thorough understanding of the mechanisms and impact of tipping behavior in these subsystems is fundamental in assessing the risks of climate change (Kemp et al., 2022).

Tipping elements are also strongly interacting, for example, the polar ice sheets and the ocean circulation, and hence tipping in one subsystem (the leading system) may lead to tipping in another (the following system), in a so-called tipping cascade (Dekker et al., 2018). This rises the possibility of domino effects, causing the climate system to collapse while the threshold of one subsystem only has been crossed (Klose et al., 2021). However, the collapse of one subsystem may also stabilize others (Ciemer et al., 2021; Swingedouw, Fichefet, Huybrechts, et al., 2008; Weaver et al., 2003) and thereby prevent such dramatic scenarios.

Using expert elicitation, Kriegler et al. (2009) qualitatively assessed the risk of such cascading events in a context of global warming. In a more quantitative assessment, Wunderling et al. (2021) studied the interactions between tipping of the GIS, the AMOC, the WAIS and the Amazon rainforest using a highly idealized model of coupled dynamical systems, each capturing the tipping through back-to-back saddle-node bifurcations. Here, the GIS, AMOC, and WAIS stood out as the protagonists of a potential large-scale cascading. However, the Wunderling et al. (2021) approach lacks a connection to the underlying physical processes, and their interactions.

The aim of this study is to couple physically motivated conceptual models of the three tipping elements. Within a new coupled model, we study similar issues as Wunderling et al. (2021), where the GIS and AMOC were described respectively as potential initiator and mediator of cascading, while the role of the WAIS was less certain. We focus on the conditions under which cascading can occur or not, and especially on regimes in which the AMOC can remain stable when interacting with tipping polar ice sheets under global warming.

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2. Modeling Coupled Tipping Elements

A conceptual inter-hemispheric model composed of the GIS, the AMOC and the WAIS subsystems is presented in Figure 1. The individual model components and their coupling are described in the below paragraphs.

2.1. The GIS

Over the last decades, satellite measurements have revealed a significant acceleration of ice loss of the GIS (The IMBIE Team, 2020), where the decreasing surface mass balance (SMB) plays a crucial role (Enderlin et al., 2014; Goelzer et al., 2013). A critical global mean surface temperature increase threshold of 0.8–3.0°C has been suggested (Armstrong McKay et al., 2022), above which the GIS would be committed to melting. An important mechanism to destabilize an ice sheet is the height-SMB feedback (Levermann & Winkelmann, 2016), according to which the thinning of an ice mass enhances melting as its surface reaches lower altitudes, associated with higher temperatures. Based on early warning signals, Boers and Rypdal (2021) claim that the height-SMB feedback might already have brought parts of the western GIS close to a tipping point.

To represent the GIS, we consider an isothermal ice sheet lying on a fixed bedrock (Greve & Blatter, 2009). The evolution of the ice thickness is given by the contribution of the transport inside the ice dome involving the ice flux, along with the SMB. The problem is simplified by using the shallow-ice approximation and considering a radially symmetric ice cap resting on a flat circular bed at sea level, with a no-ice condition at the boundary. The height-dependent SMB is defined using the precipitation rate and equilibrium line altitude, which depend on the regional temperature anomaly $\Delta\tau_N$ with respect to the present-day annual mean value. Finally, ice loss is converted to a meltwater flux F_N directly inserted in the northern Atlantic box. More details about the GIS model are provided in Section S1 in Supporting Information S1.

2.2. The AMOC

From long-term observations of sea surface temperature, it has been suggested (Caesar et al., 2018, 2021) that a slowing down of the AMOC has occurred over the last century. Global warming and associated changes in the hydrological cycle are overal destabilizing (Bakker et al., 2016) due to the salt-advection feedback. A tipping point ranging from 1.4°C to 8.0°C of global warming is suggested in the literature (Armstrong McKay et al., 2022), although with low confidence level. Also, an increased freshwater input in the deep water formation region, caused by GIS melting, is destabilizing (Jackson & Wood, 2018). Based on global climate models, Jackson and Wood (2018) have suggested a critical extra freshwater input of about 0.1 Sv, corresponding to the high end of that associated with a GIS decay (Lenaerts et al., 2015). The impact of freshwater input in the southern region, however, remains uncertain as there are numerous competing feedbacks (Swingedouw, Fichefet, Goosse,

