



LETTER

Novel, resistant microalgal polyethers: An important sink of organic carbon in the marine environment?F. GELIN,¹ I. BOOGERS,¹ A. A. M. NOORDELOOS,¹ J. S. SINNINGHE DAMSTÉ,¹ P. G. HATCHER,² and J. W. DE LEEUW¹¹Netherlands Institute for Sea Research (NIOZ), P.O. Box 59, 1790 AB Den Burg, The Netherlands²Pennsylvania State University, Fuel Science Program, University Park, PA 16802, USA

(Received December 14, 1995; accepted in revised form January 23, 1996)

Abstract—Five out of seven marine microalgal species investigated were found to biosynthesize nonhydrolysable, mainly aliphatic, biomacromolecules (algaenans). The molecular structure of the algaenan isolated from the microalga *Nannochloropsis salina* of the class Eustigmatophyceae was determined by solid state ¹³C NMR spectroscopy, Curie point pyrolysis–gas chromatography–mass spectrometry, and chemical degradations with HI and RuO₄. The structure is predominantly composed of C₂₈–C₃₄ linear chains linked by ether bridges. The algaenan isolated from a second eustigmatophyte (*Nannochloropsis* sp.) was structurally similar. Algaenans isolated from two chlorophytes also possess a strongly aliphatic nature, as revealed by the dominance of alkenes/alkanes in their pyrolysates. Accordingly, we propose that the aliphatic character of numerous Recent and ancient marine kerogens reflects selectively preserved algaenans and that these algaenans may act as a source of *n*-alkanes in marine crude oils.

1. INTRODUCTION

A wide variety of novel, insoluble, nonhydrolysable aliphatic biomacromolecules has recently been isolated from several lacustrine and terrestrial primary producers. These macromolecules have been recognised as algaenans in common, ubiquitous fresh-water green microalgae (Philp and Calvin, 1976; Berkaloﬀ et al., 1983; de Leeuw et al., 1991; Derenne et al., 1992b; de Leeuw and Largeau, 1993), as cutans and suberans in higher plant cuticles (Nip et al., 1986; Tegelaar et al., 1989a) and periderm tissue (Collinson et al., 1994), respectively, as tegmens in inner seed coats of fresh-water plants (van Bergen et al., 1994), as sporopollenins in spores and pollen grains (Guilford et al., 1988; Hemsley et al., 1993), and as polycadinenes in resins (van Aarssen et al., 1990). Because of their chemical stability and their resistance against bacterial decay, these macromolecules are selectively preserved (Zeliber et al., 1988; Nip et al., 1989) and make up a significant part of organic carbon in nonmarine sediments ranging in age from Recent to over 300 million years. They eventually can act as the source of high-wax crude oils upon burial and thermal cracking (Tegelaar et al., 1989b). Similarly, it has been proposed that the bulk of marine sedimentary organic matter (OM) derives from marine algaenans (Tegelaar et al., 1989c), but evidence for that has been rather circumstantial (Hatcher et al., 1983; Derenne et al., 1992c) and the existence of a marine algaenan was only shown for an uncommon monospecific chlorophyte (Derenne et al., 1992a). Accordingly, the possible occurrence of algaenans in five common marine green microalgae (Chlorophyceae) and two marine microalgae of the class Eustigmatophyceae has now been investigated.

2. Experimental**2.1. Culture Conditions**

All strains were obtained from the Culture Collection of the Natural Environment Research Council (U.K.). Algae were grown in batch cultures [in a medium described in Veldhuis and Admiraal (1987)] in a light/dark cycle of 16/8 h at 15°C. Cultures were kept axenic. Algaenans were isolated from the dry algal biomass by ultrasonic extraction with different organic solvents and by successive hydrolyses with KOH, HCl, and H₂SO₄. Extracts, saponified extracts, and organic fractions isolated after the KOH and HCl hydrolyses of the extracted algal biomass were analysed by GC and GC-MS.

2.2. Solid State ¹³C NMR

The NMR spectrum was obtained at a ¹³C frequency of 25.2 MHz using cross polarization with magic-angle spinning. Pertinent experimental parameters are the following: contact time = 1 ms, pulse delay = 1 s, acquisition length = 0.5 K and line broadening = 30 Hz. The chemical shifts scale is referred to tetramethylsilane at 0 ppm.

2.3. Curie Point Pyrolysis–Gas Chromatography–Mass Spectrometry

Samples were pyrolysed using a FOM-4LX Curie point pyrolysis unit directly coupled to a Hewlett Packard 5890 series II gas chromatograph fitted with a 25 m × 0.32 mm CP-Sil 5 (film thickness 0.12 μm) fused silica capillary column, temperature programmed from 0 to 320°C (hold time 10 min) at a rate of 3°C/min. Structural assignments were made based on MS and retention time data. MS conditions: VG Autospec Ultima instrument, ionization energy 70 eV, mass range 40–800, cycle time 1.8 s.