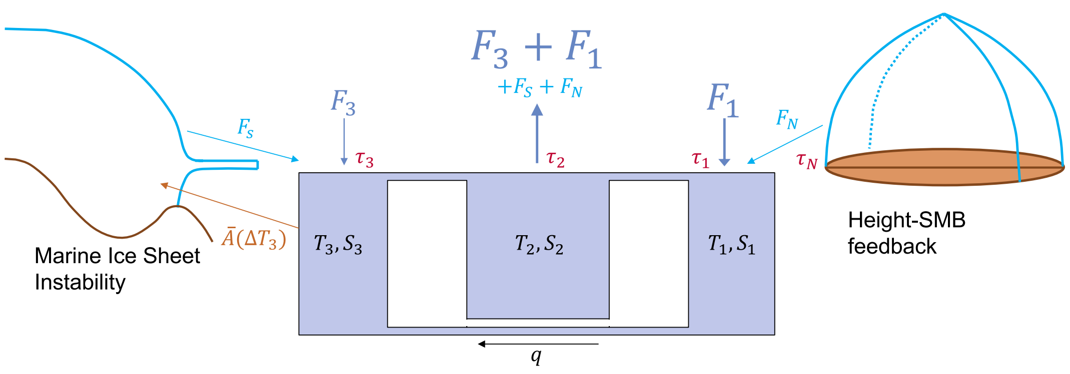

Figure 1. Representation of the coupled model. The West Antarctic Ice Sheet (WAIS) is represented by a single marine ice sheet in the Antarctic region. The Atlantic Meridional Overturning Circulation is depicted by three boxes for the southern (under 30°S), tropical (30°S to 30°N) and northern (above 30°N) Atlantic Ocean, each one coming with their own temperatures and salinities, forced by precipitation fluxes $F_{1,3}$ and background temperatures $\tau_{1,2,3}$. The Greenland Ice Sheet is represented by a radially symmetric ice dome in the Arctic region. Both ice sheets interact with the ocean through meltwater fluxes $F_{N,S}$, and the southern Atlantic Ocean temperature T_3 interacts with the WAIS through the depth integrated ice viscosity parameter $\bar{A}(\Delta T_3)$.

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& Loutre, 2008), potentially restraining the AMOC weakening under global warming (Swingedouw, Fichefet, Huybrechts, et al., 2008) or leading to an AMOC recovery (Weaver et al., 2003). Recently, based again on early warning indicators, Boers (2021) claims that the AMOC is close to tipping.

For the AMOC, we use the three-box model of Rooth (Lucarini & Stone, 2005; Rooth, 1982; Scott et al., 1999), describing the AMOC driven by the pole-to-pole density difference. The first box represents the northern Atlantic Ocean, the second the tropical region and the third the southern Atlantic Ocean. Temperatures and salinities are changed through advective transport due to the AMOC strength q, defined positive for a present-day, northern sinking configuration. The temperature $T_{1,2,3}$ of each box is relaxed to a background temperature $\tau_{1,2,3}$, at a relaxation time scale of about 25 years. Salinities $S_{1,2,3}$ are forced by surface freshwater fluxes $F_{1,3,N,S}$, including precipitation and meltwater input at the poles, compensated by evaporation in the tropics (see Figure 1), yielding conservation of total salt content for the Atlantic Ocean. More details about the AMOC model are provided in the Section S2 in Supporting Information S1.

The stability of the Rooth model in a northern sinking state under varying freshwater or temperature forcing has already been investigated (Lucarini & Stone, 2005; Scott et al., 1999). On one hand, at a total freshwater input in the northern box of $F_{1,c} = 0.86$ Sv, the model undergoes a subcritical Hopf bifurcation above which only the southern sinking state remains stable, while increasing the freshwater input in the southern box strengthens the circulation due to the salt-advection feedback. On the other hand, increasing the inter-hemispheric forcing temperature asymmetry $\tau_1 - \tau_3$ weakens the circulation. In both cases however, the associated critical values will be highly rate dependent, which will be discussed in Section 3.