2.4. Chemical Degradations

Oxidation with RuO₄ followed the procedure described by Boucher et al. (1990). The acid products were derivatized with diazomethane to form methylester derivatives.

Table 1. Summary of the presence and nature of algaenans in the microalgae investigated.

Name	Algaenan ^a	Main series of pyrolysis products of algaenan ^b				Chemical nature
		<i>n</i> -alk-1-enes/ <i>n</i> -alkanes	<i>n</i> -alkan-2-ones, mid-chain ketones ^c	alkyl- benzenes	alkyl- phenols	
<i>Brachiomonas submarina</i>	-- ^c		--			--
<i>Chlorella spaerckii</i>	1.2	C ₇ -C ₄₁ (C ₁₄)	C ₇ -C ₂₂ (C ₁₉)	C ₆ -C ₁₀ (C ₇)	C ₆ -C ₉ (C ₇)	aliphatic/aromatic
<i>Chlorococcum</i> sp.	1.0	C ₇ -C ₃₄ (C ₂₆)	C ₇ -C ₂₅ (C ₁₉)	C ₆ -C ₁₀ (C ₇)	C ₆ -C ₉ (C ₇)	aliphatic/aromatic
<i>Nannochloris</i> sp.	~1	--	--	C ₆ -C ₁₀ (C ₇)	C ₆ -C ₉ (C ₇)	aromatic
<i>Stichococcus bacillaris</i>	-- ^c		--			--
<i>Nannochloropsis salina</i>	1-2 ^d	C ₇ -C ₃₆ (C ₁₆ , C ₂₈)	C ₇ -C ₂₂ (C ₁₈ , C ₁₉) C ₂₀ -C ₃₄ (C ₂₉)	C ₆ -C ₂₂ (C ₇)	--	aliphatic
<i>Nannochloropsis</i> sp.	1.5	C ₇ -C ₃₃ (C ₁₅ , C ₂₆)	C ₇ -C ₂₅ (C ₁₉) C ₂₂ -C ₃₂ (C ₂₆)	C ₆ -C ₂₂ (C ₇)	--	aliphatic

(^a) In % of the dry algal biomass. (^b) Carbon number range of the indicated series of the algaenans pyrolysis products. The maxima are indicated in brackets. (^c) The entire biomass was dissolved after HCl hydrolysis. (^d) The range indicated reflect 3 different batch cultures of this alga. (^e) Series of mid-chain *n*-alkanones (only present in the pyrolysates of the *Nannochloropsis* algaenans) are indicated below the series of *n*-alkan-2-ones.

Treatment with hydroiodic acid (HI) followed the procedure described by Panganamala et al. (1971). Iodide was displaced with sodium methyl thiolate in methanol for 18 h at room temperature to obtain thiomethyl ethers. The thiomethylether derivatives show a highly characteristic fragmentation pattern upon GC-MS analyses (Carlson et al., 1989).

3. RESULTS AND DISCUSSION

Freeze-dried algal cells, obtained from axenically grown batch cultures, were extracted and the residues were treated with base (KOH) and acids (HCl and H₂SO₄) to remove free lipids, ester-bound lipids, proteins, and carbohydrates, respectively. These treatments afforded 1-2% of an insoluble residue, i.e., algaenan, in the case of three chlorophytes, *Chlorella spaerckii*, *Chlorococcum* sp., and *Nannochloris* sp., and the eustigmatophytes *Nannochloropsis* sp. and *Nannochloropsis salina* (Table 1). The algaenan of the latter alga was analysed by solid state ¹³C NMR (Fig. 1). The chemical shift maximizing at 33 ppm indicates the highly aliphatic nature of this algaenan since at least 95% of the NMR response is represented by this chemical shift of CH₂ moieties.

The results of pyrolysis-gas chromatography-mass spectrometry (Py-GC-MS) of this algaenan are in full agreement with the NMR data. The pyrolysate of *N. salina* algaenan (Fig. 2a) is strongly dominated by homologous series of C₇ to C₃₆ *n*-alkanes and *n*-alk-1-enes maximizing at C₁₆ and C₂₈, revealing the presence of long polymethylene moieties in its algaenan. Series of ketones identified as co-eluting *n*-alkan-17-ones* and *n*-alkan-18-ones with up to 34 carbon atoms (Fig. 2b) indicate the presence of oxygen functionalities mainly at the C-17 and C-18 positions and the maximum carbon chain length, i.e., 34 carbon atoms, of the building blocks of the macromolecules. The presence of

such oxygen functionalities was confirmed by the highly specific distribution patterns of *n*-alkan-2-ones which maximize at C₁₈ and C₁₉ (Fig. 2b). The additional occurrence of mono-unsaturated mid-chain ketones with the unsaturation located in the shortest alkyl chain suggests the presence of another functionality, probably located at the first carbon atom. Based on earlier pyrolysis studies with model compounds and algaenans derived from freshwater microalgae, it

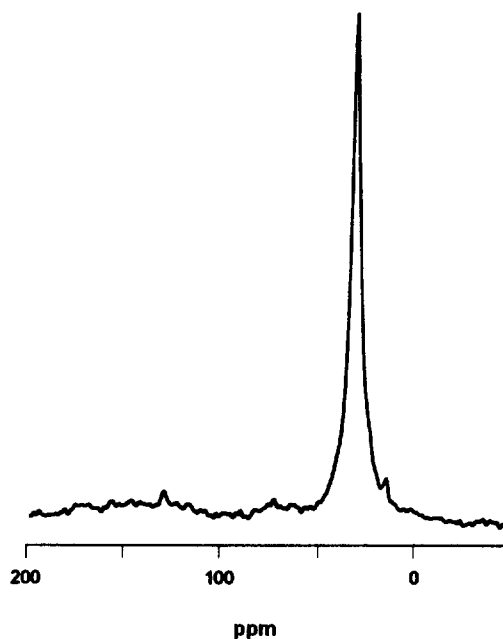


Fig. 1. Solid state ¹³C NMR spectrum of *N. salina* algaenan. The large peak at 33 ppm is assigned to long-chain polymethylene structures, the peak at 15 ppm to methyl groups, the broad region between 50 and 90 ppm is due to O-bound alkyl carbons, the region between 100 and 160 ppm corresponds to sp² carbons, and the broad resonance between 160 and 180 ppm is assigned to carbonyl or amide carbons.

* The nomenclature used for the mid-chain ketones is purposely not correct and should be, for example, C₃₀ alkan-13- and alkan-14-ones. The reason for using such numbering is due to the common, specific positions of the keto group for all the mid-chain ketones.