2.3. The WAIS

The WAIS has seen unprecedented ice loss over the last decades (The IMBIE Team, 2018), with ocean warming being the main driver (Favier et al., 2019; Joughin et al., 2014; Shepherd et al., 2004). The increased loss is likely due to the fact that a dominant part of the WAIS ice mass is grounded under sea level, making it subject to a dynamical instability known as the marine ice sheet instability (MISI), resulting in a fast retreat of the grounding line when on retrograde bed slope (Mulder et al., 2018; Schoof, 2007; Weertman, 1974). In the Amundsen sea sector, the MISI might already be initiated (Favier et al., 2014; Rignot et al., 2014), with potentially dramatic consequences for the WAIS (Feldmann & Levermann, 2015) and for the whole Antarctic continent (Garbe et al., 2020). In terms of global warming, Armstrong McKay et al. (2022) propose a tipping point ranging between 1.0°C and 3.0°C.

We consider the WAIS as one single marine ice sheet (Schoof, 2007) under depth-integrated shallow-shelf approximation, represented by a rapidly sliding, two-dimensional and symmetric marine ice sheet. A floating ice shelf is included as boundary condition at the grounding line, such that the position of the grounding line can be tracked. We consider the SMB constant and uniform, ignoring any melting contribution, as we expect dynamical ice loss to dominate when the MISI occurs. The bifurcation structure of this model with respect to the depth-integrated ice viscosity parameter $\bar{A}$ is known (Mulder et al., 2018; Schoof, 2007) and consists of two back-to-back saddle-node bifurcations inducing the MISI, resulting in a fast retreat of WAIS as this parameter exceeds the critical value of $\bar{A}_c = 2.87 \cdot 10^{-25}$ Pa^{-3}s^{-1}. In the coupled model, we consider $\bar{A}$ as a linear function of the southern Atlantic Ocean temperature anomaly ΔT_3

$$\bar{A}(\Delta T_3) = \frac{\bar{A}^0}{T_3^0}\left[T_3^0 + c_S \Delta T_3\right]. \quad (1)$$

where c_S is a non-dimensional coupling parameter and the parameters $\bar{A}^0$ and T_3^0 indicate values at reference state, translating into a critical value $\Delta T_{3,c}$ decreasing as c_S increases. Although no straightforward link can be established between T_3 and the regional ocean temperature, let us note that the range $c_S = [0.1, 1.0]$ corresponds to the range $\Delta T_{3,c} = [0.1, 1.0]$ °C, consistent with model projections for the regional ocean warming with respect to present-day likely to trigger a WAIS tipping (Garbe et al., 2020). Finally, ice loss is converted into a meltwater flux F_S, from which we assume only a fraction $f = 0.27$ to enter the southern Atlantic Ocean, considering the rest to be lost in the Pacific Ocean. More details about the WAIS model and the estimation of f are provided in the Section S3 in Supporting Information S1.

3. Results

In this section, we will systematically use the initial state such that the AMOC is in a stable northern sinking configuration similar to present-day (Lucarini & Stone, 2005), and with ice sheets yielding realistic present-day values for ice volumes and meltwater fluxes (see Section S4 in Supporting Information S1). To investigate the coupled model under global warming, we linearly increase surface temperatures over the GIS and Atlantic Ocean during 100 years, after which temperature is held constant, that is, for $j \in \{N, 1, 2, 3\}$ (with t in years),

$$\tau_j(t) = \tau_j(0) + \gamma_j \frac{\Delta\tau_2}{100} t. \quad (2)$$

where amplification parameters γ_j are used to represent the phenomena of polar amplification (Cai et al., 2021; Hahn et al., 2021; Holland & Landrum, 2021), here with respect to the equatorial warming $\Delta\tau_2$. Those are estimated from results of Hahn et al. (2021), where many CMIP5 and CMIP6 models were used and compared to assess the (zonally averaged) amplification as a function of latitude when forced by a CO_2 quadrupling, and chosen to be $\gamma_N = 2$, $\gamma_1 = 1.3$, $\gamma_2 = 1.0$, and $\gamma_3 = 1.0$ for respectively the North Pole and the three oceanic boxes. For those values, the forcing can be expressed in terms of the global warming with respect to the pre-industrial period $\Delta\tau_G \approx 1.1\Delta\tau_2 + 1.1$ alone, obtained by averaging over the Earth's surface and accounting for the present-day warming.

To determine whether cascading occurs or not, we first focus on the AMOC when no ice sheets are involved or, in other words, when $c_S = \gamma_N = 0$. In this case, applying the forcing (2), we find a critical value $\Delta\tau_{G,c} = 9.2°$C at which the AMOC destabilizes, thereby tipping to the southern sinking configuration. Next, we couple only the GIS to the AMOC, that is, setting $c_S = 0$. The critical value $\Delta\tau_{G,c}$, above which the AMOC destabilizes decreases to 6.9°C. Indeed, as the GIS in this model reaches its critical warming level already at $\Delta\tau_G = 1.8°$C, the AMOC is destabilized not only by rising temperatures but also by additional meltwater input into the northern box. We note that while the tipping point of the GIS sits in the middle of the range proposed by Armstrong McKay et al. (2022), the AMOC tipping point exceeds it by approximately one degree, which remains reasonable given the low confidence associated with this projection. In any case, this situation clearly represents a tipping cascade as both systems tip while only the critical threshold of the GIS has been crossed.

Finally, choosing non-zero values for c_S, we couple the WAIS to the system. We repeat the global warming experiments with $\Delta\tau_G = 7.1°$C for two different WAIS-coupling values, $c_S = 0.2$ and $c_S = 0.8$, best illustrating the different effects of the coupling. For this level of warming, the GIS systematically tips at about year 10, while T_3 is increased by approximately 5°C within the first 150 years, far above the critical value triggering the MISI for both c_S values.

In the case of low coupling ($c_S = 0.2$, Figure 2a), the WAIS tips at about year 30, and the resulting meltwater flux is not large enough to compensate for the destabilizing effect of freshwater input in the north. Hence, the AMOC tips at about 400 years, resulting in another drastic rise of T_3. However, the subsequent acceleration of the WAIS collapse happens too late, as the AMOC is then already in a reversed circulation regime. Higher coupling ($c_S = 0.8$, Figure 2b) results in a more abrupt WAIS collapse triggered earlier, at about year 10. In this case, the meltwater flux from the WAIS is strong enough to maintain the AMOC in a northern sinking configuration. It is worth noting however that, while the circulation shift has been avoided, the AMOC strength is committed to a long term decrease of about 20% due to global warming.

The cases in Figure 2 are shown as the red crosses in Figure 3a, where the final state of the AMOC is shown in part of the $(\Delta\tau_G, c_S)$ parameter plane. In the yellow region, the AMOC is destabilized to the southern sinking state while, in the blue region, it remains in a northern sinking configuration. As expected, the critical value of warming leading to AMOC tipping $\Delta\tau_{G,c}$ (the boundary between the yellow and blue region) increases with increasing c_S, that is, when the WAIS more strongly reacts to ocean warming. Over the c_S interval [0.1,1.0], $\Delta\tau_{G,c}$ ranges between 6.9°C and 7.4°C, and implied transition times for both ice sheets remain consistent with those proposed by Armstrong McKay et al. (2022). Hence, this creates the possibility of preventing a collapse of the AMOC under conditions for which the WAIS tips fast enough. Notably, this range of warming is above those implied by any SSP scenario for the next century (Lee et al., 2021), for which the AMOC would hence remain stable regardless of the value of c_S. However, the same experiment conducted using a longer temperature increase time of 300 years (Figure 3b) leads to a higher tipping point of $\Delta\tau_{G,c} = 10.1°$C for the AMOC alone. Consequently, the coupling to the ice sheets results in $\Delta\tau_{G,c}$ ranging between 7.6°C and 8.3°C, consistent with expected warming levels implied by both SSP5-8.5 and SSP3-7.0 extended scenarios after 300 years (Lee et al., 2021). Representative cases (red crosses in Figure 3b) are illustrated on the Figure S2 in Supporting Information S1. In both experiments, the length of the interval on which $\Delta\tau_{G,c}$ varies appears to linearly depends on the fraction f of the WAIS meltwater flux reaching the southern Atlantic Ocean (see Figure S1 in Supporting Information S1).