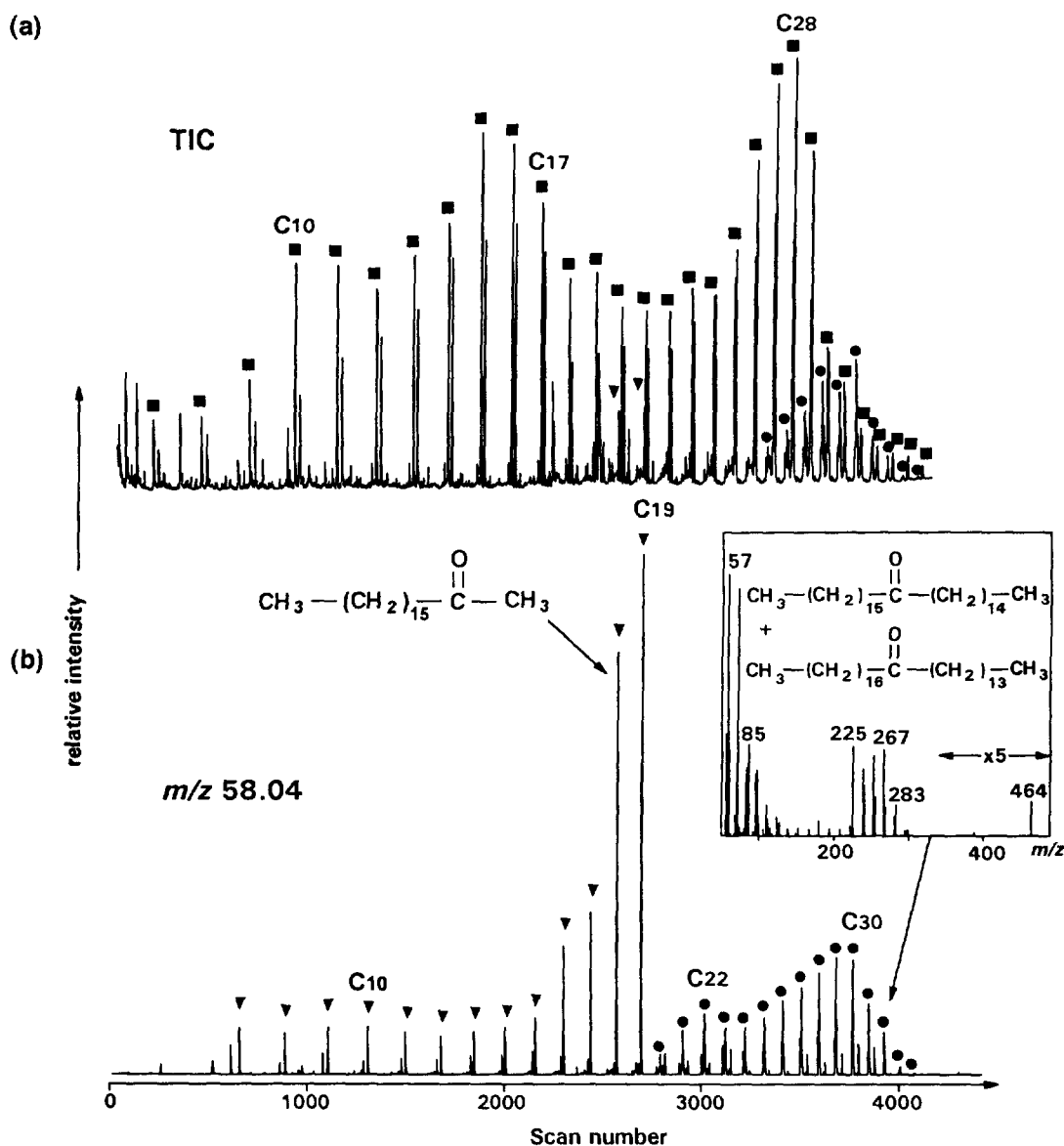


FIG. 2. Analysis of *N. salina* algaenan by Curie-point pyrolysis-gas chromatography-mass spectrometry; (a) total ion current trace (TIC) and (b) mass chromatograms of m/z 58.04 revealing the distribution of ketones. Filled squares indicate the n -alk-1-ene/ n -alkane doublets, filled triangles indicate the n -alkan-2-ones and filled circles indicate the mid-chain ketones. Numbers indicate the total numbers of carbon atoms. The insert in the bottom figure shows the mass spectrum of the coeluting C_{32} alkan-15- and -16-ones, identified based on MS data from previous studies (de Leeuw et al., 1981; Gatellier et al., 1993).

is concluded that both functionalities represent ether linkages (Gelin et al., 1993, 1994). This is confirmed by the presence of a small but recognizable peak for alkyl-O carbons in the ^{13}C NMR spectrum at approximately 70 ppm (Fig. 1).

Chemical degradation of the *N. salina* algaenan with ruthenium tetroxide (RuO_4) further supports the presence of ether bonds at the ω -17 or ω -18 positions and likely indicates ether bonds at the C-1 position as indicated by the generation of C_{24} - C_{34} ω -17- and ω -18-keto-alkanoic acids (Fig. 3a). A series of C_{24} - C_{34} ω -17,18-diketo-alkanoic acids was also produced by RuO_4 oxidation of this algaenan and reveals the presence of three ether linkages located at the C-1, ω -

17 and ω -18 positions of the polymethylene carbon chains (Fig. 3b). The third functionality is probably involved in cross-linking so that a three-dimensional network is built up, helping to explain the insolubility of the algaenan. The C_8 - C_{18} n -alkanedioic and C_{12} - C_{18} n -alkanoic acids which represent the major products of the RuO_4 degradation result from mid-chain oxidations at the vicinal ether positions (Fig. 3c,d). The long-chain fatty acids, maximizing at C_{29} , probably result from the oxidation of ether-linked alkyl moieties. The prevalence of ether-linked long-chain alkyl units in the *N. salina* algaenan structure was further demonstrated by the release of predominantly $1,\omega$ -17- and $1,\omega$ -18- C_{28} - C_{36}

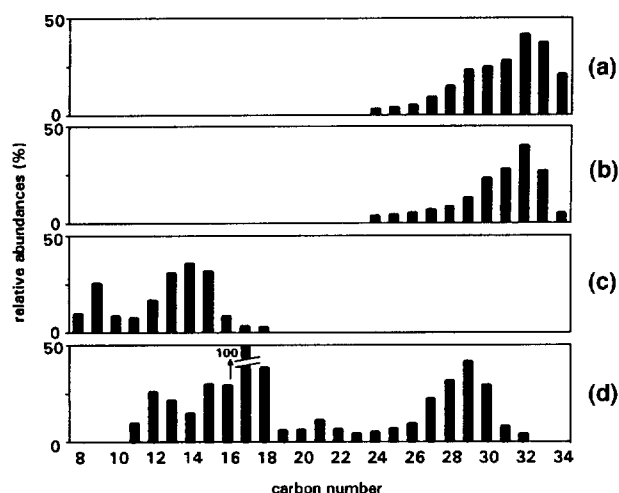


FIG. 3. Histograms displaying the major series of oxidation products derived from RuO_4 degradation of *N. salina* algaenan; (a) co-eluting n - ω ¹⁷- and n - ω ¹⁸-keto-alkanoic acids, (b) n - ω ^{17,18}-diketo-alkanoic acids, (c) n -alkanedioic acids, and (d) n -alkanoic acids.

diiodo-alkanes after hydroiodic acid treatment, known to selectively cleave ether bonds (Panganamala et al., 1971).