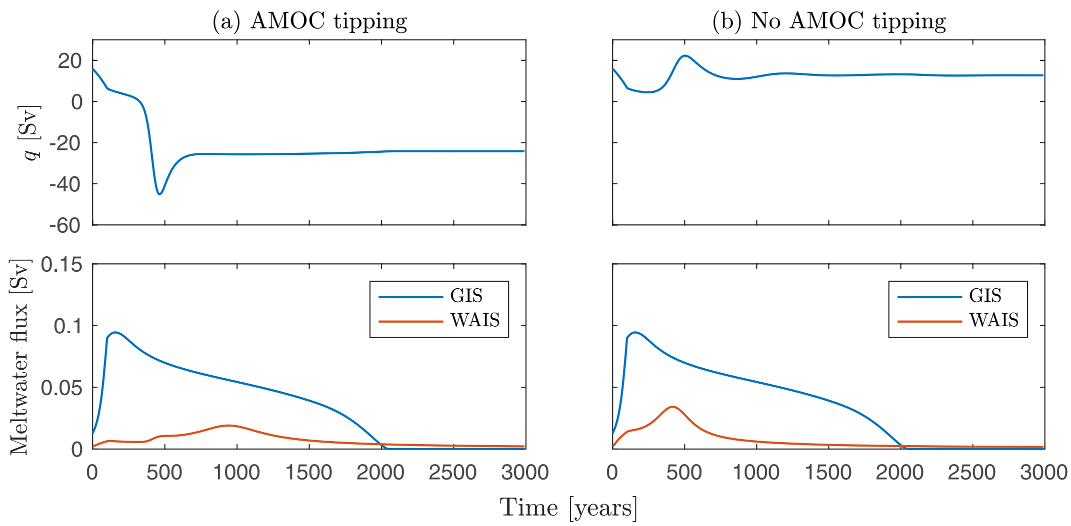

Figure 2. Transient behavior of the Atlantic Meridional Overturning Circulation strength q and the ice sheet meltwater fluxes under a linear global warming $\Delta\tau_G = 7.1°C$ lasting 100 years, for different couplings: (a) $c_S = 0.2$ and (b) $c_S = 0.8$.

In the global warming experiments so far, the destabilization of both ice sheets is induced within a few decades. Moreover, at initial state, all tipping elements are in equilibrium while in reality, some of them might already be engaged in a transient, for example, the GIS or WAIS. To gain more insight into the influence of the different delays and rates of change in the coupled system, we perform additional sensitivity experiments by forcing only the ice sheets, while the AMOC reacts solely to the implied meltwater fluxes. Similar to what is seen in so-called hosing experiments (Boulton et al., 2014; Jackson & Wood, 2018; Rahmstorf et al., 2005), we explore more extreme forcings and investigate which of those hypothetical scenarios results in an AMOC collapse.

First, we apply a linear increase of the regional surface temperature in Greenland τ_N lasting 100 years, and look for the critical value of $\Delta\tau_N$ leading to a southern sinking state of the AMOC. At the critical value of $\Delta\tau_{N,c} = 22.3°C$, the AMOC tipping occurs at a GIS melting totally in about 500 years. With this forcing, the GIS meltwater flux reaches 0.33 Sv at its peak, which is less than the forcing required to reach the Hopf bifurcation of the Rooth model. Hence, the AMOC collapse cannot be explained by bifurcation tipping. However, as the GIS collapses, the meltwater flux increases fast enough to trigger a rate-induced tipping (Ashwin et al., 2012).

Next, we add the WAIS to assess the stability of the AMOC when interacting with both polar ice sheets. To explore the combined effect of tipping rates and their delay in time, we force both ice sheets independently. At a time t we initiate a forcing of the GIS, linearly increasing τ_N by 23°C in 100 years. By choosing a slightly larger forcing than in the previous experiment, we reduce the potential AMOC stabilizing region occurring as a consequence of the WAIS tipping. After a time delay Δt, we initiate a forcing of the WAIS, applying a linear increase of T_3 by 7°C (affecting the WAIS only), in 100 years. Here, the exact value of T_3 increase is not crucial as the WAIS tipping response will anyway be determined by the coupling parameter c_S.

The final state of the AMOC in the parameter space (Δt, c_S) is shown in Figure 3c. Below $c_S \approx 0.1$ (hence above $\Delta T_{3,c} = 1.0°C$), the AMOC always tips whenever the WAIS forcing is initiated. In this case, no WAIS meltwater flux can stabilize the AMOC against the high GIS meltwater input. However, as the coupling parameter c_S increases, a region of stability appears (blue). In this region, the lowest values of c_S require a strongly negative time delay Δt to prevent the AMOC tipping. There, the slower WAIS tipping provides a lower but sufficiently sustained meltwater input, such that the peak of the MISI coincides with the fast GIS tipping. As c_S increases, the stabilizing region rapidly encompasses shorter delays, including positive ones from $c_S \approx 0.3$. Note however that, at strong coupling, a WAIS tipping triggered too soon will result in most of the WAIS meltwater content to be released too long before the GIS tipping. Finally, it appears that there is a critical time delay at about $\Delta t = 200$ years, from which no WAIS tipping can causally interfere with the destabilization of the AMOC, due to the strong hysteresis behavior of the Rooth model. Representative cases (red crosses in Figure 3c) are illustrated on the Figure S3 in Supporting Information S1.