Pyrolysis of the algaenan of *Nannochloropsis* sp. showed very similar distributions of alkenes, alkanes, and alkanones in its pyrolysate, indicating that both eustigmatophytes biosynthesize structurally related algaenans (Table 1).

Analyses of the extracts of the two eustigmatophytes revealed the presence of free and ester-bound C_{30} – C_{32} alken-1-ols and C_{30} – C_{36} n -alkadiols with alcohol groups at ω -17 or ω -18 and C-1. These findings are in agreement with those of a previous study of the lipids of *N. salina* and other *Nannochloropsis* species (Volkman et al., 1992). The structural similarity of these lipids with units in the algaenan strongly indicates that the specific C_{30} – C_{34} diols are the main building blocks of the *Nannochloropsis* algaenan. Since their identification in Black Sea sediments (de Leeuw et al., 1981), C_{30} – C_{32} diols have shown to be ubiquitous in marine sediments in relatively high abundances (e.g., Smith et al., 1983; ten Haven et al., 1987a–c, 1992; Morris and Brassell, 1988; McCaffrey et al., 1991; Hoefs et al., 1995a), implying that their macromolecular counterparts can be important contributors to the OM of marine sediments.

The abundant alkene/alkane patterns noticed in the pyrolysates of the eustigmatophyte algaenans were also observed, though with different distributions, in the pyrolysates of algaenans of the marine chlorophytes, *C. spaerckii* and *Chlorococcum* sp. (Table 1). Because homologous series of alkenes and alkanes often dominate pyrolysates of marine kerogens isolated from both Recent (Klok et al., 1984; Eglinton et al., 1994) and ancient (Derenne et al., 1990; Douglas et al., 1991) sediments, it is concluded that a common source, which could be algaenans, although not necessarily derived from the four microalgae investigated here, are probably important contributors to marine sedimentary OM. Further support of this comes from stable carbon isotope studies of alkenes and alkanes isolated from pyrolysates of marine kerogens. Their $\delta^{13}\text{C}$ values are rather uniform and thus seem

to indicate a common source, i.e., algaenan (Eglinton, 1994; Collister et al., 1994; Hoefs et al., 1995b). However, very specific alkene/alkane distribution patterns, such as those found in the pyrolysates of the *Nannochloropsis* algaenans, have not been found so far in the pyrolysates of marine kerogens. This is readily explained by assuming that a suite of structurally diverse algaenans derived from different algae make up these kerogens. Although these different algaenans may be isotopically distinct, they all produce a full range of n -alkanes upon pyrolysis. This will lead to averaged carbon isotopic compositions for all n -alkanes which will not differ much from each other. These considerations indicate that the contribution of algaenan-producing microalgae to marine OM should be traced via specific biomarkers, i.e., diols in eustigmatophytes, rather than through specific pyrolysis patterns of their algaenans.

Aromatic and phenolic compounds were significantly present in the pyrolysates of the two chlorophyte algaenans already mentioned and were predominant in that of *Nannochloris* sp. (Table 1). The origin of the aromatic moieties observed in algaenan pyrolysates is still unclear, whereas the phenols are likely generated by the thermal cleavage of polyphenolic macromolecules such as phlorotannin (van Heemst et al., 1995).

Our results strongly indicate that specific marine eustigmatophyte and chlorophyte microalgae biosynthesize algaenans. These biomacromolecules, due to their highly resistant nature, are thought to escape from mineralization and become selectively preserved in marine sediments, as it has been observed previously for algaenans of freshwater microalgae in lacustrine environments (Goth et al., 1988; Tegehaar et al., 1989c; Derenne et al., 1991; Sinnighe Damsté et al., 1993). Therefore, we regard marine algaenans as a major sink of organic carbon in marine sediments. Upon burial in the Earth crust these materials may act as an important source of n -alkanes in marine petroleum.

Acknowledgments—We thank Prof. J. M. Hayes, Dr. T. I. Eglinton, and a number of anonymous referees for helpful comments on earlier drafts of this paper. This work was partly supported by a PIONIER grant to JSSD by the Netherlands Organisation for Scientific Research (NWO). This is NIOZ contribution 3015.

Editorial handling: J. D. Macdougall

REFERENCES

- Berkaloff C., Casadevall E., Largeau C., Metzger P., Peracca S., and Virlet J. (1983) The resistant polymer of the walls of the hydrocarbon-rich alga *Botryococcus braunii*. *Phytochemistry* **22**, 389–397.
- Boucher R. J., Standen G., Patience R. L., and Eglinton G. (1990) Molecular characterization of kerogen from Kimmeridge clay formation by mild selective chemical degradation and solid state ^{13}C -NMR. In *Advances in Organic Geochemistry 1989* (ed. B. Durand and F. Behar); *Org. Geochem.* **16**, 951–958.
- Carlson D. A., Roan C.-S., Yost R. A., and Hector J. (1989) Dimethyl disulfide derivatives of long chain alkenes, alkadienes and alkatrienes for gas chromatography/mass spectrometry. *Anal. Chem.* **61**, 1564–1571.
- Collinson M. E., van Bergen P. F., Scott A. C., and de Leeuw J. W. (1994) The oil-generating potential of plants from coal and coal-bearing strata through time: a review with new evidence from carboniferous plants. In *Coal and Coal-bearing Strata as Oil-*

- prone Source Rocks? (ed. A. C. Scott and A. J. Fleet); *Geol. Soc. Special Publ.* **77**, 31–70.
- Collister J. W., Lichtfouse E., Hieshima G., and Hayes J. M. (1994) Partial resolution of sources of *n*-alkanes in the saline portion of the Parachute Creek Member, Green River Formation (Piceance Creek Basin, Colorado). *Org. Geochem.* **21**, 645–659.
- de Leeuw J. W. and Largeau C. (1993) A review of macromolecular organic compounds that comprise living organisms and their role in kerogen, coal and petroleum formation. In *Organic Geochemistry, Principles and Applications* (ed. M. H. Engel and S. A. Macko), pp. 23–72. Plenum Press.
- de Leeuw J. W., Rijpstra W. I. C., and Schenck P. A. (1981) The occurrence and identifications of C₃₀, C₃₁ and C₃₂ alkan-1,15-diols and alkan-15-one-1-ols in Unit I and Unit II Black Sea sediments. *Geochim. Cosmochim. Acta* **45**, 2281–2285.
- de Leeuw J. W., van Bergen P. F., van Aarssen B. G. K., Gatellier J.-P. L. A., Sinninghe Damsté J. S., and Collinson M. E. (1991) Resistant biomacromolecules as major contributors to kerogen. *Phil. Trans. R. Soc. Lond. B* **333**, 329–337.
- Derenne S., Largeau C., Casadevall E., Sinninghe Damsté J. S., Tegelaar E. W., and de Leeuw J. W. (1990) Characterization of Estonian Kukersite by spectroscopy and pyrolysis: Evidence for abundant alkyl phenolic moieties in an Ordovician, marine, type II/I kerogen. In *Advances in Organic Geochemistry 1989* (ed. B. Durand and F. Behar); *Org. Geochem.* **16**, 873–888.
- Derenne S., Largeau C., Casadevall E., Berkaloff C., and Rousseau B. (1991) Chemical evidence of kerogen formation in source rocks and oil shales via selective preservation of thin resistant outer walls of microalgae: origin of ultralaminae. *Geochim. Cosmochim. Acta* **55**, 1041–1050.
- Derenne S., Largeau C., Berkaloff C., Rousseau B., Wilhelm C., and Hatcher P. G. (1992a) Non-hydrolysable macromolecular constituents from outer walls of *Chlorella fusca* and *Nanochlorum eucaryotum*. *Phytochemistry* **31**, 1923–1929.
- Derenne S., Largeau C., and Hatcher P. G. (1992b) Structure of *Chlorella fusca* algaenan. Relationships with ultralaminae in lacustrine kerogens. Species-environment-dependant variations in the composition of fossil ultralaminae. *Org. Geochem.* **18**, 417–422.
- Derenne S., Le Berre F., Largeau C., Hatcher P., Connan J., and Raynaud J. F. (1992c) Formation of ultralaminae in marine kerogens via selective preservation of thin resistant outer walls of microalgae. In *Advances in Organic Geochemistry 1991* (ed. C. Eckhart and J. R. Maxwell); *Org. Geochem.* **19**, 345–350.
- Douglas A. G., Sinninghe Damsté J. S., Fowler M. G., Eglinton T. I., and de Leeuw J. W. (1991) Unique distributions of hydrocarbons and sulphur compounds released by flash pyrolysis from the fossilised alga *Gloeocapsomorpha prisca*, a major constituent in one of four Ordovician kerogens. *Geochim. Cosmochim. Acta* **55**, 275–291.
- Eglinton T. I. (1994) Carbon isotopic evidence for the origin of macromolecular aliphatic structures in kerogen. *Org. Geochem.* **21**, 721–735.
- Eglinton T. I., Irvine J. E., Vairavamurthy A., Zhou W., and Manowitz B. (1994) Formation and diagenesis of macromolecular organic sulphur in Peru margin sediments. In *Advances in Organic Geochemistry 1993* (ed. N. Telnæs, et al.); *Org. Geochem.* **22**, 781–799.
- Gatellier J.-P. L. A., de Leeuw J. W., Sinninghe Damsté J. S., Derenne S., Largeau C., and Metzger P. (1993) Very early diagenesis of a resistant cell wall biomacromolecule of the green alga *Botryococcus braunii* (Race A) as revealed by spectroscopic and pyrolysis investigations of a Coorongite. *Geochim. Cosmochim. Acta* **57**, 2053–2068.