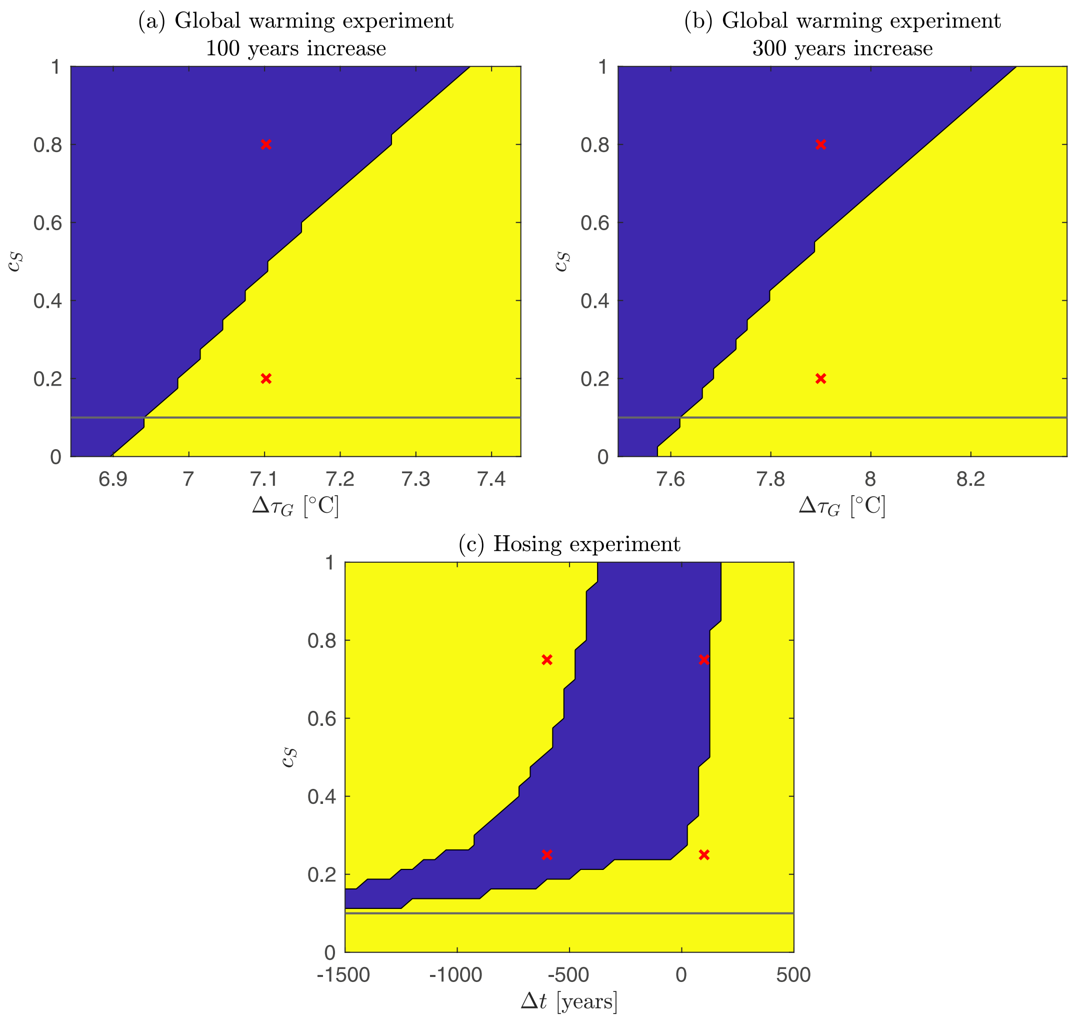

Figure 3. (a–b) Final state of the Atlantic Meridional Overturning Circulation (AMOC) depending on the warming $\Delta\tau_G$ and coupling parameter c_S, using a linear global warming lasting respectively (a) 100 years and (b) 300 years (c) Same for the hosing experiment but depending on the time delay Δt and coupling parameter c_S. The yellow area stands for reversed circulation (AMOC tipping), while blue area stands for northern sinking circulation (no AMOC tipping). The gray line indicates the lowest expected value for c_S at 0.1, corresponding to the critical ocean warming $\Delta T_{3,c} = 1.0°C$. Red crosses represent parameter configurations used in (a) Figure 2b (b) Figure S2 in Supporting Information S1 and (c) Figure S3 in Supporting Information S1.