- Gelin F., Gatellier J.-P. L. A., Sinninghe Damsté J. S., Derenne S., Largeau C., Metzger P., and de Leeuw J. W. (1993) Mechanisms of flash pyrolysis of ether lipids isolated from the green microalga *Botryococcus braunii*. *J. Anal. Appl. Pyrolysis* **27**, 155–168.
- Gelin F., Sinninghe Damsté J. S., Derenne S., Largeau C., Metzger P., and de Leeuw J. W. (1994) Scope and limitations of flash pyrolysis-gas chromatography/mass spectrometry as revealed by the thermal behaviour of high-molecular-weight lipids derived from the green microalga *Botryococcus braunii*. *J. Anal. Appl. Pyrolysis* **28**, 183–204.
- Goth K., de Leeuw J. W., Püttman W., and Tegelaar E. W. (1988) Origin of Messel Oil Shale kerogen. *Nature* **336**, 759–761.
- Guilford W. J., Schneider D. M., Labovitz J. and Opella S. J. (1988) High resolution solid state ¹³C NMR spectroscopy of sporopollenin from different plant taxa. *Plant. Physiol.* **86**, 134–136.
- Hatcher P. G., Spiker E. C., Szeverenyi N. M., and Maciel G. E. (1983) Selective preservation and origin of petroleum-forming aquatic kerogen. *Nature* **305**, 498–501.
- Hemsley A. R., Barrie P. J., Chaloner W. G., and Scott A. C. (1993) The composition of sporopollenin: its contribution to living and fossil spore systematics. *Grana* (suppl. 1), 2–11.
- Hoefs M. J. L., Sinninghe Damsté J. S., and de Leeuw J. W. (1995a) Organic geochemistry of Arabian Sea surface sediments: paleoenvironmental implications. In *Organic Geochemistry: Developments and Applications to Energy, Climate, Environment and Human History* (ed. J. O. Grimalt and C. Dorronsoro), pp. 158–160.
- Hoefs M. J. L., van Heemst J. D. H., Gelin F., Koopmans M. P., de Leeuw J. W., and Sinninghe Damsté J. S. (1995b) Alternative biological sources for 1,2,3,4-tetramethylbenzene in flash pyrolysates of kerogen. *Org. Geochem.* (in press).
- Klok J., Baas M., Cox H. C., de Leeuw J. W., Rijpstra W. I. C., and Schenck P. A. (1984) Qualitative and quantitative characterization of the total organic matter in a recent sediment (Part II). In *Advances in Organic Geochemistry 1983* (ed. P. A. Schenck et al.); *Org. Geochem.* **6**, 265–278.
- McCaffrey M. A., Farrington J. W., and Repeta D. J. (1991) The organic geochemistry of Peru margin surface sediments: II. Paleoenvironmental implications of hydrocarbon and alcohol profiles. *Geochim. Cosmochim. Acta* **55**, 483–498.
- Morris R. J. and Brassell S. C. (1988) Long-chain alkanediols: biological markers for cyanobacterial contributions to sediments. *Lipids* **23**, 256–258.
- Nip M., Tegelaar E. W., Brinkhuis H., de Leeuw J. W., Schenck P. A., and Holloway P. J. (1986) Analysis of modern and fossil plant cuticles by Curie point Py-GC and Curie point Py-GC-MS: Recognition of a new, highly aliphatic and resistant biopolymer. In *Advances in Organic Geochemistry 1985* (ed. D. Leythaeuser and J. Rullkötter); *Org. Geochem.* **10**, 769–778.
- Nip M., de Leeuw J. W., Schenck P. A., Windig W., Meuzelaar H. L. C., and Crelling J. C. (1989) A flash pyrolysis and petrographic study of cutinite from the Indiana paper coal. *Geochim. Cosmochim. Acta* **53**, 671–683.
- Panganamala R. V., Sievert C. F., and Cornwell D. G. (1971) Quantitative estimation and identification of O-glycerol as alkyl iodides and their hydrocarbon derivatives. *Chem. Phys. Lipids* **7**, 336–344.
- Philp R. P. and Calvin M. (1976) Possible origin for insoluble organic (kerogen) debris in sediments from insoluble cell-wall materials of algae and bacteria. *Nature* **262**, 134–136.
- Sinninghe Damsté J. S., De Las Heras F. X. C., van Bergen P. F., and de Leeuw J. W. (1993) Characterization of Tertiary Catalan lacustrine oil shales: Discovery of extremely organic sulphur-rich Type I kerogens. *Geochim. Cosmochim. Acta* **57**, 389–415.
- Smith D. J., Eglinton G., and Morris R. J. (1983) Occurrence of long-chain alkan-diols and alkan-15-one-1-ols in a Quaternary Sapropel from the Eastern Mediterranean. *Lipids* **18**, 902–905.
- Tegelaar E. W. et al. (1989a) Scope and limitations of several pyrolysis methods in the structural elucidation of a macromolecular plant constituent in the leaf cuticle of *Agave Americana* L. *J. Anal. Appl. Pyrolysis* **15**, 29–54.
- Tegelaar E. W., Mattheizing R. M., Jansen J. B. H., Horsfield B., and de Leeuw J. W. (1989b) Possible origin of *n*-alkanes in high-wax crude oils. *Nature* **342**, 529–531.
- Tegelaar E. W., de Leeuw J. W., Derenne S., and Largeau C. (1989c) A reappraisal of kerogen formation. *Geochim. Cosmochim. Acta* **53**, 3103–3106.
- ten Haven H. L., Baas M., de Leeuw J. W., and Schenck P. A. (1987a) Late Quaternary Mediterranean sapropels. I—On the origin of organic matter in sapropel S₇. *Mar. Geol.* **75**, 137–156.
- ten Haven H. L., Baas M., de Leeuw J. W., Schenck P. A., and Brinkhuis H. (1987b) Late Quaternary Mediterranean sapropels

- II. Organic geochemistry and palynology of S₁ sapropels and associated sediments. *Chem. Geol.* **64**, 149–167.
- ten Haven H. L., Baas M., Kroot M., de Leeuw J. W., Schenck P. A., and Ebbing J. (1987c) Late Quaternary Mediterranean sapropels: III. Assessment of source of input and palaeotemperature as derived from biological markers. *Geochim. Cosmochim. Acta* **51**, 803–810.
- ten Haven H. L., Eglinton G., Farrimond P., Kohnen M. E. L., Poynter J. G., Rullkötter J., and Welte D. H. (1992) Variations in the content and composition of organic matter in sediments underlying active upwelling regimes: a study from ODP Legs 108, 112, and 117. In *Upwelling Systems: Evolution Since the Early Miocene* (ed. C. P. Summerhayes et al.); *Geol. Soc. Special Publ. No. 64*, pp. 229–246.
- van Aarssen B. G. K., Cox H. C., Hoogendoorn P., and de Leeuw J. W. (1990) A cadinene biopolymer in fossil and extant dammar resins as a source for cadinanes and bicadinanes in crude oils from South East Asia. *Geochim. Cosmochim. Acta* **54**, 3021–3031.
- van Bergen P. F., Collinson M. E., Sinninghe Damsté J. S., and de Leeuw J. W. (1994) Chemical and microscopical characterization of inner seed coats of fossil water plants. *Geochim. Cosmochim. Acta* **58**, 231–239.
- van Heemst J. D. H., Peulvé S., de Leeuw J. W., Sicre M.-A., and Saliot A. (1995) Algal polyphenolic resistant macromolecules in marine dissolved and particulate organic matter. In *Organic Geochemistry: Developments and Applications to Energy, Climate, Environment and Human History* (ed. J. O. Grimalt and C. Dorronsoro); pp. 940–942.
- Veldhuis M. J. W. and Admiraal W. (1987) The influence of phosphate depletion on the growth and colony formation of *Phaeocystis pouchetii* (Hariot) Lagerheim. *Mar. Biol.* **95**, 47–54.
- Volkman J. K., Barrett S. M., Dunstan G. A., and Jeffrey S. W. (1992) C₃₀–C₃₂ alkyl diols and unsaturated alcohols in microalgae of the class Eustigmatophyceae. *Org. Geochem.* **18**, 131–138.
- Zeliber J. L., Romankiw L., Hatcher P. G., and Colwell R. R. (1988) Comparative analysis of the chemical composition of mixed and pure cultures of green algae and their decomposed residues by ¹³C nuclear magnetic resonance spectroscopy. *Appl. Environ. Microbiol.* **54**, 1051–1060.