4. Summary and Discussion

In this paper, we present a conceptual model to study the interaction of three tipping elements (WAIS, AMOC and the GIS) of the climate system. Under global warming, coupling the GIS to the AMOC drastically destabilizes the AMOC, making the GIS a potential initiator of global cascading as suggested by Wunderling et al. (2021). On the other hand, coupling the WAIS to the AMOC has a stabilizing effect on the AMOC, especially in the case of a relatively fast and early WAIS tipping. By considering two different warming time scales, our model suggests these competing feedbacks to be especially relevant regarding warming projections based on extended SSP5-8.5 and SSP3-7.0 scenarios (Lee et al., 2021).

By considering the stability of the AMOC when affected by meltwater fluxes only, we identified two key components to prevent an AMOC collapse, that is, interrupting a tipping cascade: the tipping rate of ice masses and the time delay between these tipping phenomena. While a comparatively slow tipping of the WAIS could keep the AMOC stable when triggered hundreds of years before the GIS tipping, it turns out that a faster WAIS tipping is more efficient to avoid an AMOC collapse for shorter delays, which is probably a more realistic scenario when thinking about climate change. In any case, our results rely on the fact that a freshwater input in the southern

Atlantic Ocean stabilizes the AMOC, a behavior which is shared by many box model representations of the AMOC (Cimatoribus et al., 2012; Rahmstorf, 1996; Rooth, 1982).

Of course, the model contains many idealizations and hence we argue below why we think these results are robust when more detailed physical processes are included. First, it is known that the stability of the AMOC in the Rooth model is very sensitive to the inter-hemispheric temperature forcing asymmetry, here implied by the amplification coefficients used to define climate change (Lucarini & Stone, 2005). While other choices of these parameters would affect the magnitude of the GIS and WAIS influence on the AMOC stability, we expect our results to remain robust as long as the warming remains destabilizing. A more accurate assessment of those amplification coefficients spanning the Atlantic Ocean alone would be an improvement to the quantitative results of our study.

Second, the description of the influence of the oceanic temperature on the WAIS has been strongly idealized. However, we can expect our qualitative results to hold as long as this interaction remains destabilizing. To better base it on physical grounds, one would have to consider sub-shelf melting and calving processes, interacting with the ice shelf stability through buttressing (Haseloff & Sergienko, 2018, 2022) and lateral drag (Schoof et al., 2017). Also, a better assessment of the fraction f of the WAIS freshwater flux reaching the southern Atlantic Ocean would be a direct improvement, which involves resolving the dynamics associated to the Antarctic Circumpolar Current and is beyond the scope of this study. Nonetheless, the apparent linear behavior of the critical warming with respect to f supports our results, as the stabilizing effect remains substantial when f varies around our estimation.

Third, some feedbacks have been omitted. Most importantly, the stabilizing effect of an AMOC tipping on the GIS via cooling of the northern hemisphere (Jackson et al., 2015) was not included, as our model has no atmospheric component. This represents a potentially strong negative feedback which may result in a safe overshoot of the GIS tipping point (Ritchie et al., 2021). Also, the mutually destabilizing effect of sea level rise (Gomez et al., 2010) on both ice sheets has been neglected, and is far more destabilizing for the WAIS than the GIS (Wunderling et al., 2021). In our model, both of those feedbacks would most probably inhibit the AMOC destabilization further, which makes it an interesting track to explore in future research.

In conclusion, the stability of the climate system, and in particular of the AMOC, is drastically changed when considering interactions between the tipping elements in agreement with the more abstract results of Wunderling et al. (2021). We emphasized here the consequences of a potentially stabilizing effect of a WAIS tipping on the AMOC in the presence of a tipping GIS, which could have important consequences on the other tipping elements and, by extension, on the climate system as a whole. Hence, while the collapse of the WAIS will always be a dramatic event, it might prevent a larger-scale cascading tipping event to happen. This stresses the importance of getting a better understanding of the interaction between the WAIS and the AMOC and to include the effects of interacting tipping elements in future climate change projections.

Data Availability Statement

All MATLAB codes are publicly available (Sinet, 2022), at the address: https://doi.org/10.5281/zenodo.6800055, including a short manual.

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